

The terrestrial Arctic response to (and role in) local and global climate change



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Laboratory
Boulder, CO**



NCAR is sponsored by the National Science Foundation



NCAR Sea Ice Prediction Pool Results 2008-2010

Organized by Jennifer E. Kay
(jenkay@ucar.edu)

Predictions for mean September ice extent submitted
on June 1 (~20 entries each year)

Owner of best guess gets ice cream from worst guess

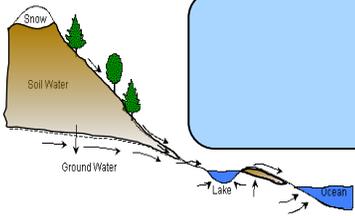
NCAR ensemble mean prediction error

2008: 0.00 (-0.11) million km²

2009: -0.63 (-0.73)

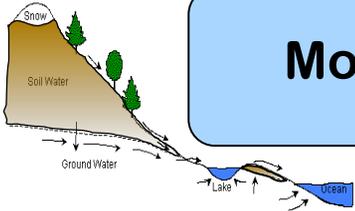
2010: -0.01 (-0.11)



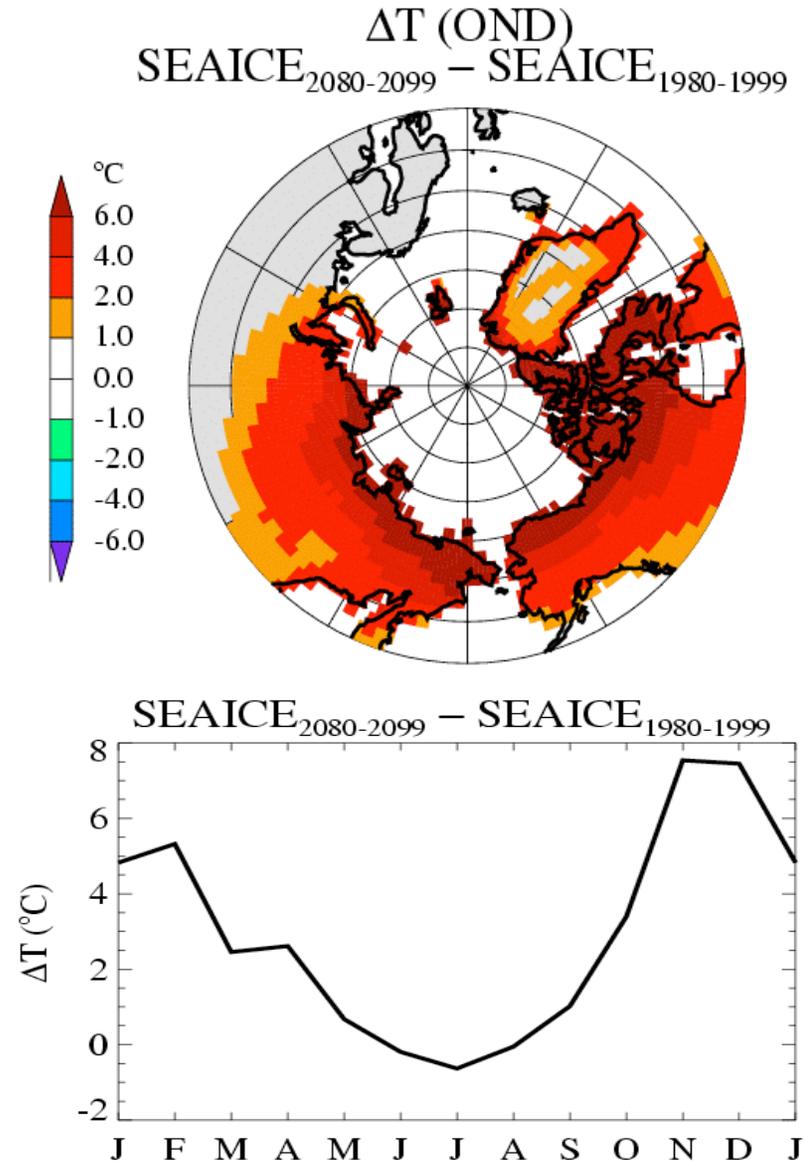
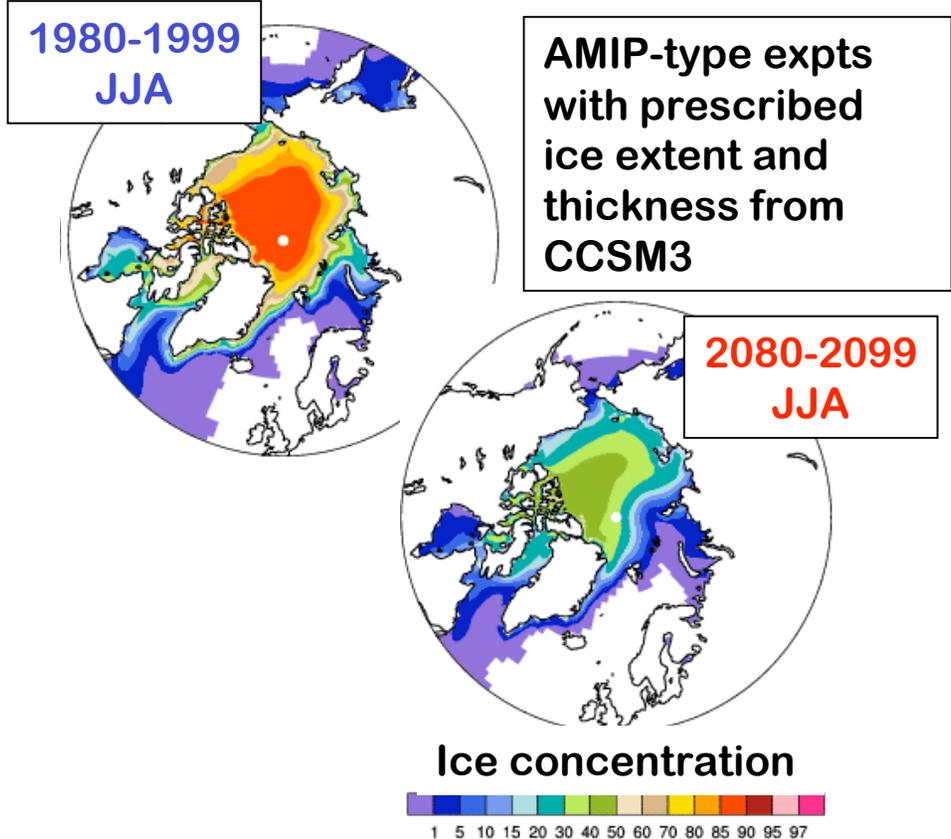


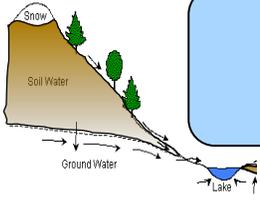
Topics for today

- **Impact of sea ice loss on terrestrial Arctic climate**
 - T, P, Snow
 - Rate of ice loss
 - Compare impact of sea ice loss to terrestrial snow extent decrease
- **Terrestrial Arctic feedbacks**
 - Example of shrub cover expansion



