

# **Numerical studies on internal and external variations of the winter polar vortex with a mechanistic circulation model**

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1. General introduction
2. Internal variations of the polar vortex
3. Response to external variations
4. Associated predictability variations
5. Concluding remarks

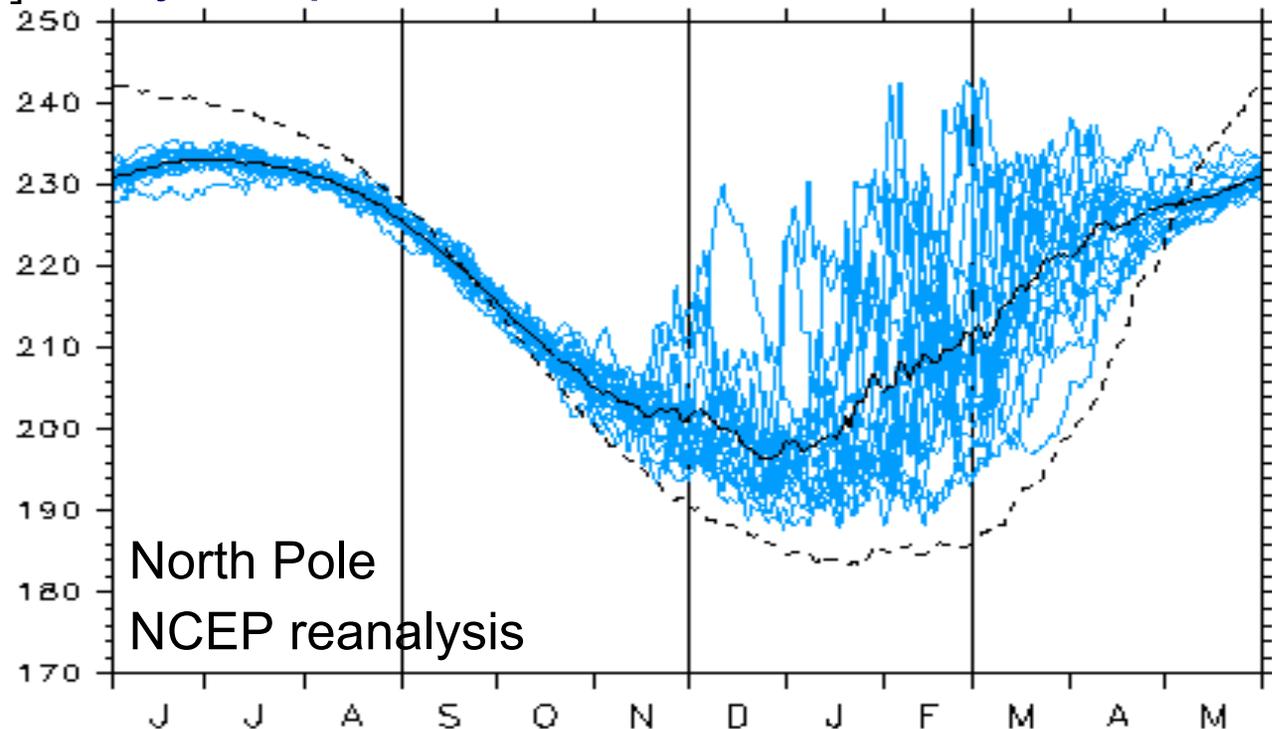
# 1. General introduction

## ❖ stratospheric sudden warming (SSW)

- breakdown of the polar vortex in the winter stratosphere

- a rise of temperature by several tens K in a few days
- several **minor warming** events in a year, while a **major warming** event occurring roughly every two years

[K] Daily temperature at 30 hPa for 1979-1997



Yoden et al.  
(2002, *JMSJ*)