Modeled impact of sea ice loss on terrestrial Arctic climate

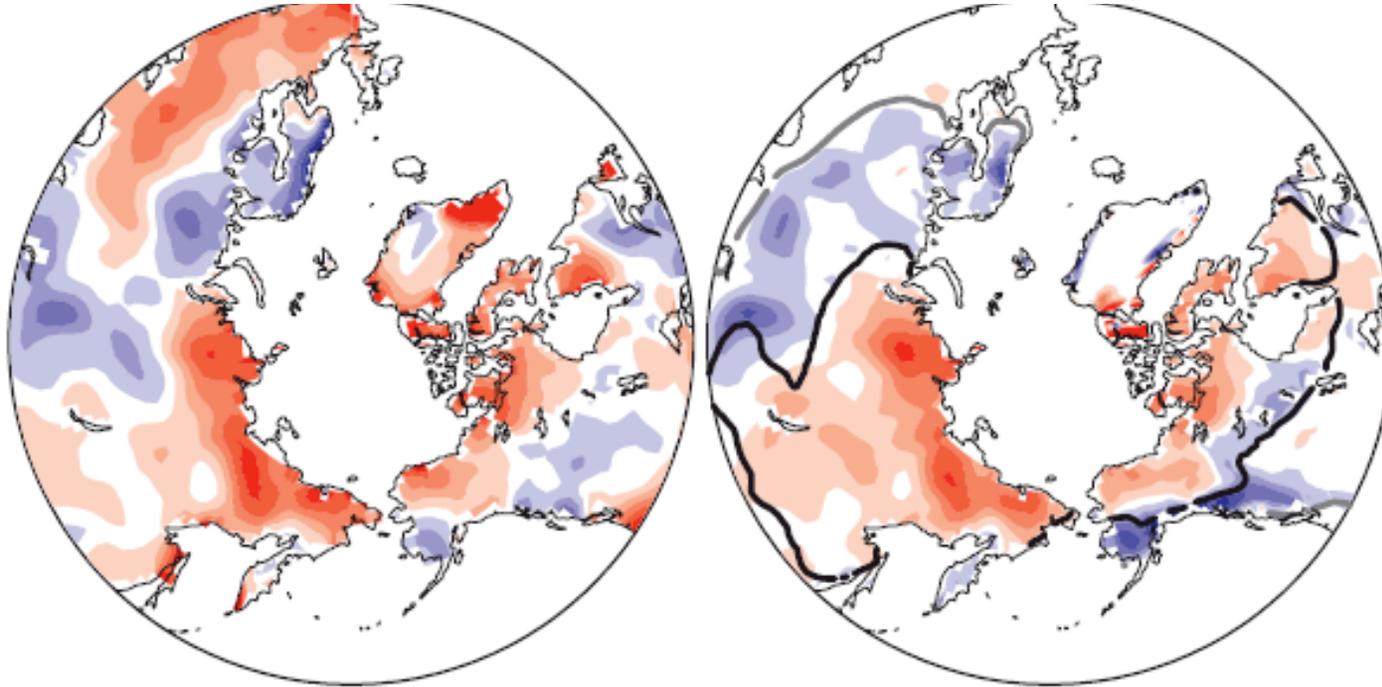


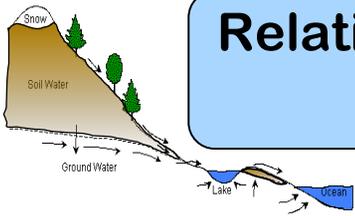


Modeled impact of sea ice loss on terrestrial Arctic climate

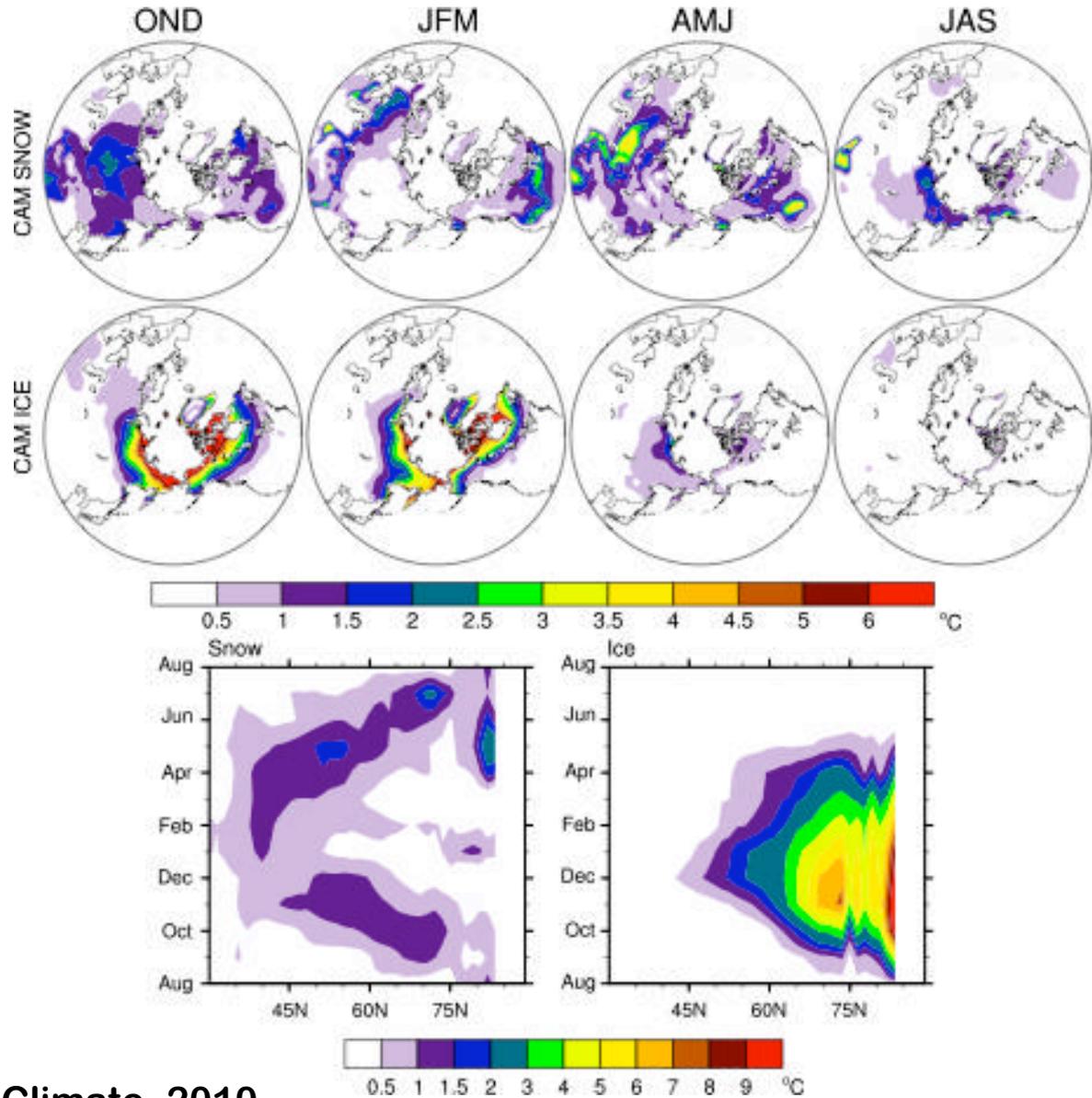
Change in accumulated Oct to Mar precip

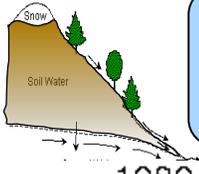
Change in March snow depth





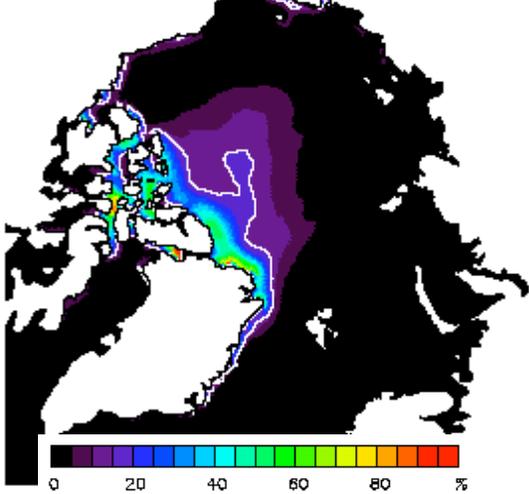
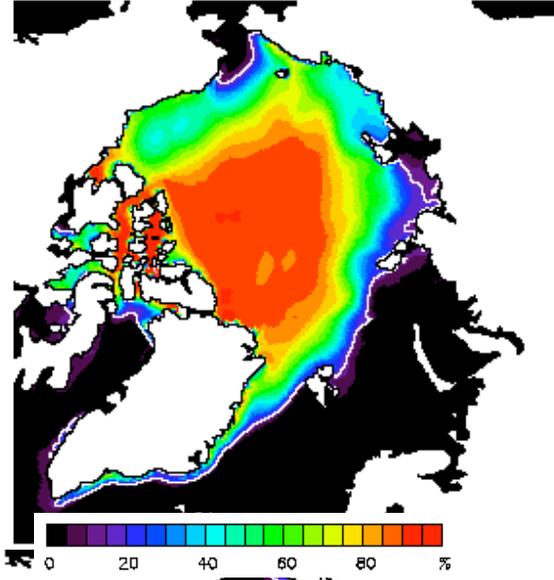
Relative influence of projected sea ice vs snow cover change on Arctic terrestrial T_{air}



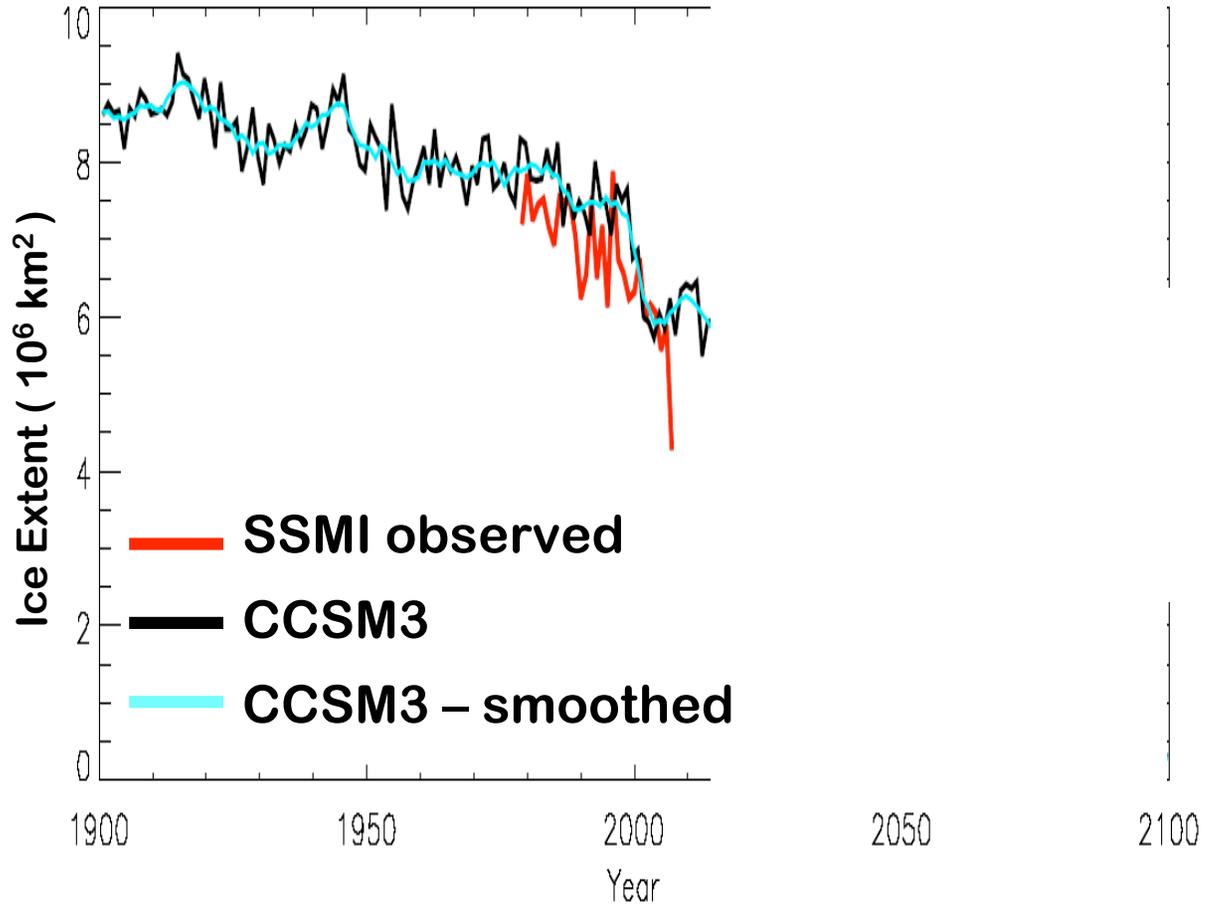


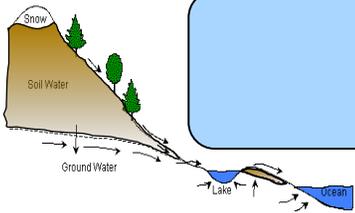
Abrupt reductions in the September sea ice cover CCSM3

1990–1999 Avg SEPT aice

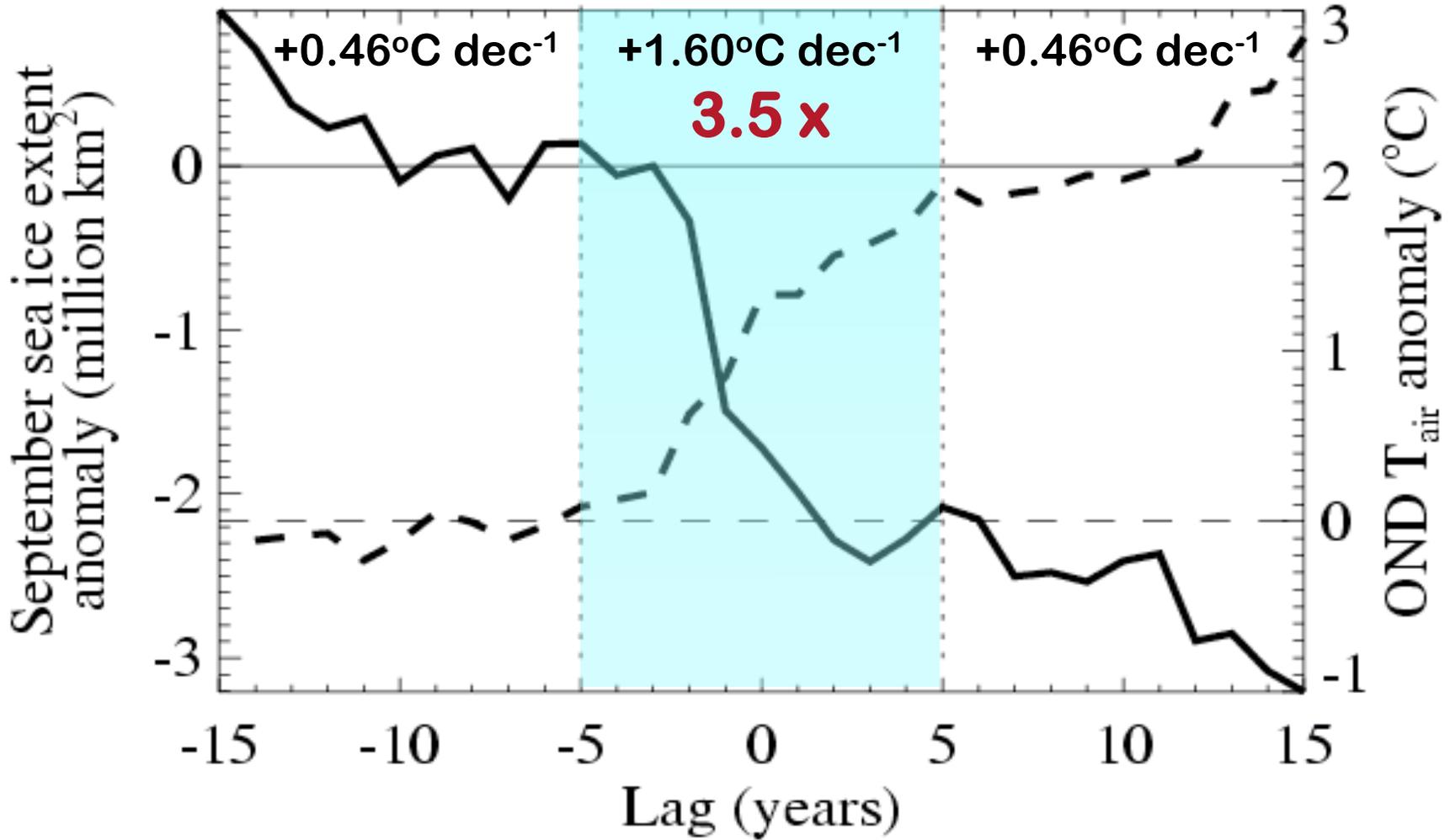


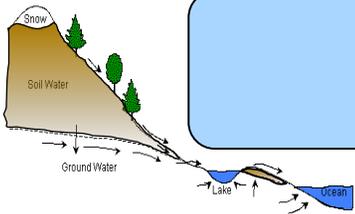
September sea ice extent



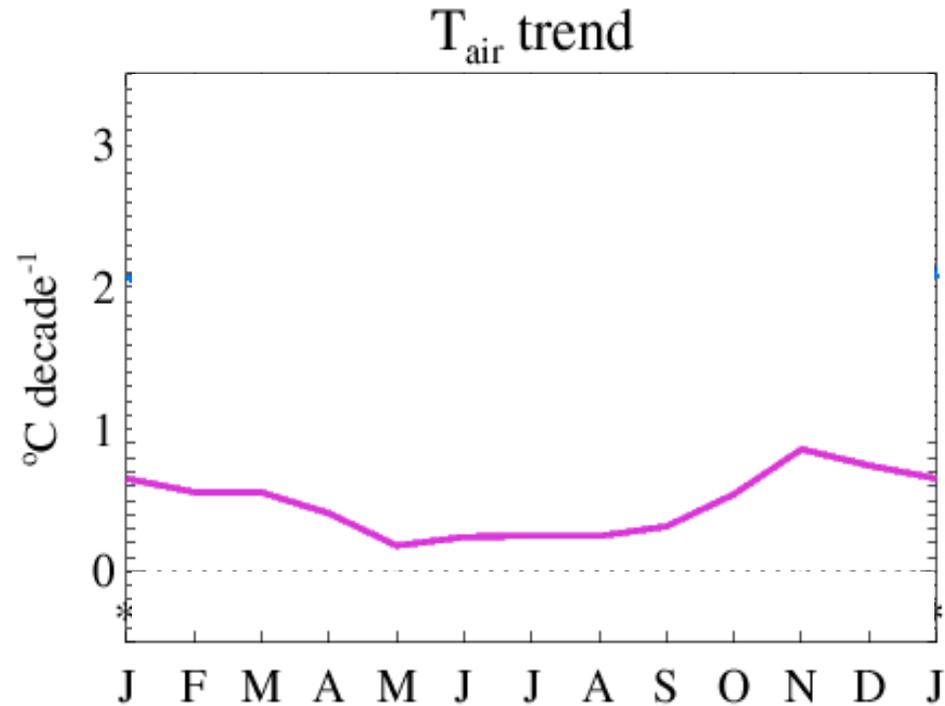
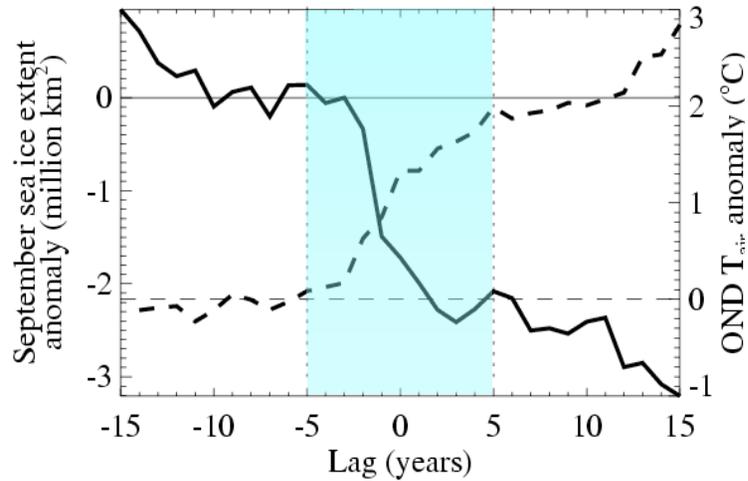