## ❖ intraseasonal and interannual variability in the stratosphere

- Labitzke diagram

- histograms of the monthly mean temperature at the North Pole

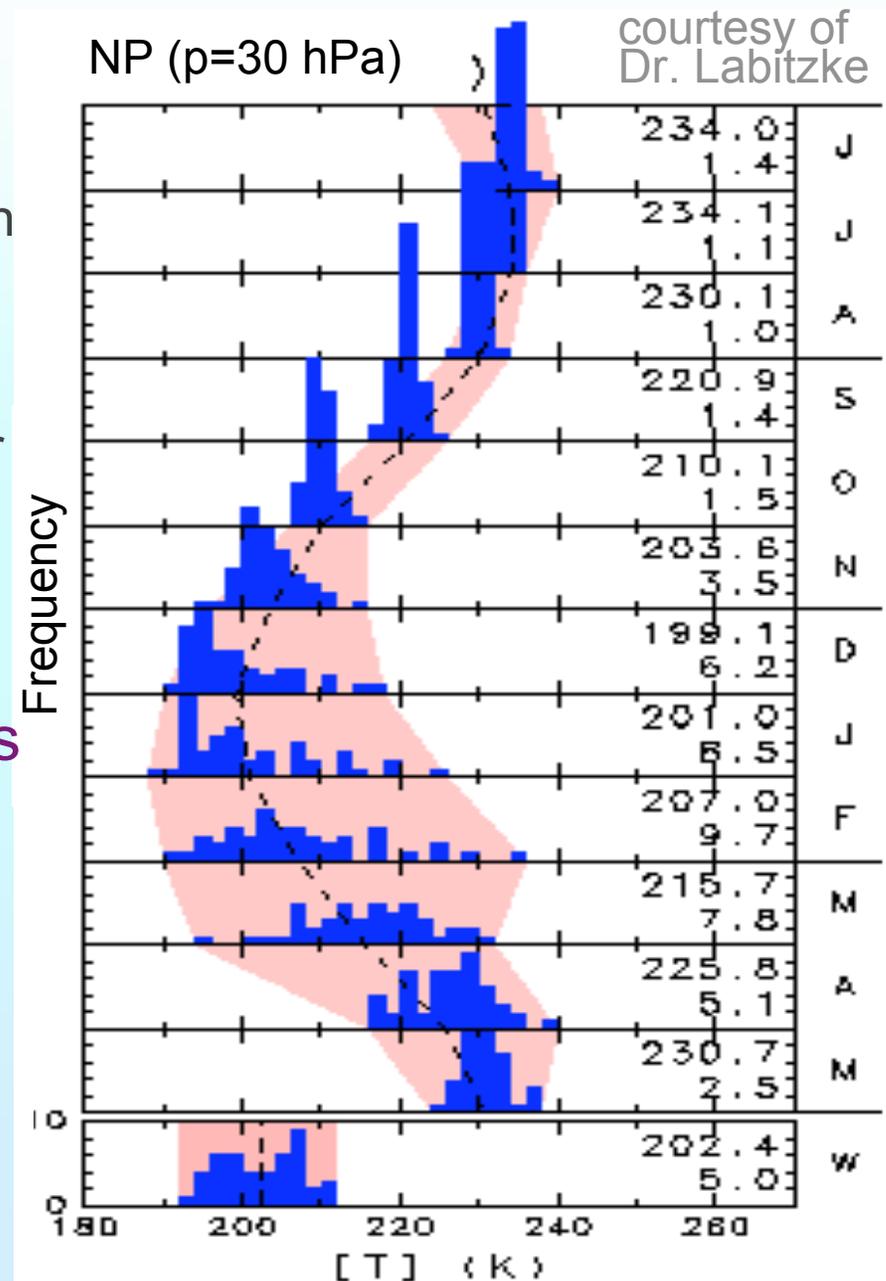
- large variability in winter

- mostly due to the occurrence or non-occurrence of an SSW

- internal variations

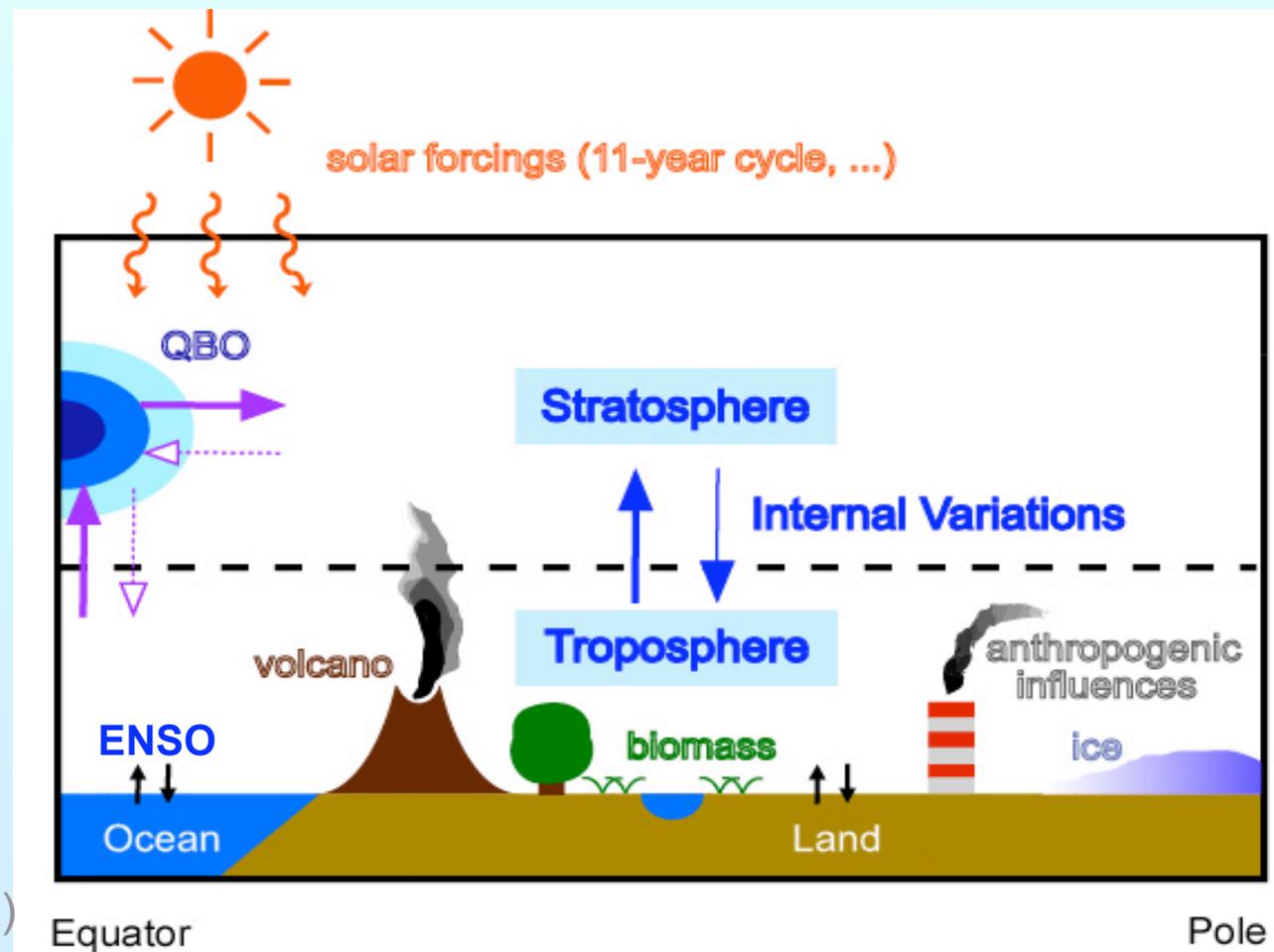
vs

responses to external forcings



- possible “external” forcings

- **out of the atmosphere:** solar forcings, volcanic eruption, biomass, anthropogenic influences (~trend)
- **in the atmosphere** but far from the poles: equatorial QBO, ENSO



Yoden et al.  
(2002, *JMSJ*)