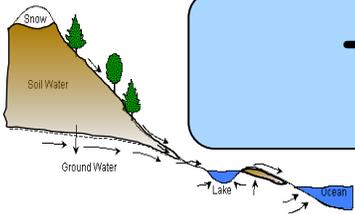
**Lagged composite:
sea ice extent and western Arctic land T_{air}**





T_{air} trend over Arctic land during rapid and moderate sea ice loss



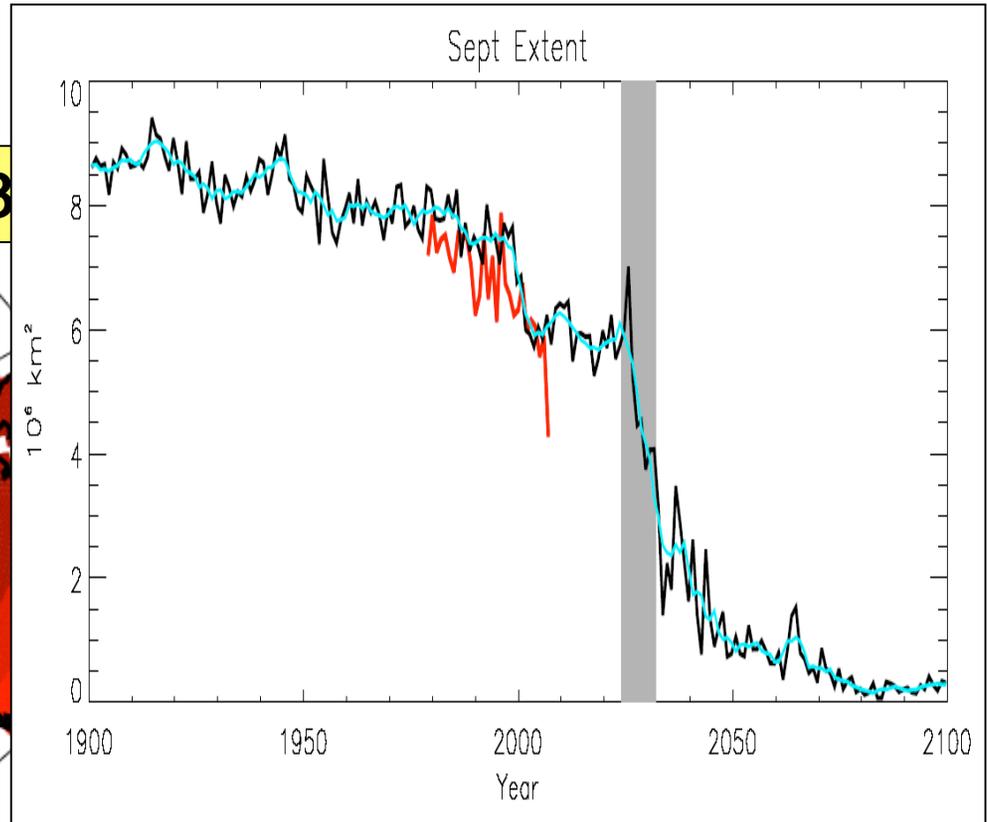
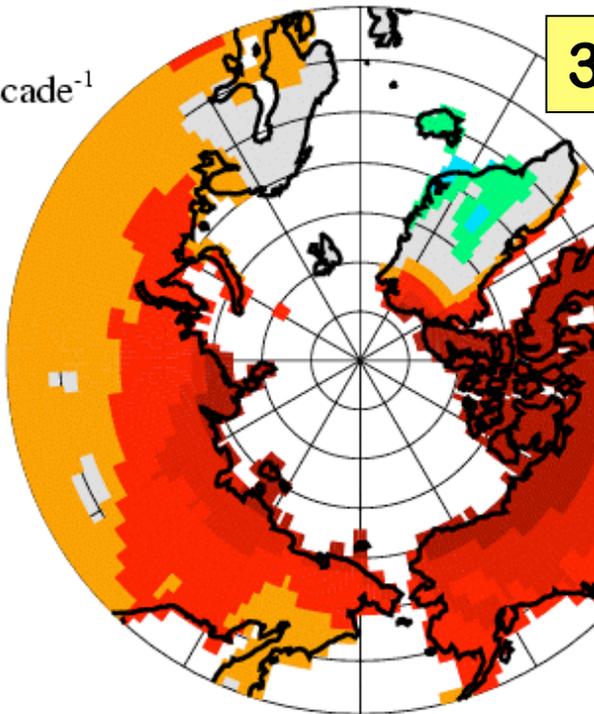
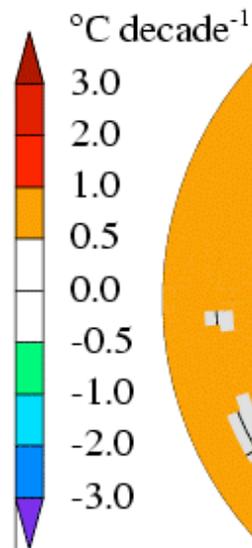


Trends during rapid vs moderate sea ice loss

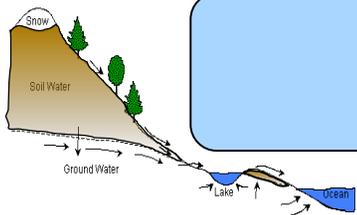
Periods of rapid sea ice loss

OND

T

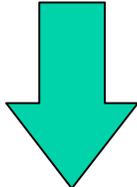
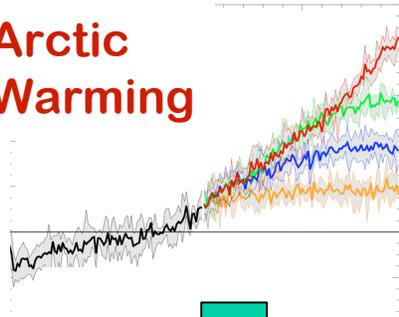


- Similar change in trends in q_{air} , $LW \downarrow$
- No statistically significant or spatially coherent change in trends in P , snow depth, or $SW \downarrow$

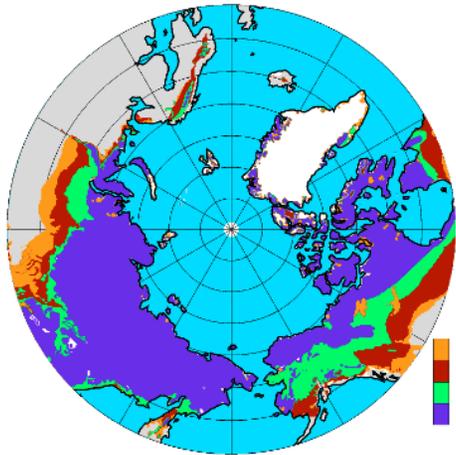


Permafrost thaw

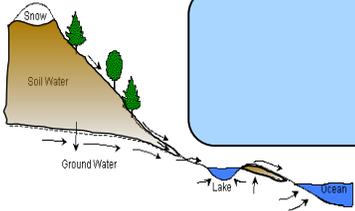
Arctic Warming



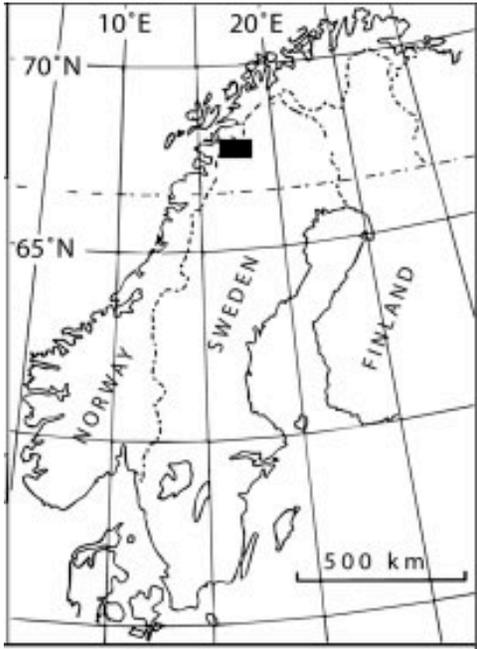
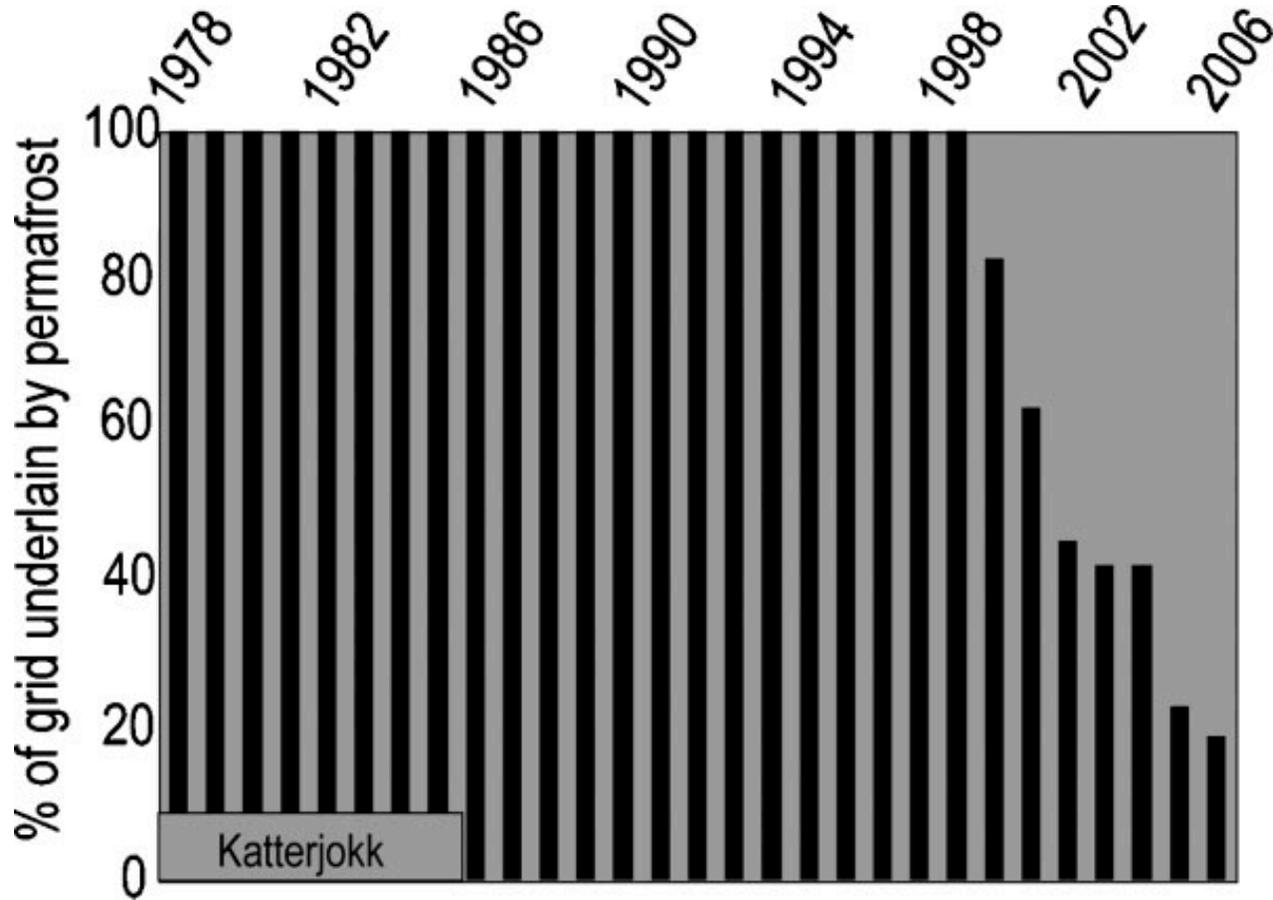
Permafrost warms and thaws



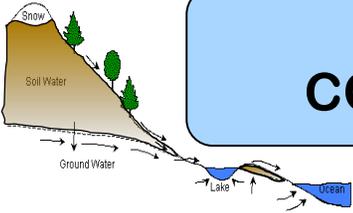
- Cont.
- Disc.
- Spor.
- Isol.