- possible “external” forcings

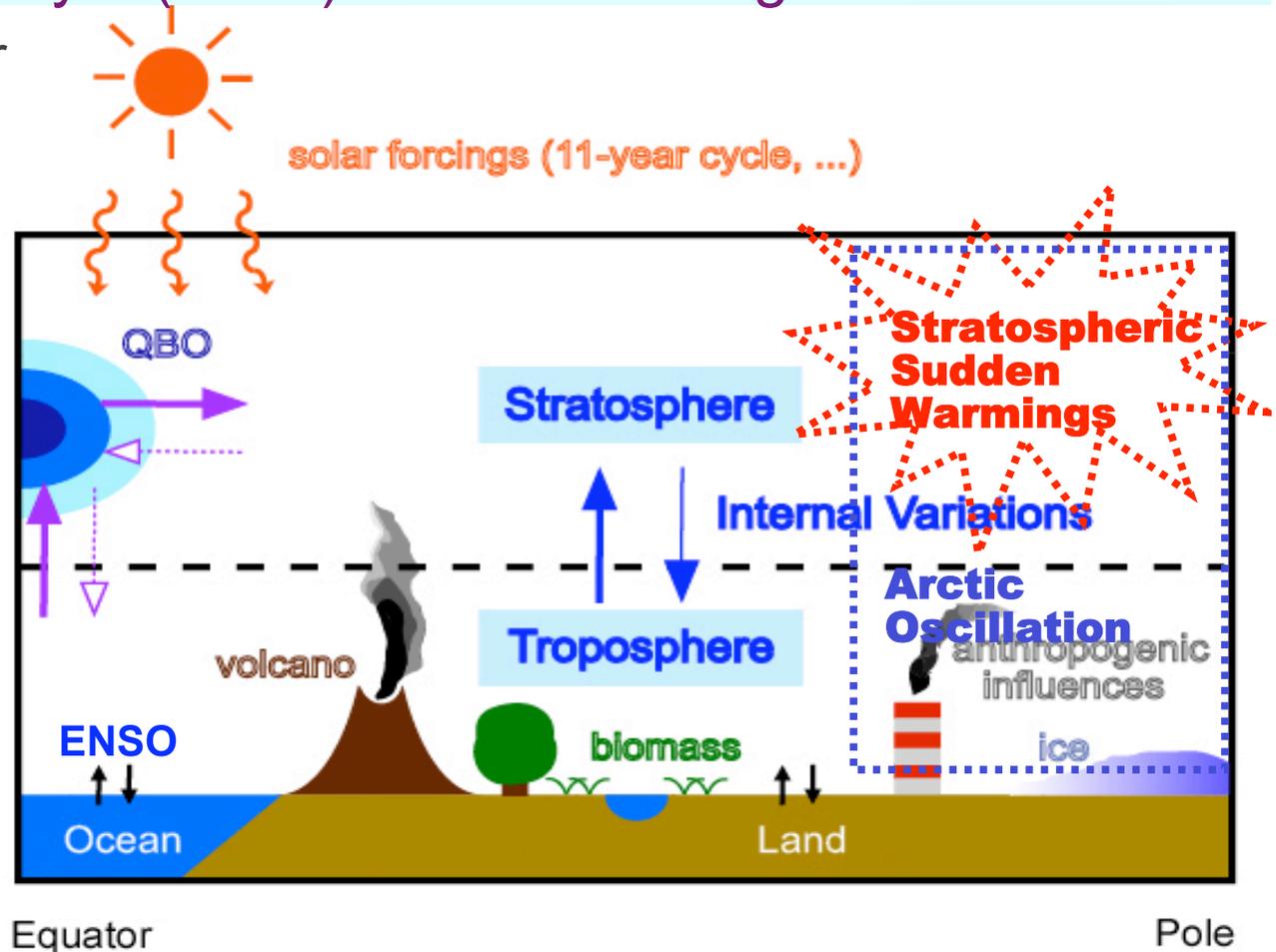
- out of the atmosphere: solar forcings, volcanic eruption, biomass, anthropogenic influences (~trend)
- in the atmosphere but far from the poles: equatorial QBO, ENSO

- SSW and associated variations could be a key process which may amplify a (small) external forcing

- highly nonlinear
- c.f. stochastic resonance

- An example: QBO influence

Yoden et al.  
(2002, *JMSJ*)

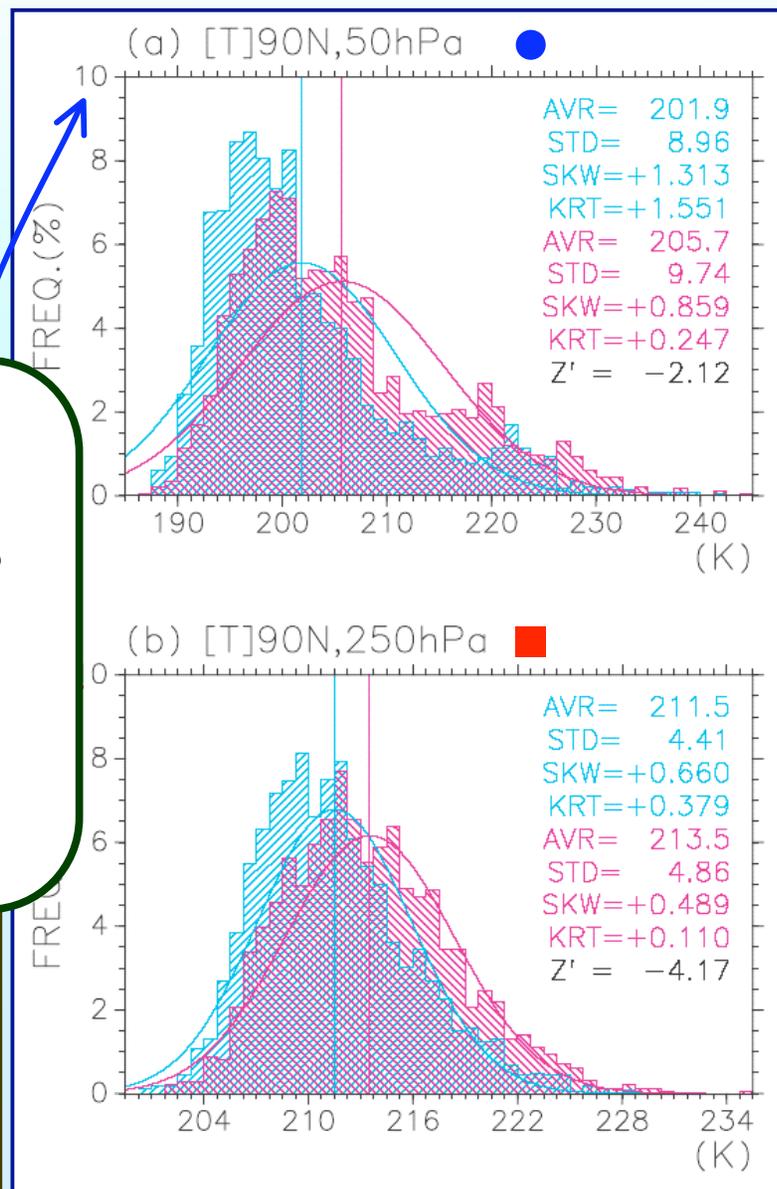
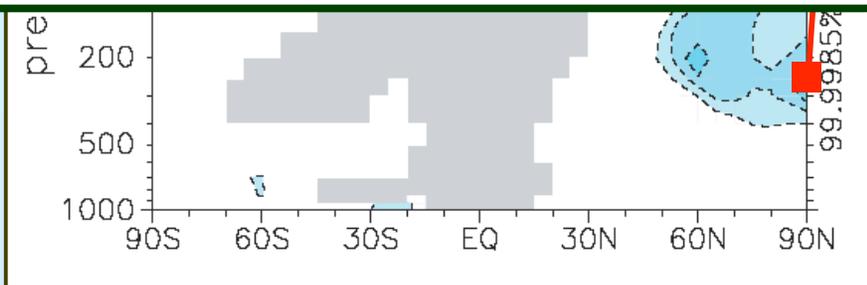


- Naito and Yoden (2005, SOLA)  
 “Statistical analysis of the QBO effects on the extratropical stratosphere and troposphere”

- NCEP/NCAR reanalysis data:  
 ~2,000 days for each of  
 Westerly or Easterly phase



Highly significant differences between W'ly and E'ly phases but heavy overlapping of PDFs due to internal variations



## ❖ some difficulty in observational studies

- data length is limited

  - at most 50 years (e.g., Berlin data since 1950s)