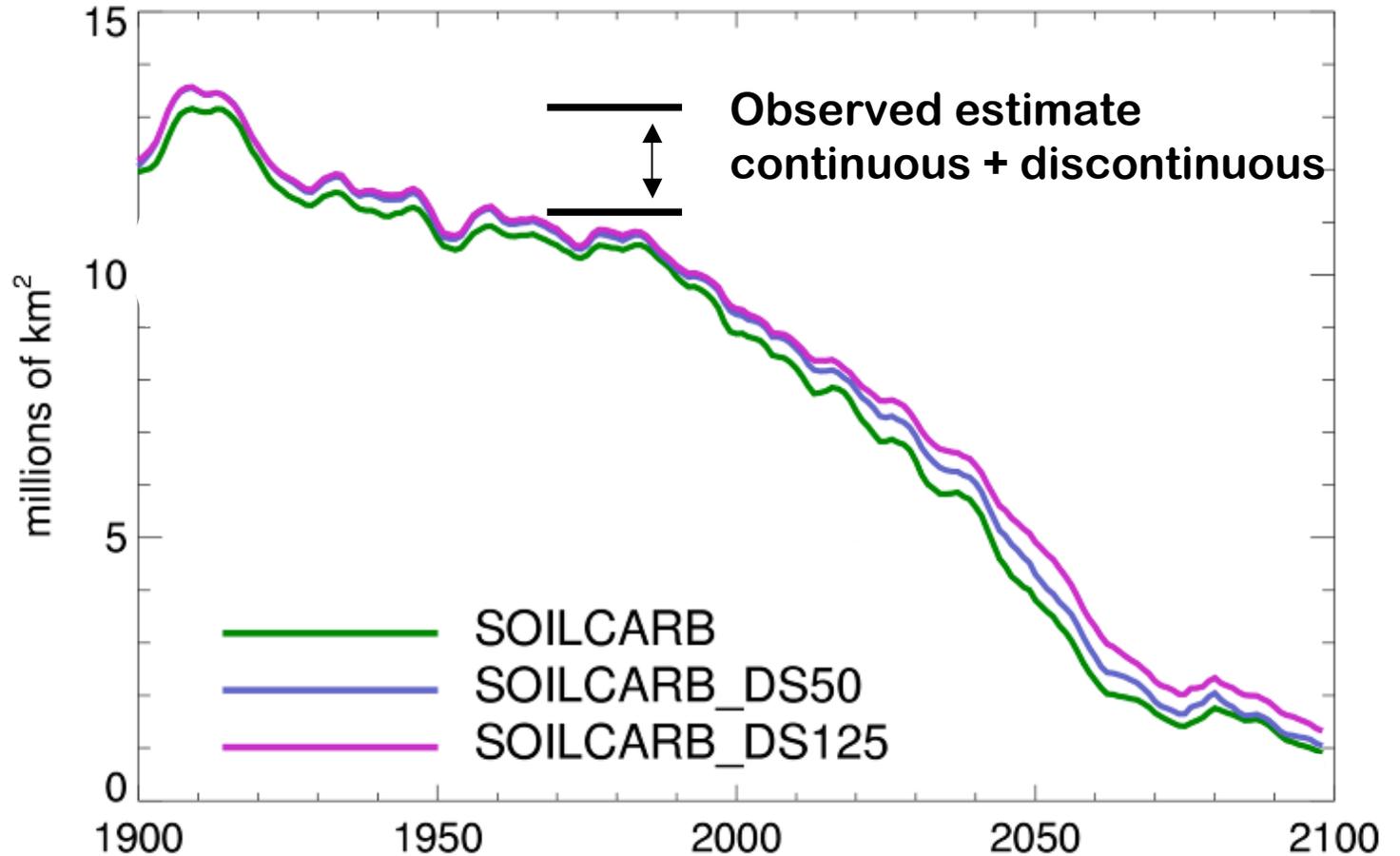
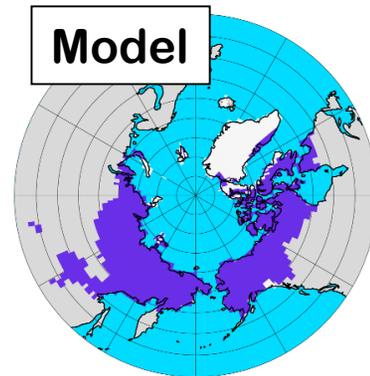
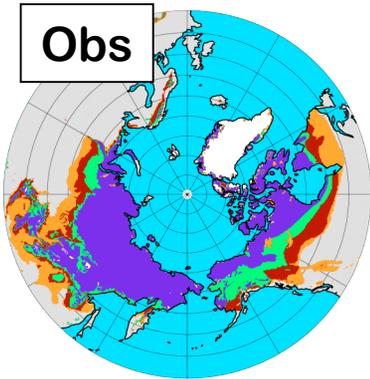
Observed permafrost degradation



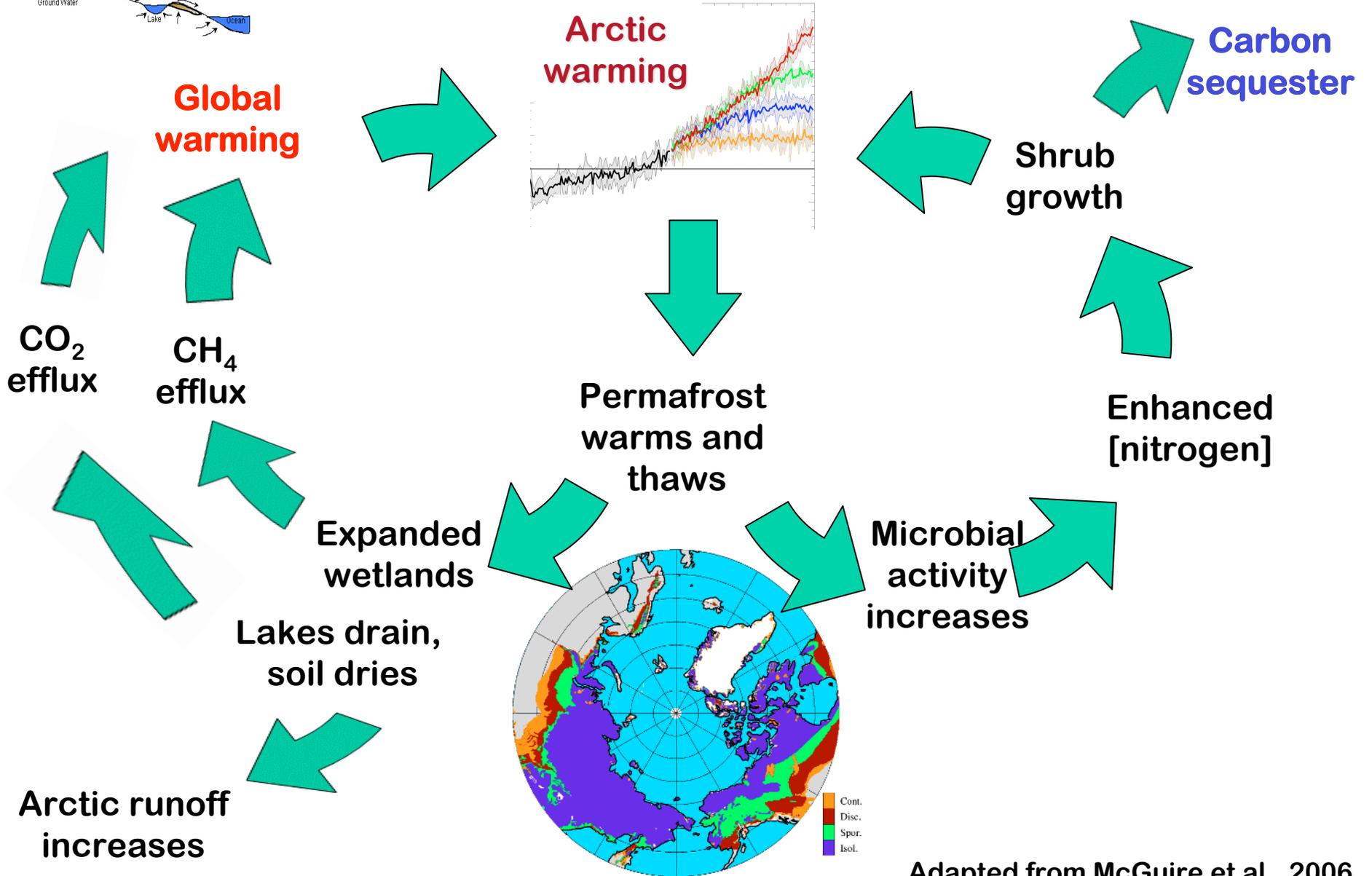
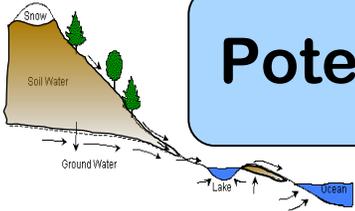
Projected near-surface permafrost degradation CCSM3 20C+A1B forcing, land as in Lawrence et al. 2008



Integrated area with permafrost within 3m of surface

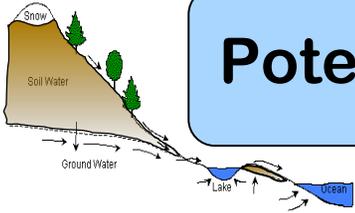


Potential Arctic terrestrial climate-change feedbacks

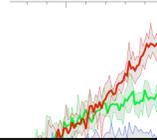


Adapted from McGuire et al., 2006

Potential Arctic terrestrial climate-change feedbacks



Arctic warming



Carbon sequester

Global warming

CO₂ efflux

CH₄ efflux

Expanded wetlands

Lakes drain, soil dries

Arctic runoff increases

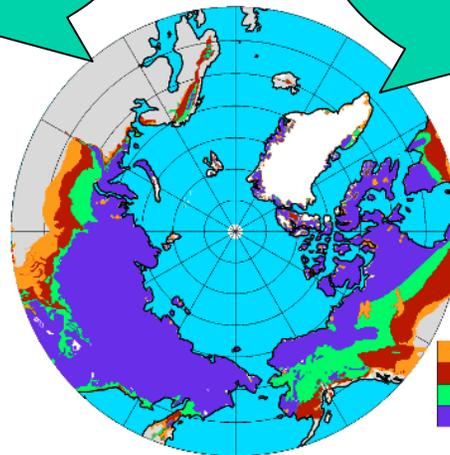
Hydrology

Western Siberia lake area:
 +12% in Continuous
 -6 to -9% in Disc., Isol., Spor.
 (Smith et al., 2005)

Increases in runoff and enhanced winter baseflow
 (Dai et al., 2009; St Jacques and Sauchen, 2009)

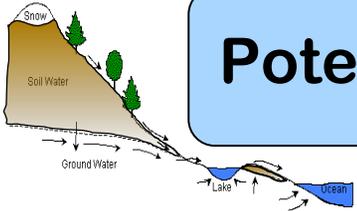
Permafrost

Microbial activity increases



Cont.
Disc.
Spor.
Isol.

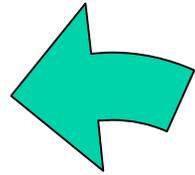
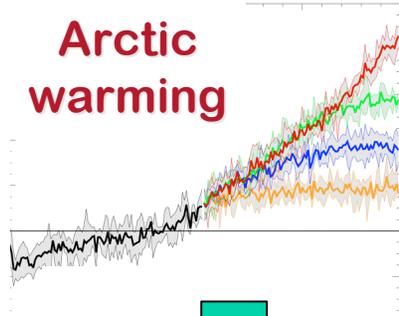
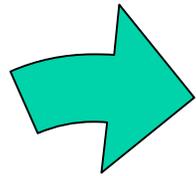
Potential Arctic terrestrial climate-change feedbacks



Global warming

CO₂ efflux

CH₄ efflux



Carbon sequester

Shrub growth

Carbon stocks in permafrost-affected soil

~ 1700 PgC (Tarnocai et al., 2009)

Atmos carbon content

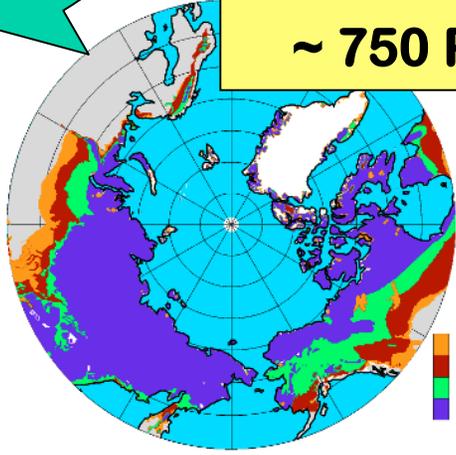
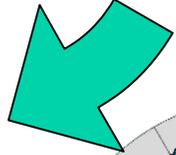
~ 750 PgC + ~9 PgC yr⁻¹

increases

Expanded wetlands

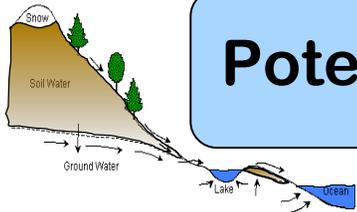
Lakes drain, soil dries

Arctic runoff increases



Adapted from McGuire et al., 2006

Potential Arctic terrestrial climate-change feedbacks



Carbon sequester

Global warming

Surface energy partitioning
Soil moisture and permafrost state affects partitioning of net radiation into sensible, latent, and ground heat fluxes

CO₂ efflux

CH₄ efflux

Permafrost warms and thaws

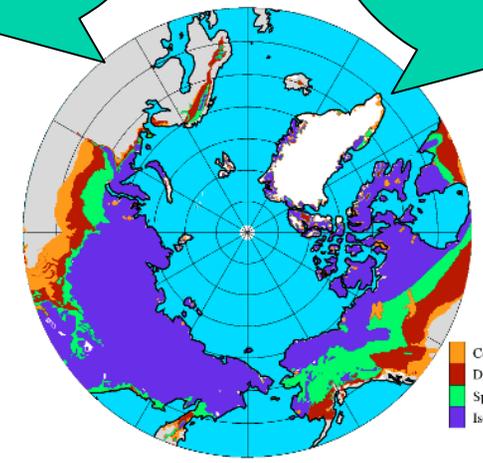
Enhanced [nitrogen]