- essential difficulty (nearly impossible) in separation of each response of the polar stratosphere variations

  - **highly nonlinear** processes with interactions among dynamics, radiation and/or chemistry with multiple time scales

## ❖ Only **numerical experiments** overcome this difficulty

- can supply much longer data to obtain statistical significance

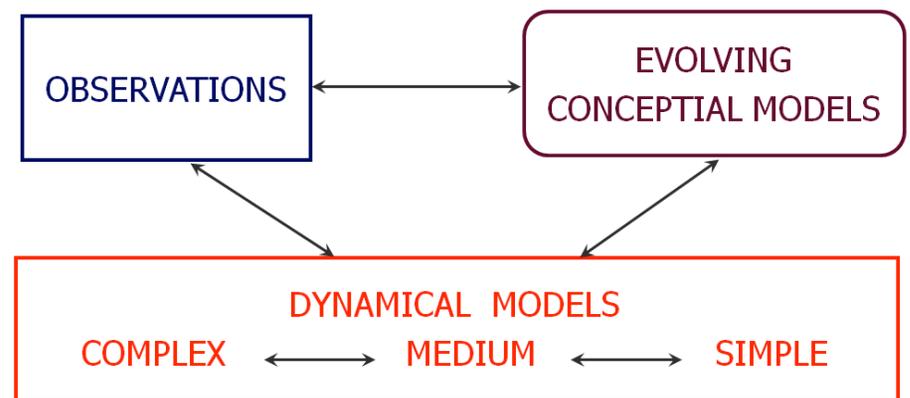
- can do the experiments that control the external forcings

## ❖ hierarchy of numerical models

- Hoskins (1983, *Q.J.R.M.S.*)

“Dynamical processes in the atmosphere and the use of models”

the optimum situation for meteorological research



❖ Over a decade in Kyoto, we have made numerical studies on internal and external variations of the winter polar vortex with a **mechanistic circulation model**

● 3-D global MCM

- GFD Dennou Club AGCM5 (1998)
- Resolution: T21L42 (surface to the mesopause)
- Simplified physical processes:
  - Newtonian heating/cooling (in some cases, under perpetual-winter condition)
  - Rayleigh friction at the surface and at the top sponge layer
  - dry atmosphere
  - idealized surface topography only in NH,  $s = 1$ ,  $amp = 1000m$

- experiments on some parameter dependence with long enough data for statistical significance tests



- experiments on some parameter dependence with long enough data (max ~ 15,000 years) for statistical significance
  - Taguchi, Yamaga and Yoden (2001)
    - SSWs in a stratosphere-troposphere (S-T) coupled system
  - Taguchi and Yoden (2002a,b,c)
    - internal variations associated with SSWs
  - Naito, Taguchi and Yoden (2003)
    - QBO effects
  - Nishizawa and Yoden (2004)
    - annular-mode variability
  - Nishizawa and Yoden (2005)
    - spurious trends due to short dataset
  - Naito and Yoden (2006)
    - QBO effects on SSWs
  - Ito, Naito and Yoden (2009)
    - QBO and 11-year solar cycle
  - Kohma, Nishizawa and Yoden (2010)
    - PJO and fast variations (SSW, VI)

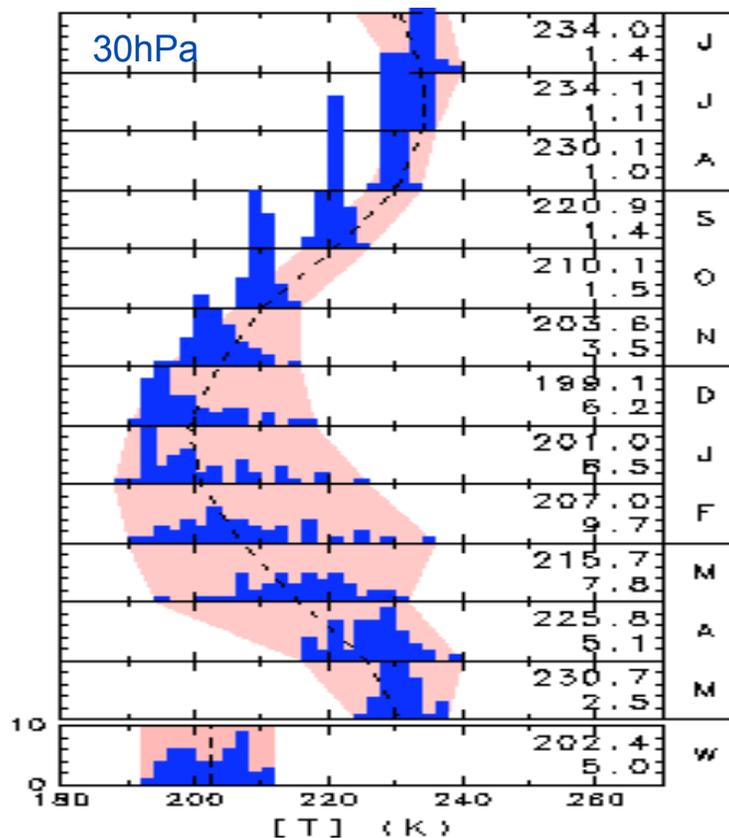
## 2. Internal variations of the polar

### vortex

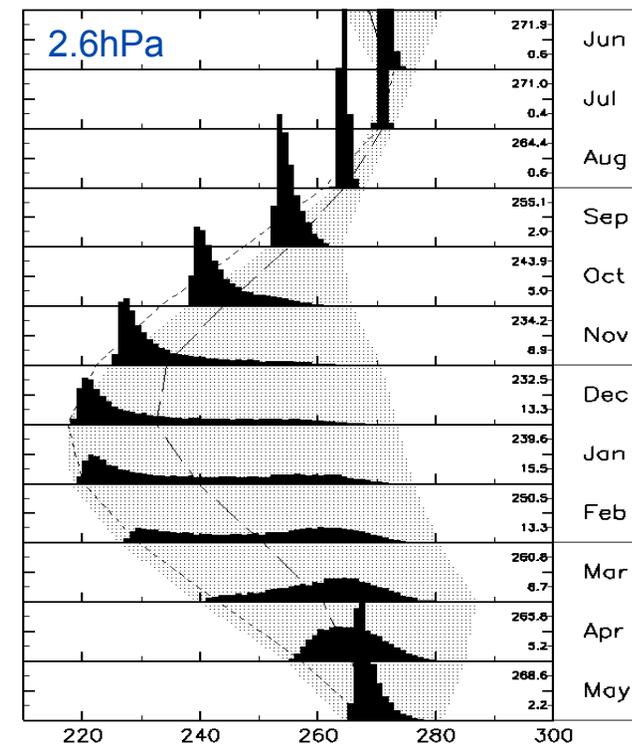
❖ seasonal dependence of internal interannual variability

- due to the occurrence of SSWs in winter stratosphere
- breakdown of the polar vortex is a highly nonlinear process under a purely periodic annual forcing

Real atmosphere (Berlin data)

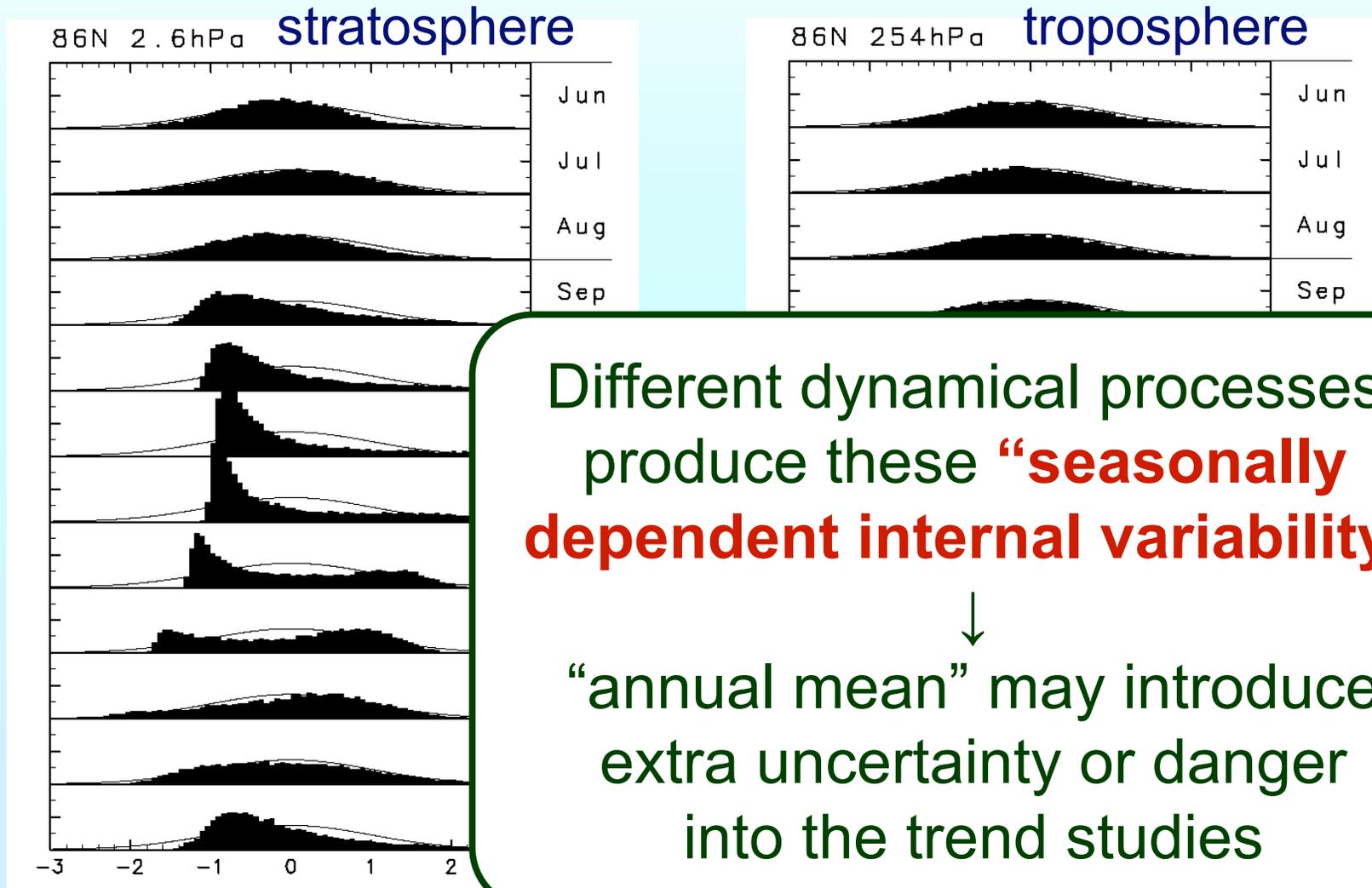


MCM(15,200years)



Labitzke diagram of the polar stratospheric [T]  
(Nishizawa and Yoden, 2005, *JGR*)

- Nishizawa and Yoden (2005, JGR )  
non-Gaussian nature of internal interannual variability
  - normalized pdfs of monthly [T] at the north pole



Different dynamical processes  
produce these **“seasonally  
dependent internal variability”**

↓  
“annual mean” may introduce  
extra uncertainty or danger  
into the trend studies