Expanded wetlands

Microbial activity increases

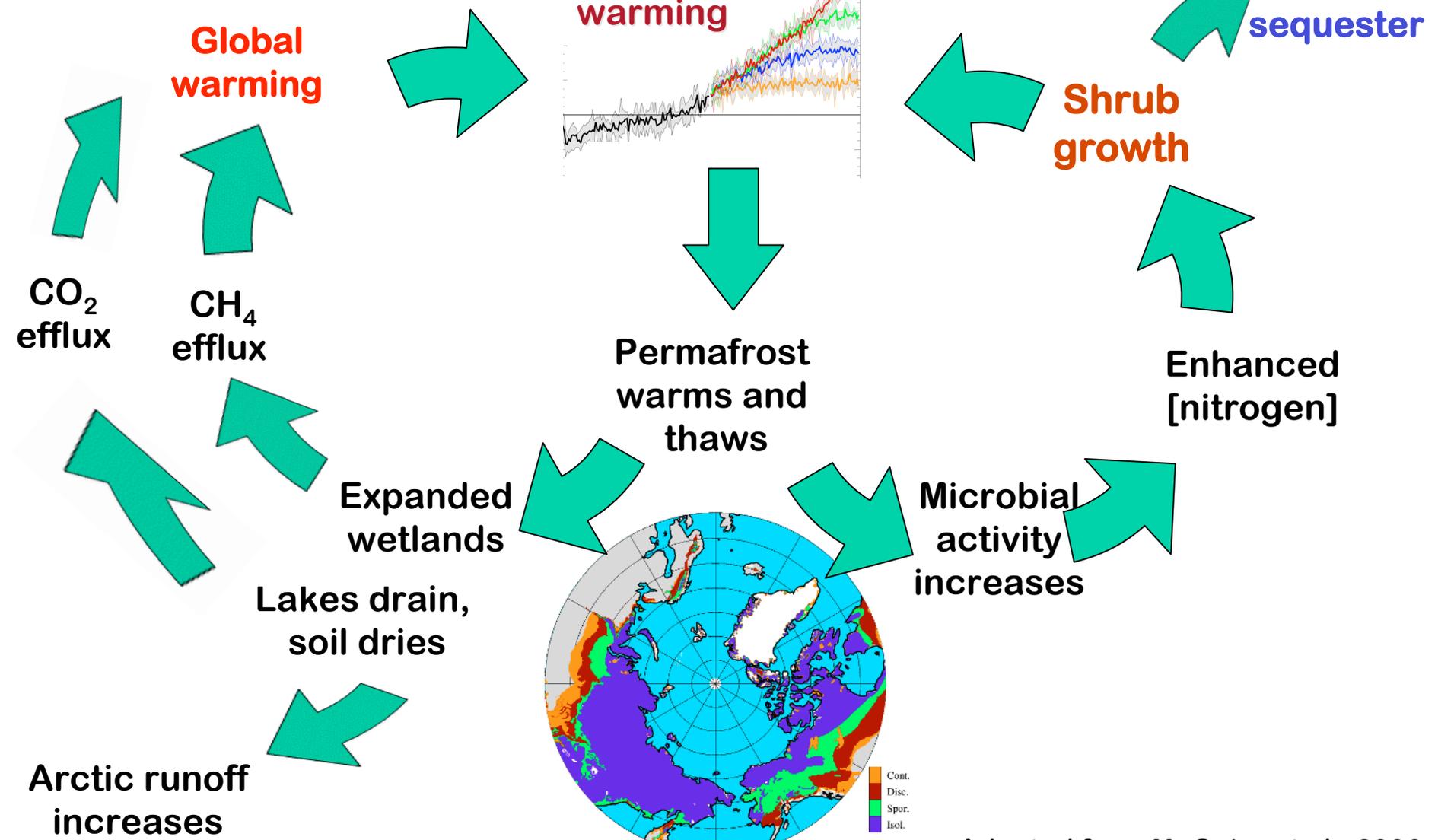
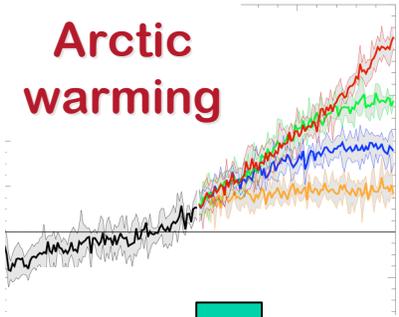
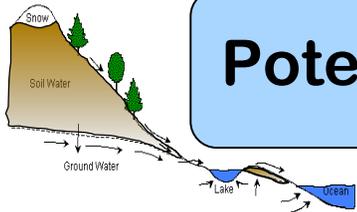
Lakes drain, soil dries

Arctic runoff increases

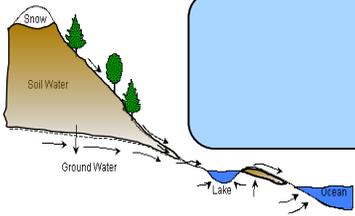


Adapted from McGuire et al., 2006

Potential Arctic terrestrial climate-change feedbacks



Adapted from McGuire et al., 2006



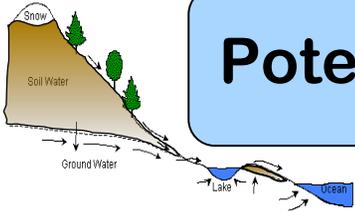
Shrub cover increasing in Arctic



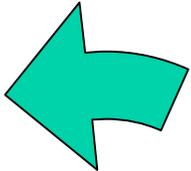
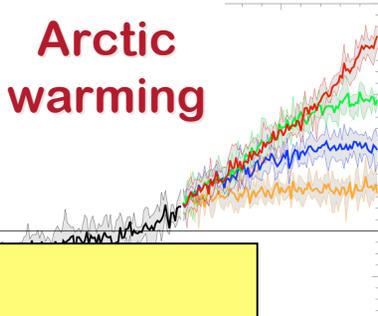
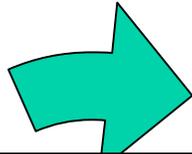
- Shrub cover increasing in N. Alaska at 1.2% per decade since 1950, 14% to 20% cover (Sturm et al. 2001)
- Shrub cover can increase much quicker in response to climate shifts than forest cover since shrubs already exist across most of the Arctic

Figure 1. Increasing abundance of shrubs in arctic Alaska. The photographs were taken in 1948 and 2002 at identical locations on the Colville River ($68^{\circ} 57.9'$ north, $155^{\circ} 47.4'$ west). Dark objects are individual shrubs 1 to 2 meters high and several meters in diameter. Similar changes have been detected at more than 200 other locations across arctic Alaska where comparative photographs are available. Photographs: (1948) US Navy, (2002) Ken Tape.

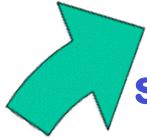
Potential Arctic terrestrial climate-change feedbacks



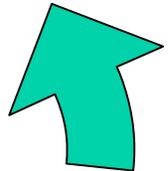
Global warming



Shrub growth



Carbon sequester



Enhanced [nitrogen]

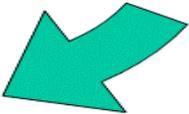
Vegetation
Radiative forcing of complete conversion tundra to shrubland
+8.9W m⁻² (4.2W m⁻² GHG)
 (Chapin et al., 2005)
Since 1950, 14% to 20% cover
 (Sturm et al., 2005)

Co
eff

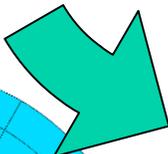
ost
nd

wetlands

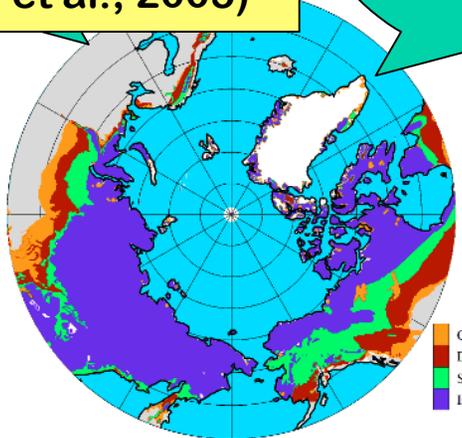
**Lakes drain,
soil dries**



**Arctic runoff
increases**



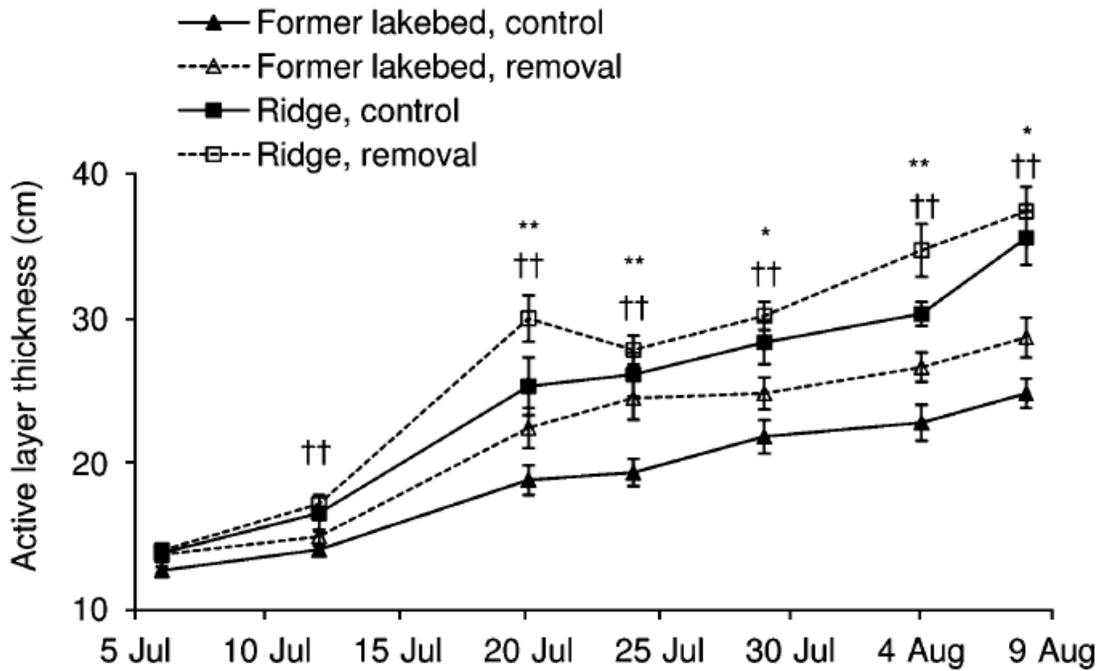
**Microbial
activity
increases**



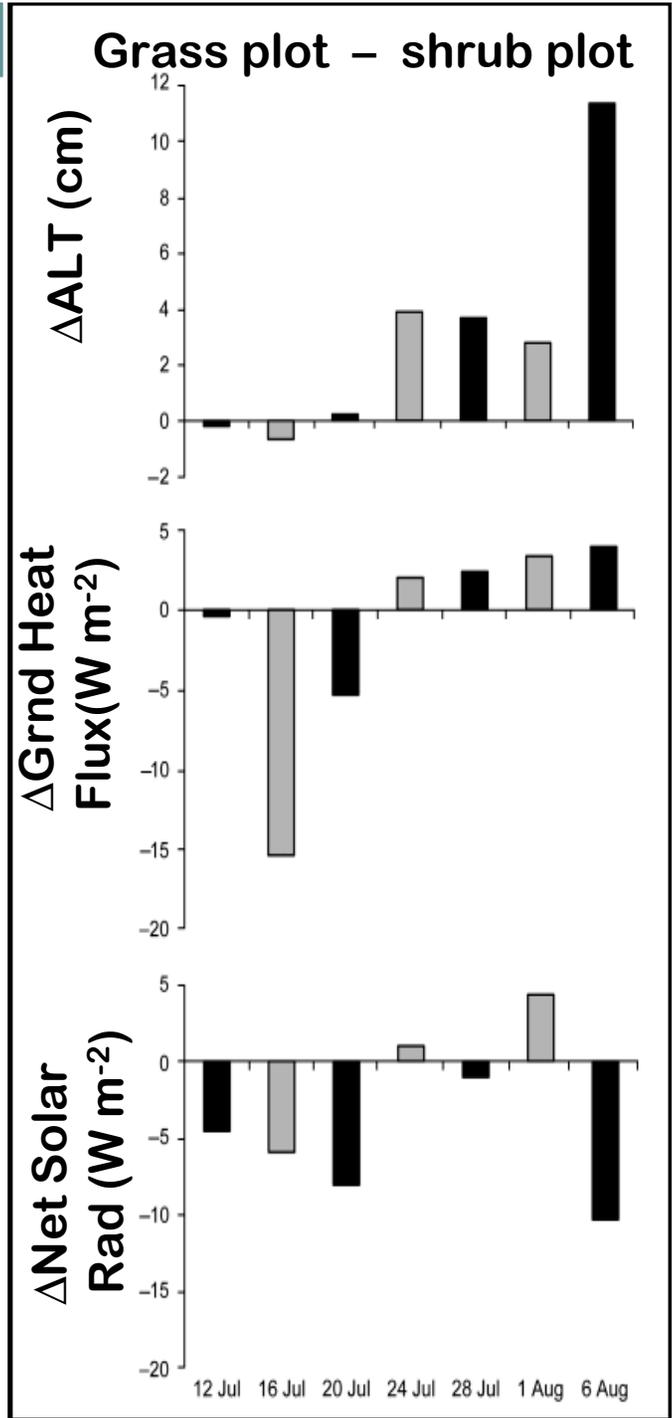
Adapted from McGuire et al., 2006

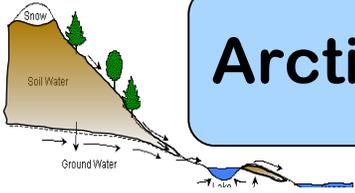
Shrub expansion may reduce summer permafrost thaw in Siberian tundra

D. BLOK*, M. M. P. D. HEIJMANS*, G. SCHAEPMAN-STRUB*†, A. V. KONONOV‡, T. C. MAXIMOV‡ and F. BERENDSE*



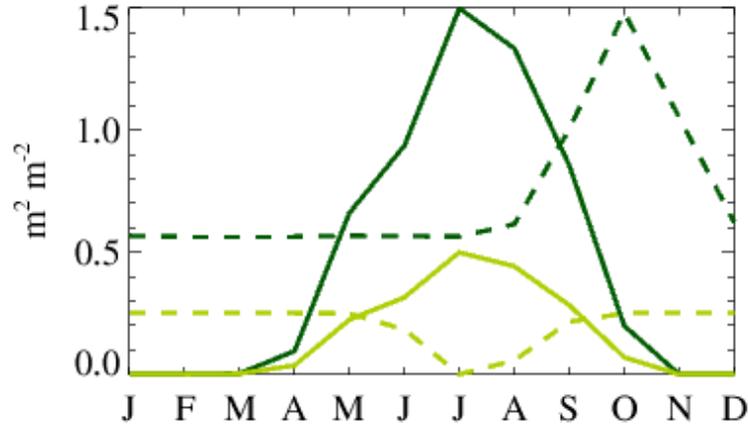
“These results suggest that the expected expansion of deciduous shrubs in the Arctic region, triggered by climate warming, may reduce summer permafrost thaw.”





Arctic tundra vegetation experiments with CLM/CESM

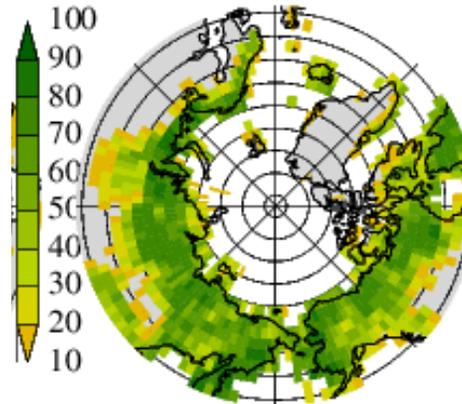
Leaf Area / Stem Area Index



- Grass (LAI)
- Shrub (LAI)
- - - Grass (SAI)
- - - Shrub (SAI)

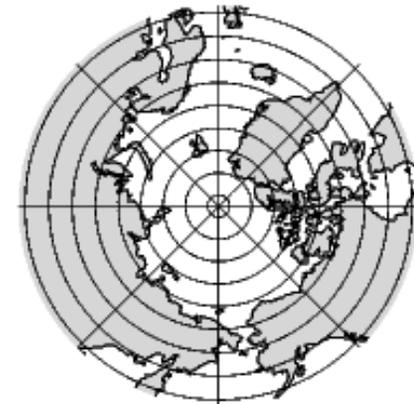
GRASS

Arctic C3 Grass

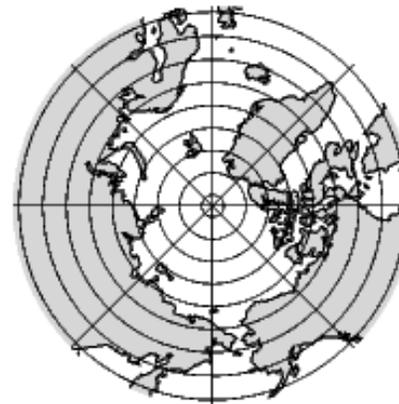


SHRUB

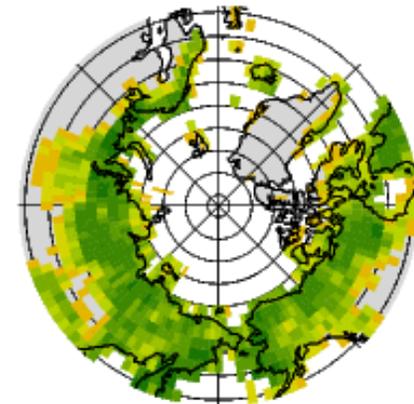
Arctic C3 Grass

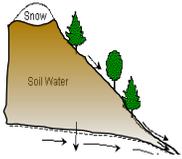


Boreal Shrub

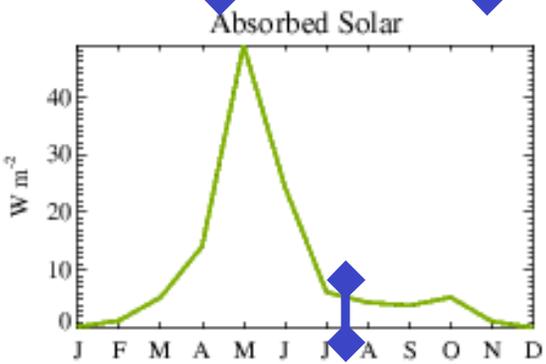
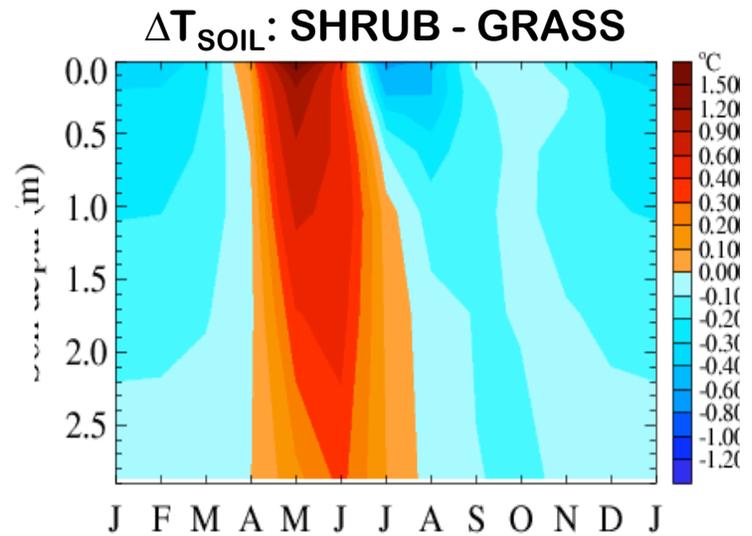
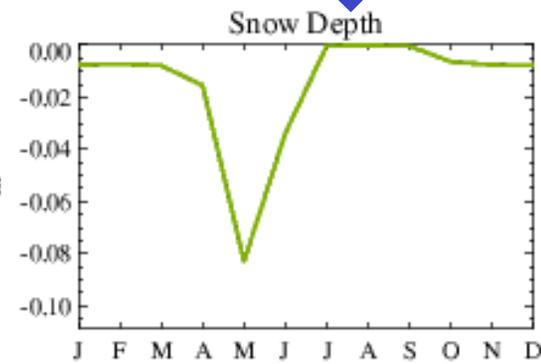
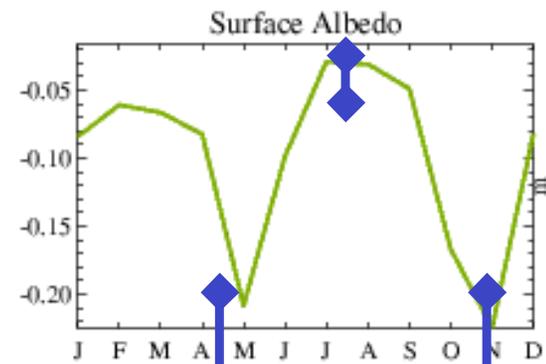
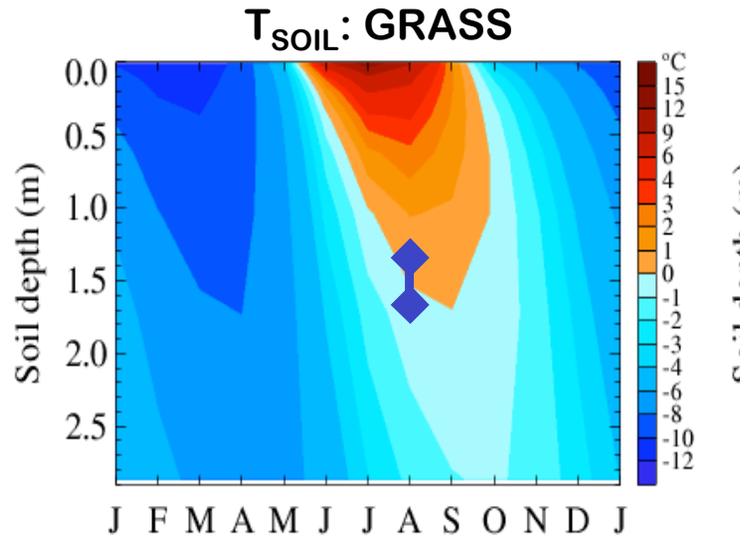
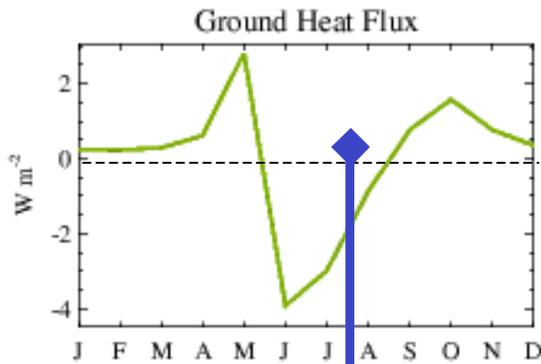
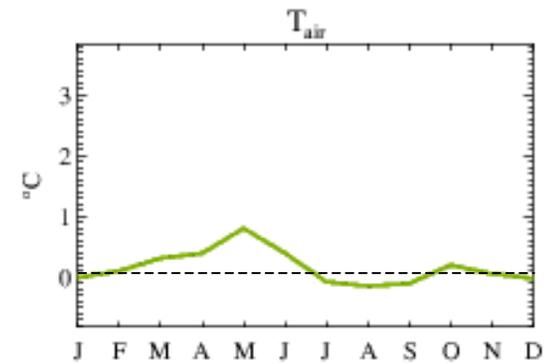


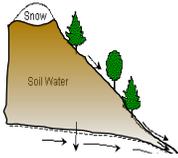
Boreal Shrub





Offline CLM4: SHRUB expt – GRASS expt

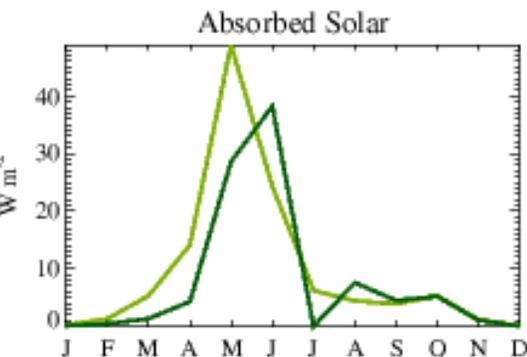




Coupled CAM4/CLM4: SHRUB expt – GRASS expt

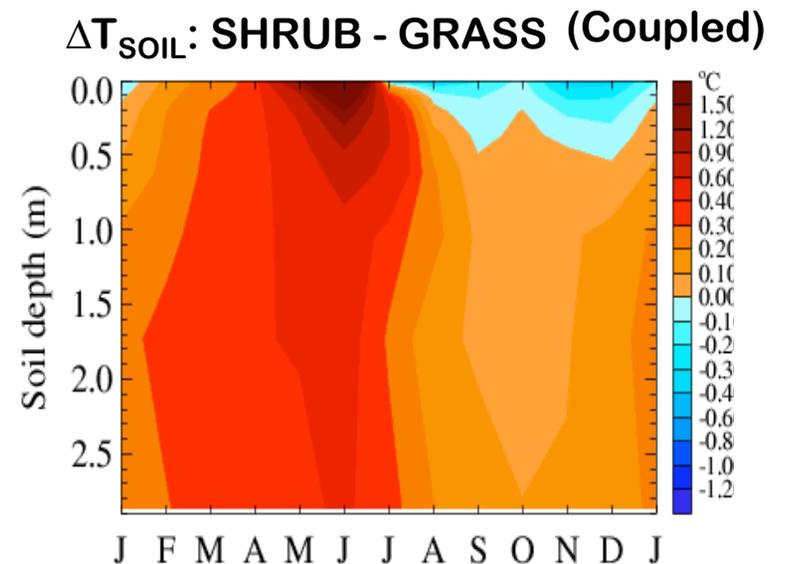
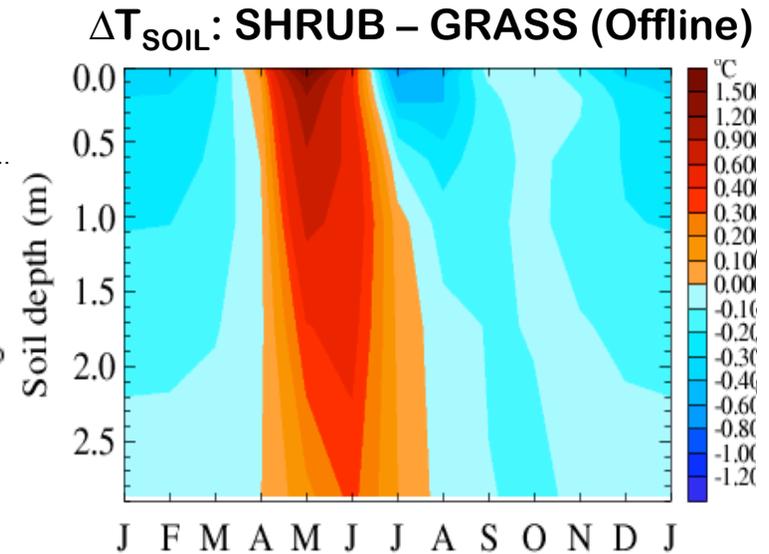
The impact of expanded shrub cover is significantly different in coupled vs offline expts (i.e., increasing vs decreasing ALT)

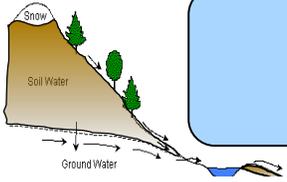
Offline expts roughly mimic a field manipulation study - i.e., surface properties change from grass to shrub but forcing (T, P, Solar, etc) is the same



SHRUB – GRASS

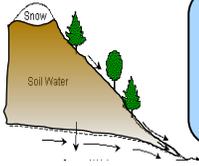
— Offline
— Coupled





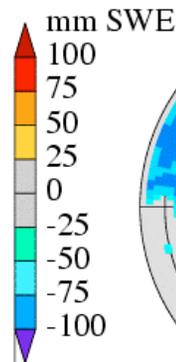
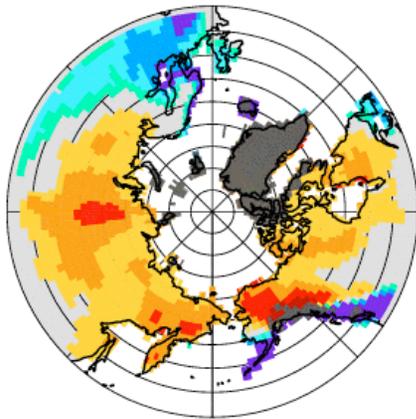
Summary