## ❖ an application: seasonally dependent detectability of a linear trend

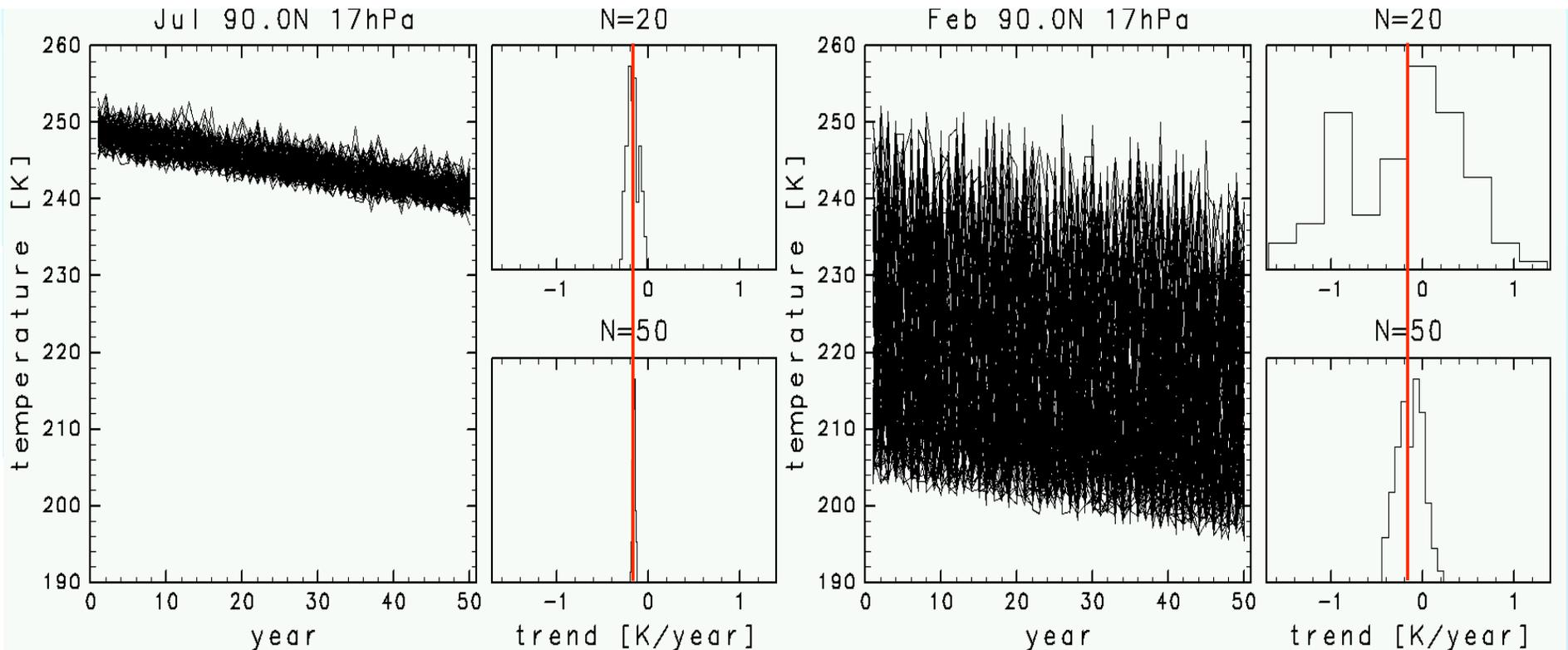
### ● a cooling trend experiment

- 96 ensembles of 50-year integration with an external linear trend  $-0.25\text{K/year}$  around  $1\text{hPa}$

Natural variability:

small in summer (July)

large in winter (Feb.)

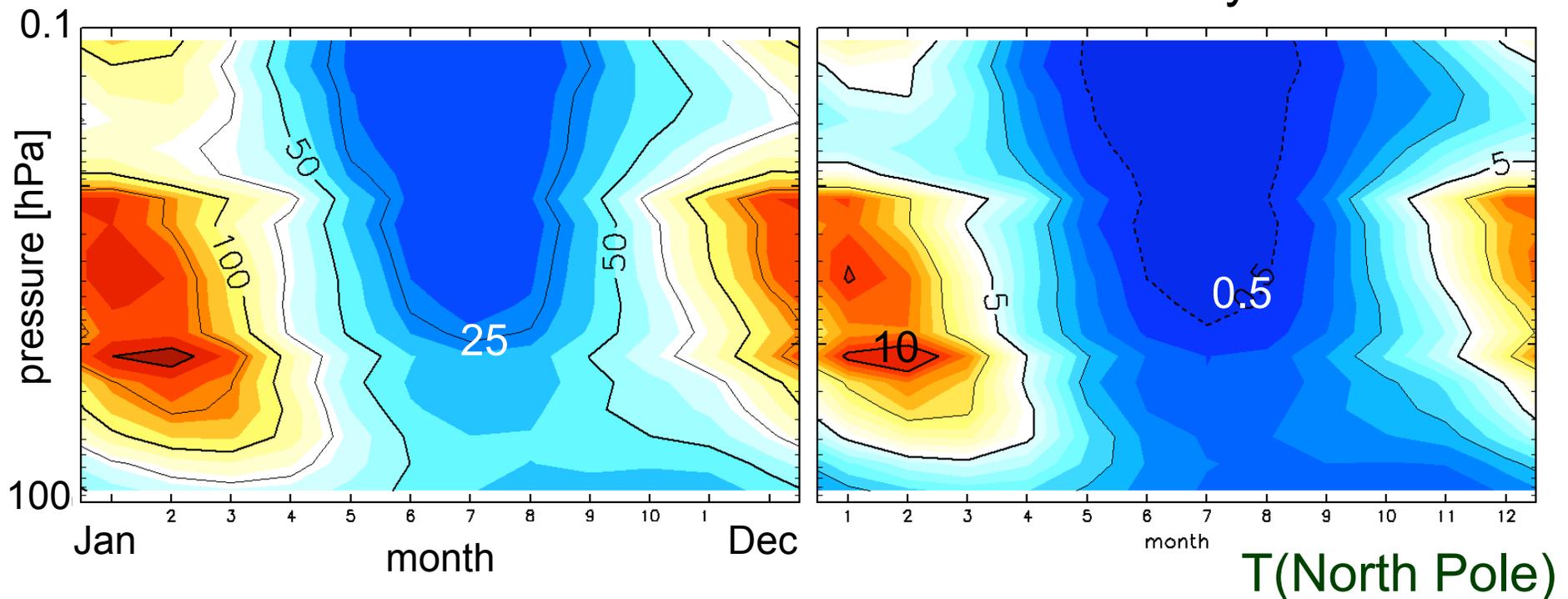


## ❖ seasonally dependent detectability

- How many years do we need to get a statistically significant trend ?
- How small trend can we detect in finite length data with a statistical significance ?

necessary data length [years]  
to detect a linear trend of  
- 0.5K/decade with 90% conf.

necessary magnitude of  
the trend [K/decade] to  
detect with a 20-year dataset



# 3. Response to external variations

## ❖ QBO effects on the occurrence of SSWs

- Naito, Taguchi and Yoden (2003, *JAS* )
- Naito and Yoden (2006, *JAS* )

➤ “QBO forcing” in the zonal momentum eq.:

$$\partial u / \partial t = \dots - \alpha_{QBO} (u - U_{QBO})$$

$U_{QBO}$ : prescribed zonal mean zonal wind of a particular phase of the QBO

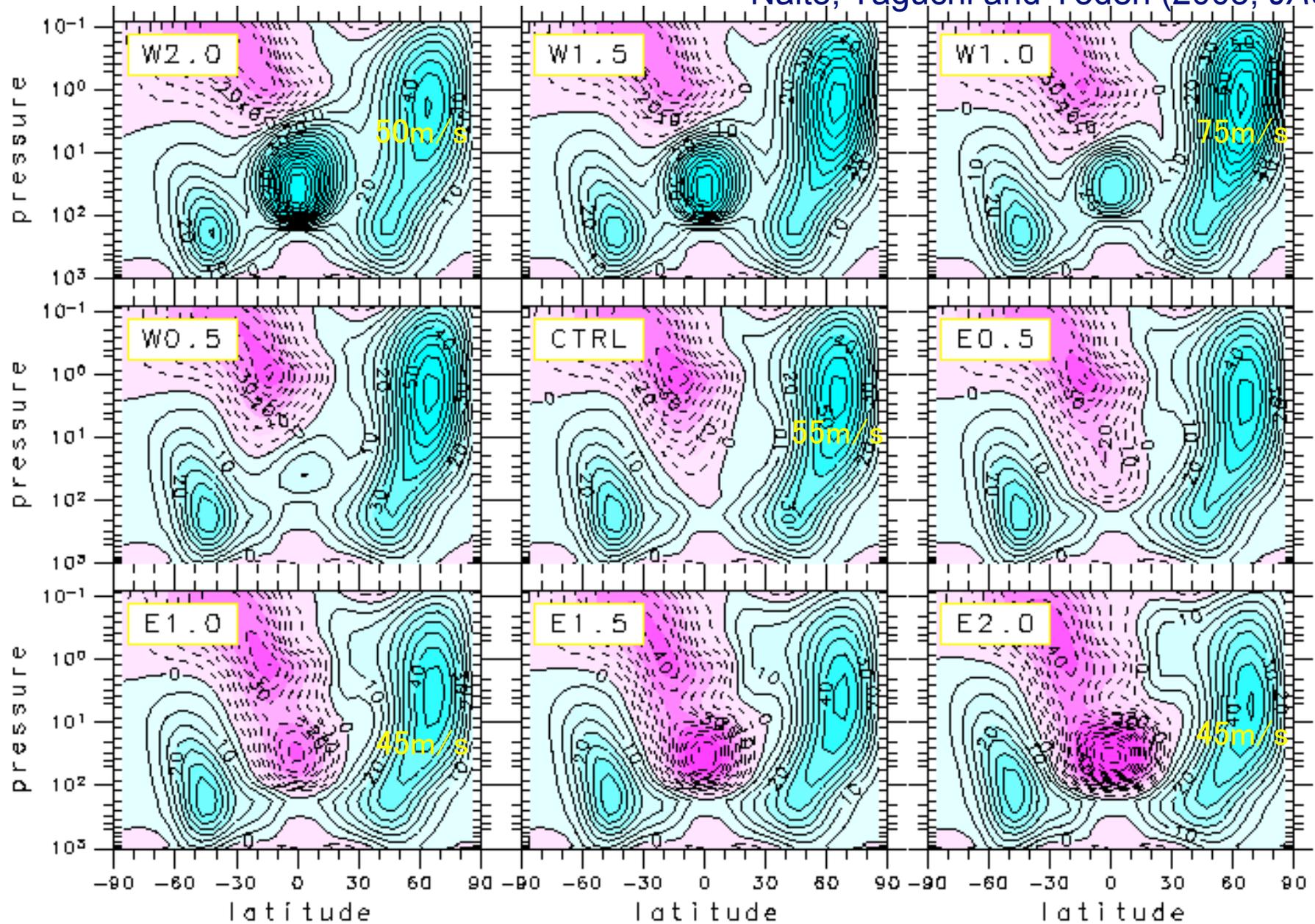
- Under a perpetual winter condition (10,800-day statistics)
- Assess the atmospheric response to a small (or finite) change in the external parameter by a statistical method



Mt. Fuji

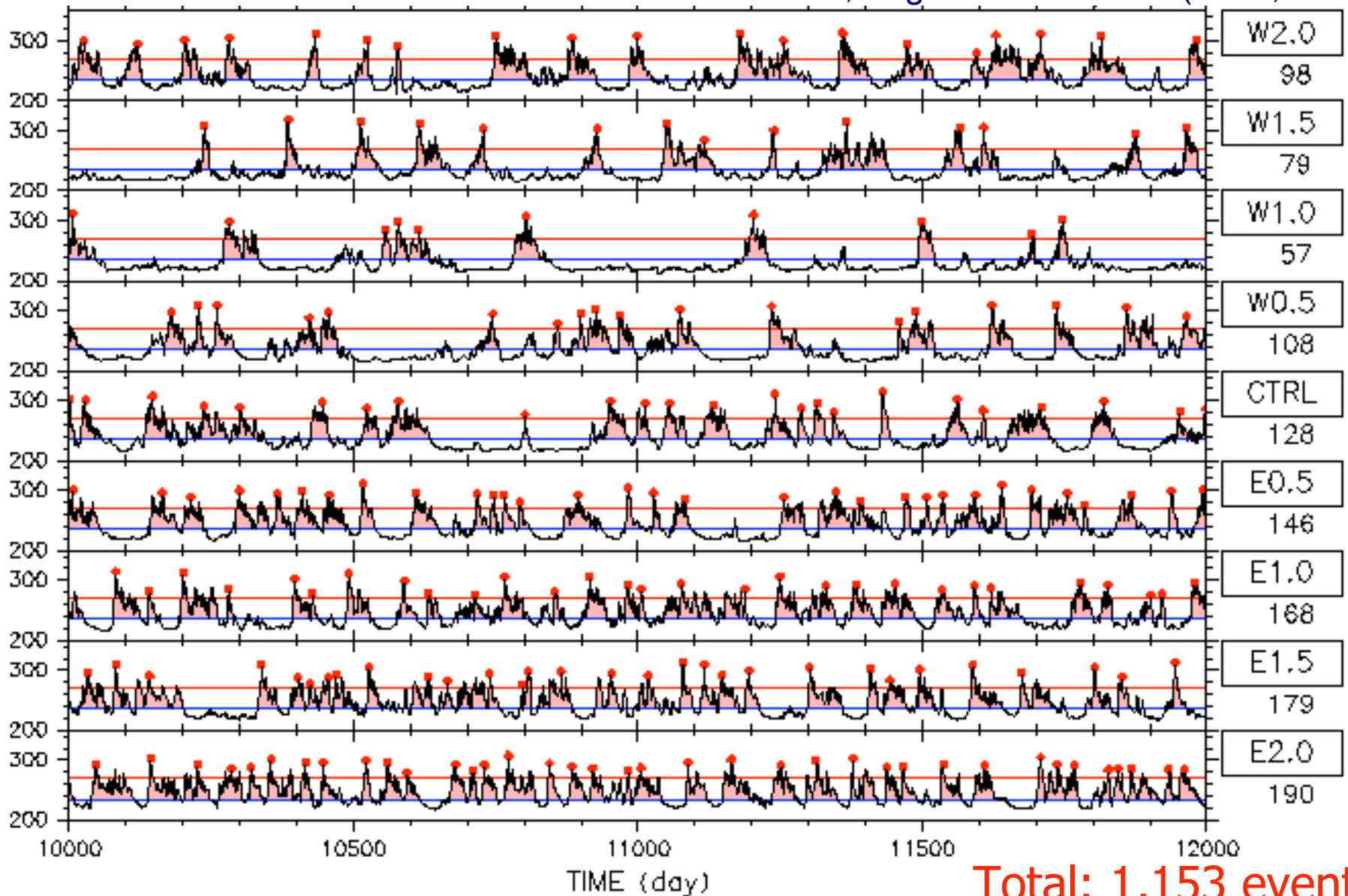
# 10,800-day mean fields of zonal-mean zonal wind [m/s]

Naito, Taguchi and Yoden (2003, JAS)



# Time series of zonal-mean temperature [K] at $\phi=86\text{N}$ , $p=2.6\text{hPa}$ for 2,000 days

Naito, Taguchi and Yoden (2003, *JAS*)