- **Sea ice loss leads to substantial warming over land, peaking in autumn and winter. In CCSM3, the sea ice loss induced warming extends over 1500km inland.**
- **The rate of sea ice loss may exert a significant control on rate of terrestrial Arctic climate change. Many terrestrial systems may be sensitive to the rate of change.**
- **Projected warming is likely to result in considerable near-surface permafrost thaw, which could initiate positive (and negative) feedbacks that may be relevant on decadal timescales.**
- **Many of these feedback mechanisms are not represented in current generation Earth System Models. Hence, models are missing processes that could be a (small) source of decadal predictability.**

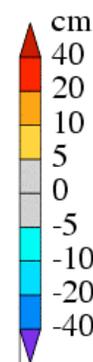
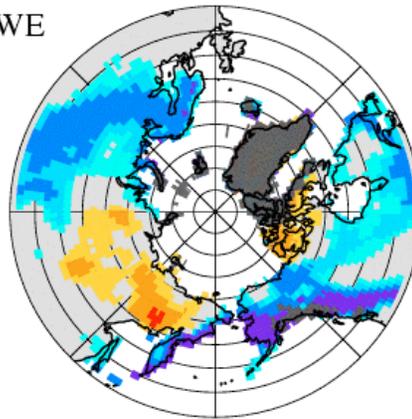


Projected snow changes in CCSM3 (2080-99 – 1950-69)

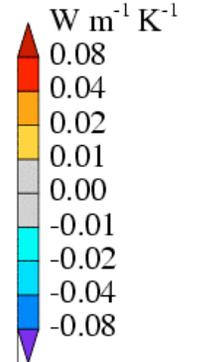
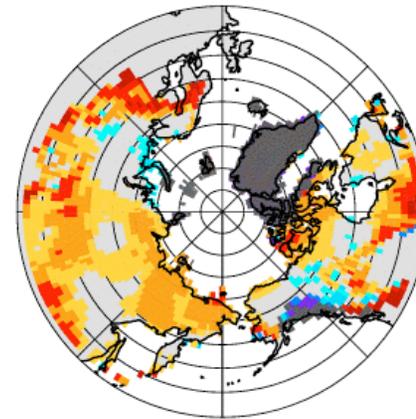
NDJFM Snowfall



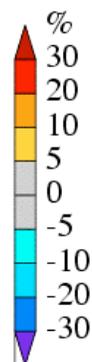
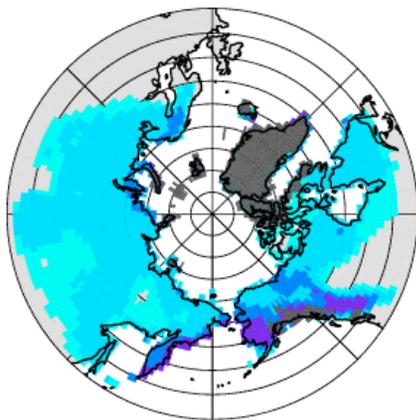
Max Snow Depth



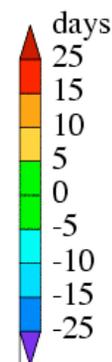
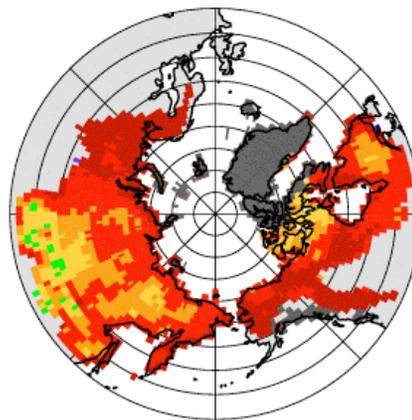
Snow Thermal Conductivity



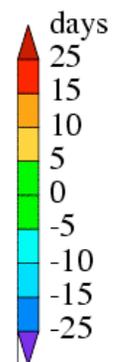
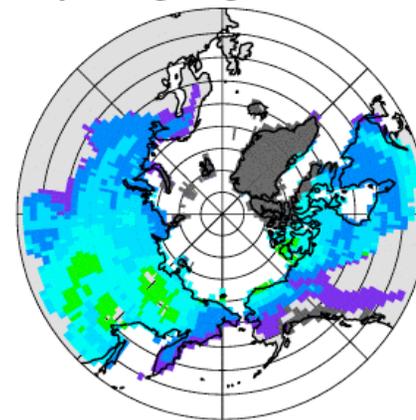
AMSO Snow Cover Fraction

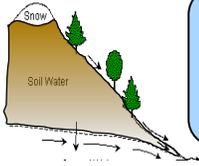


Day of First Autumn Snow Accum



Day of Spring Snow Melt





Prescribed snow experiments

Snow Season Length

HOLD_MW_DP - HOLD_DP

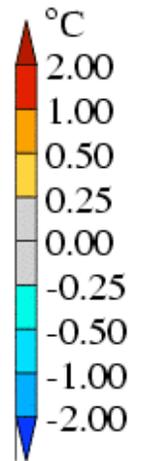
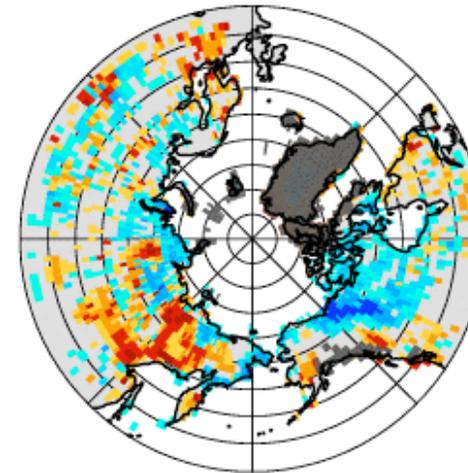
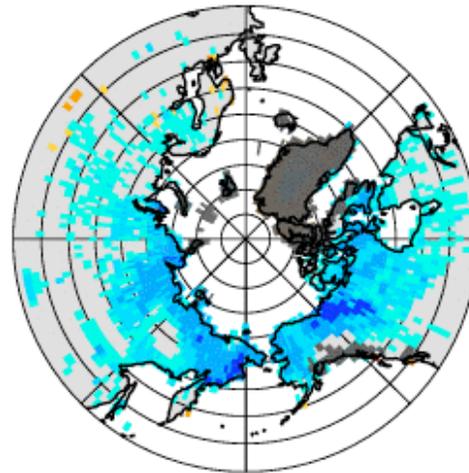
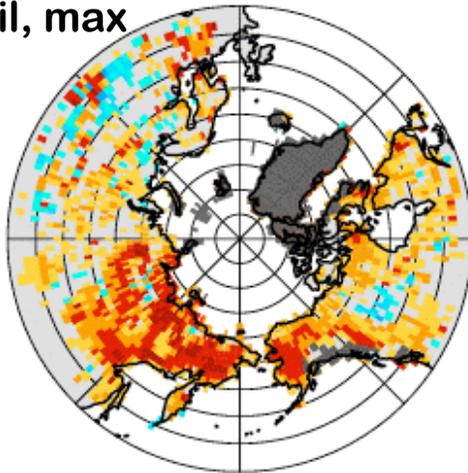
Snow Depth

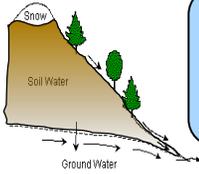
TRND_DP - HOLD_MW_DP

Combined

TRND_DP - HOLD_DP

$\Delta T_{\text{soil, max}}$

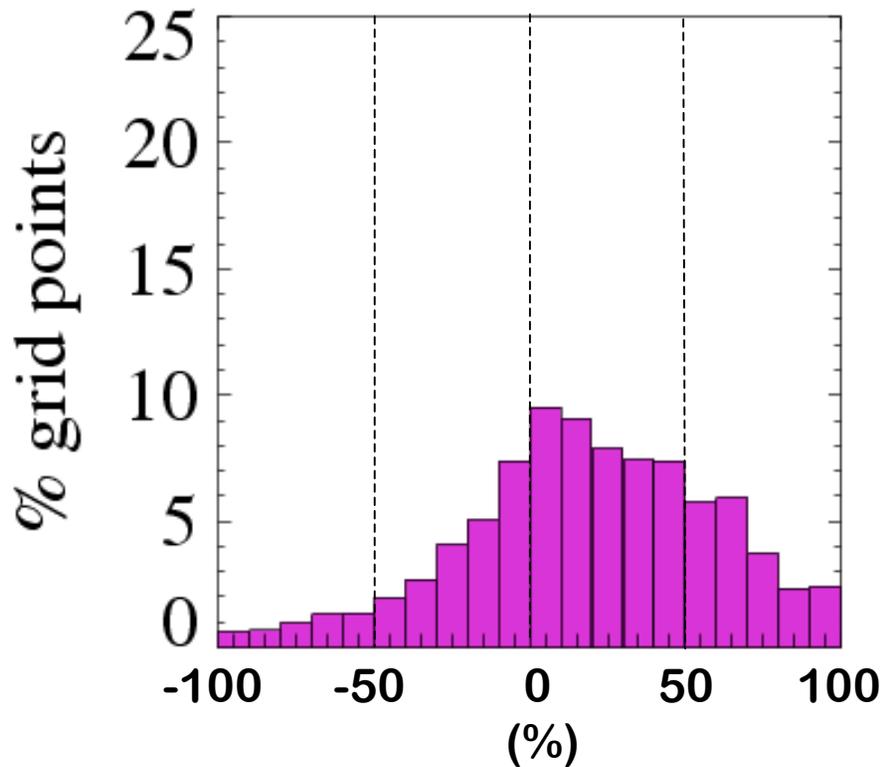




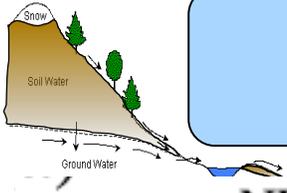
Relative influence of Δsnow vs ΔT_{air} on ΔT_{soil}

% of $\Delta T_{\text{soil,max}}$ (1m depth) attributable to snow state trends for the period ...

1950 to 2000

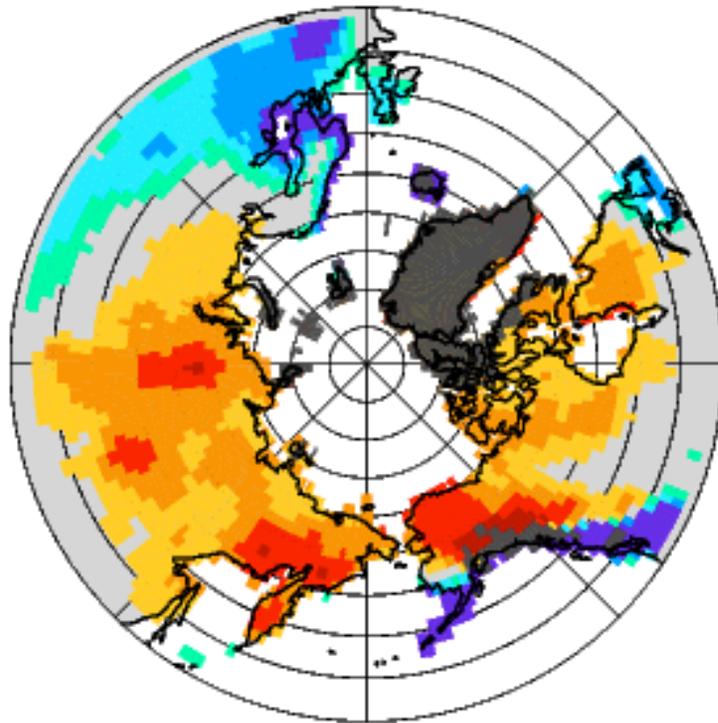


50% means that T_{soil} trend equally due to snow state and T_{air} trends



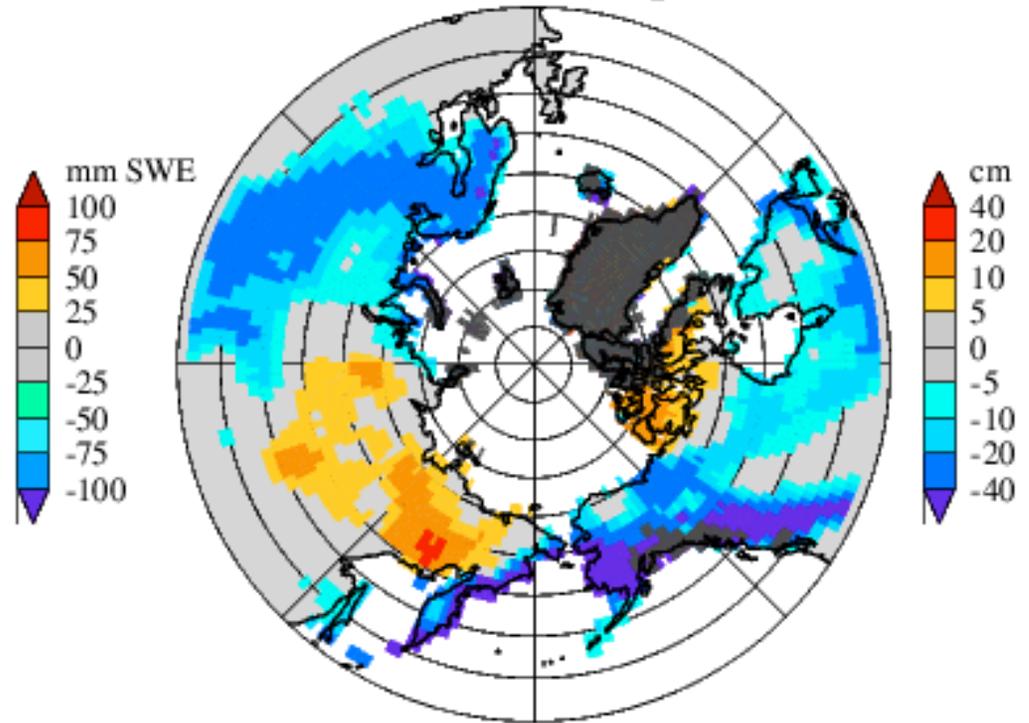
CCSM3 Projected snow changes A1B, 2080-2099 minus 1950-1969

NDJFMA Snowfall

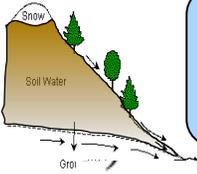


10 – 30% increase in
winter snowfall

Max Snow Depth

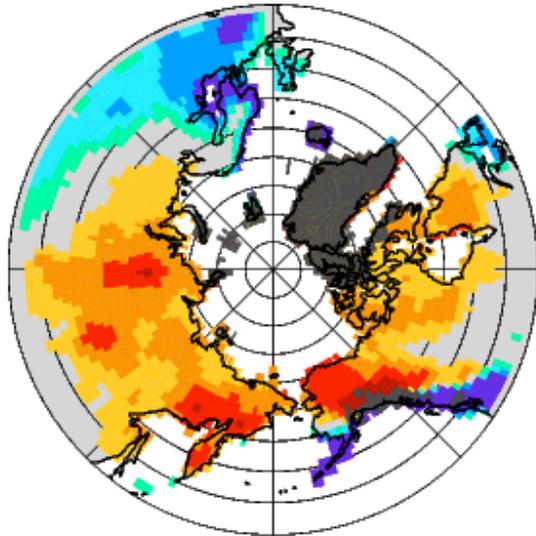


→ deeper OR shallower snowpack

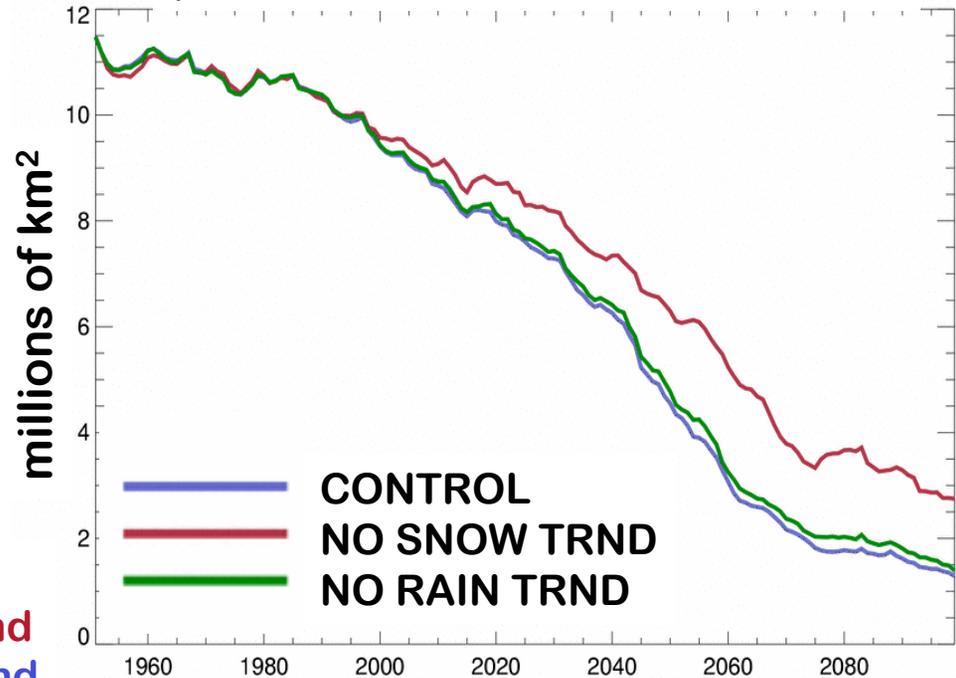


Offline controlled snowfall experiment

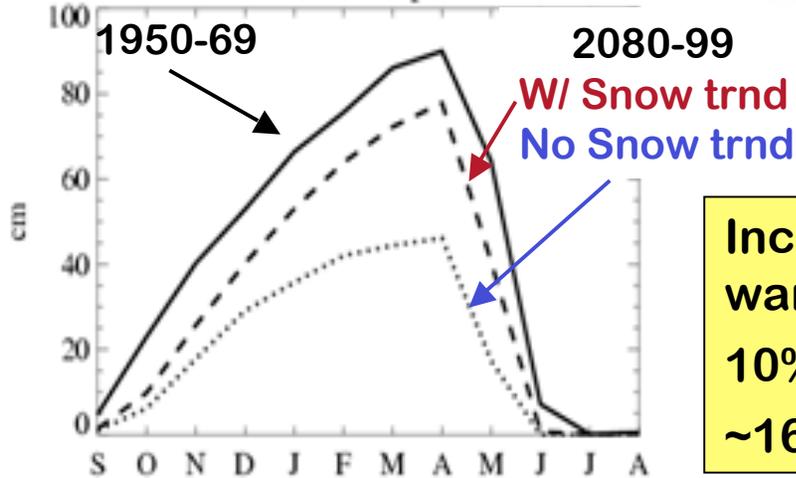
NDJFMA Snowfall



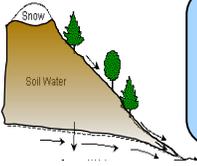
Integrated area with permafrost within 3m of surface



Snow Depth

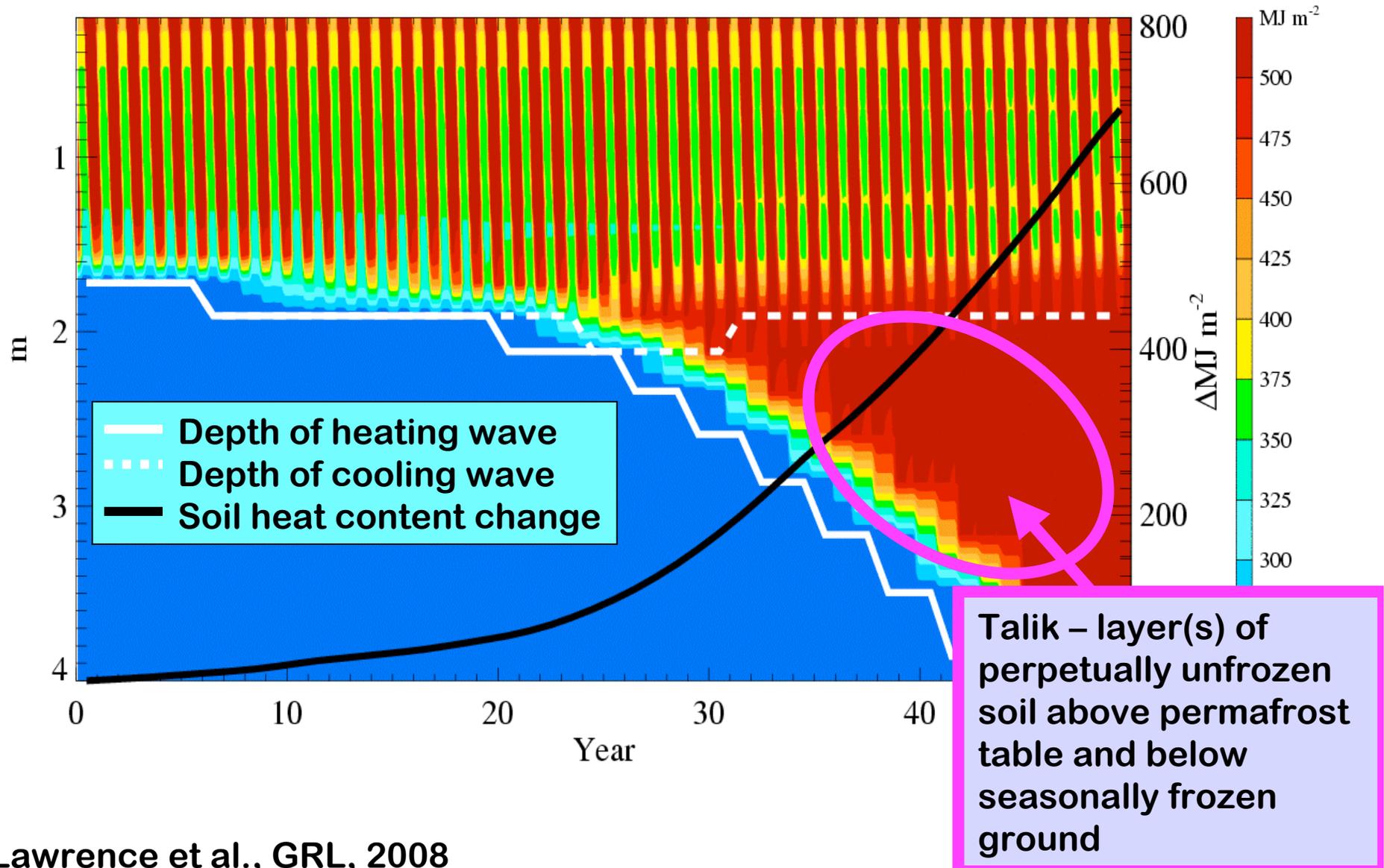


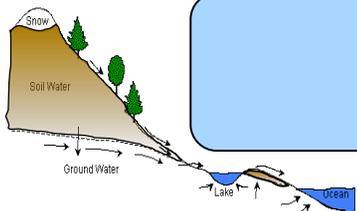
Increasing snowfall is effectively a soil warming agent
10%–30% of total soil warming at 1m depth
~16% contraction of near-surface permafrost



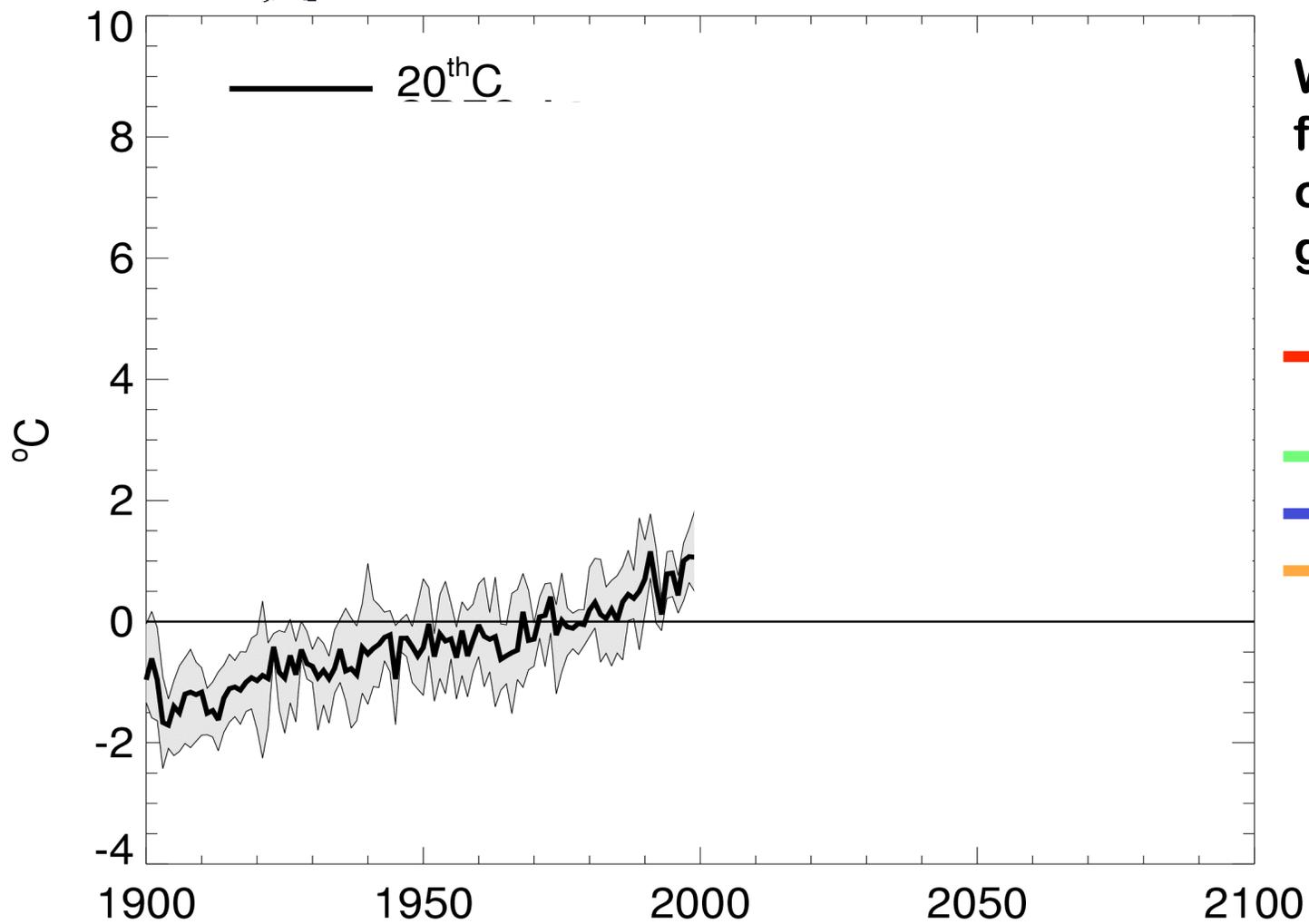
Soil heat accumulation

Warm permafrost case, LINEAR warming expt





Surface air temperature change (CCSM3): Arctic land area



**Warming is ~2x
faster in Arctic
compared to
global average**