Total: 1,153 events

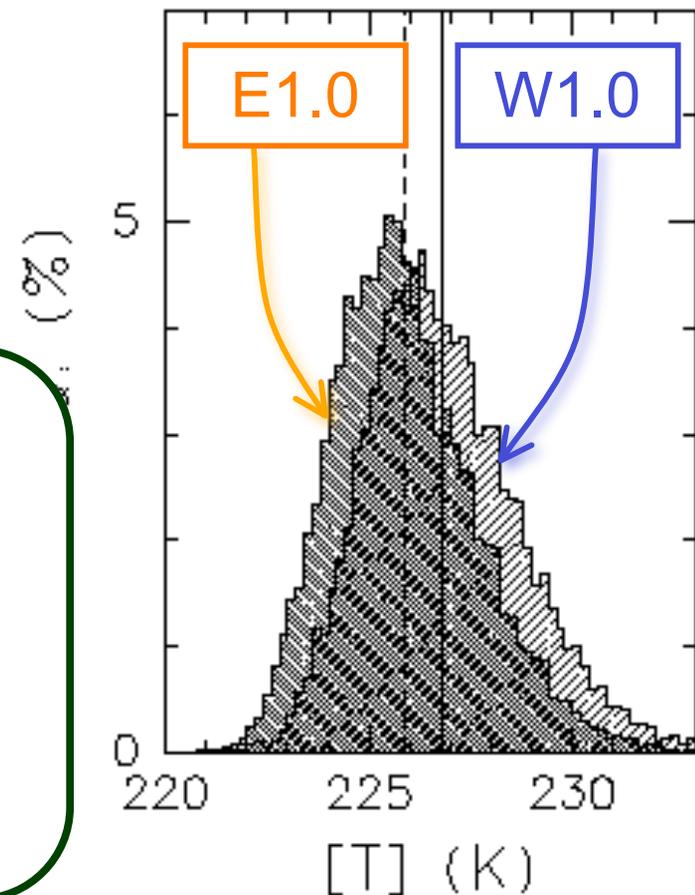
- statistical assessment of difference:  
QBO effects on the polar troposphere
- a large sample method

➤ A standard normal variable:

$$Z = \frac{[T_W] - [T_E]}{\sqrt{\frac{\sigma_W^2}{N_W} + \frac{\sigma_E^2}{N_E}}} = \frac{226.8 - 225.8}{\sqrt{1.87^2 + 1.75^2}} = 40.6$$

Highly significant differences between W'ly and E'ly phases but heavy overlapping of PDFs due to internal variations

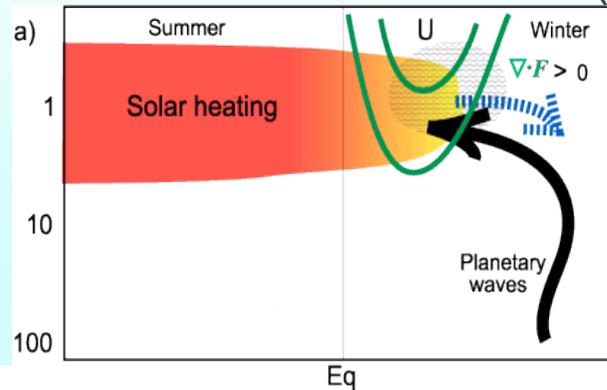
Frequency distributions of zonal-mean temperature [K] (86N, 449hPa, 10800 days)



# ❖ solar effect in the presence of QBO

## ● Ito, Naito and Yoden (2009, *GRL*)

- Examination of Labitzke's relationship (1987, 2006) and Koderu and Kuroda's idea (2002)



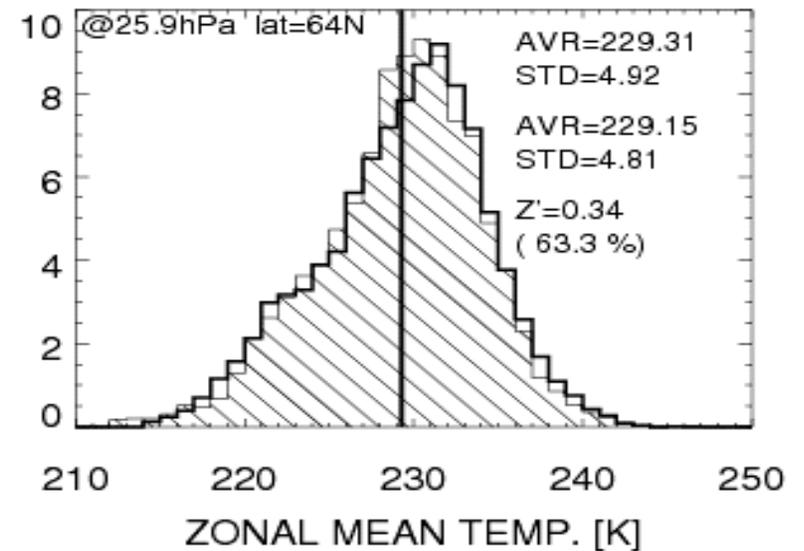
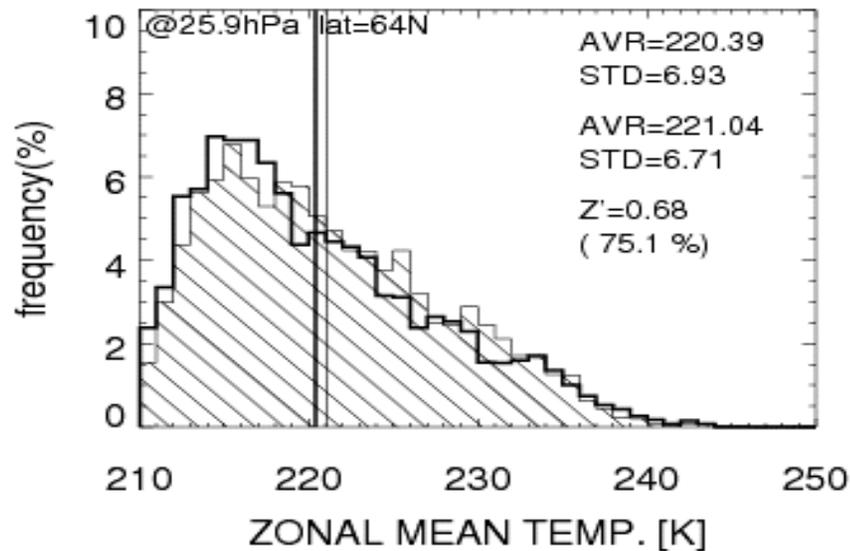
Labitzke (1987)	W	E
Max	W	C
Min	C	W

### QBO Westerly

### QBO Easterly

SOLAR:MIN  
 VS  $\Delta T_{\max} = +2.4\text{K}$

SOLAR:MIN  
 VS  $\Delta T_{\max} = +2.4\text{K}$



## 4. Associated predictability variations

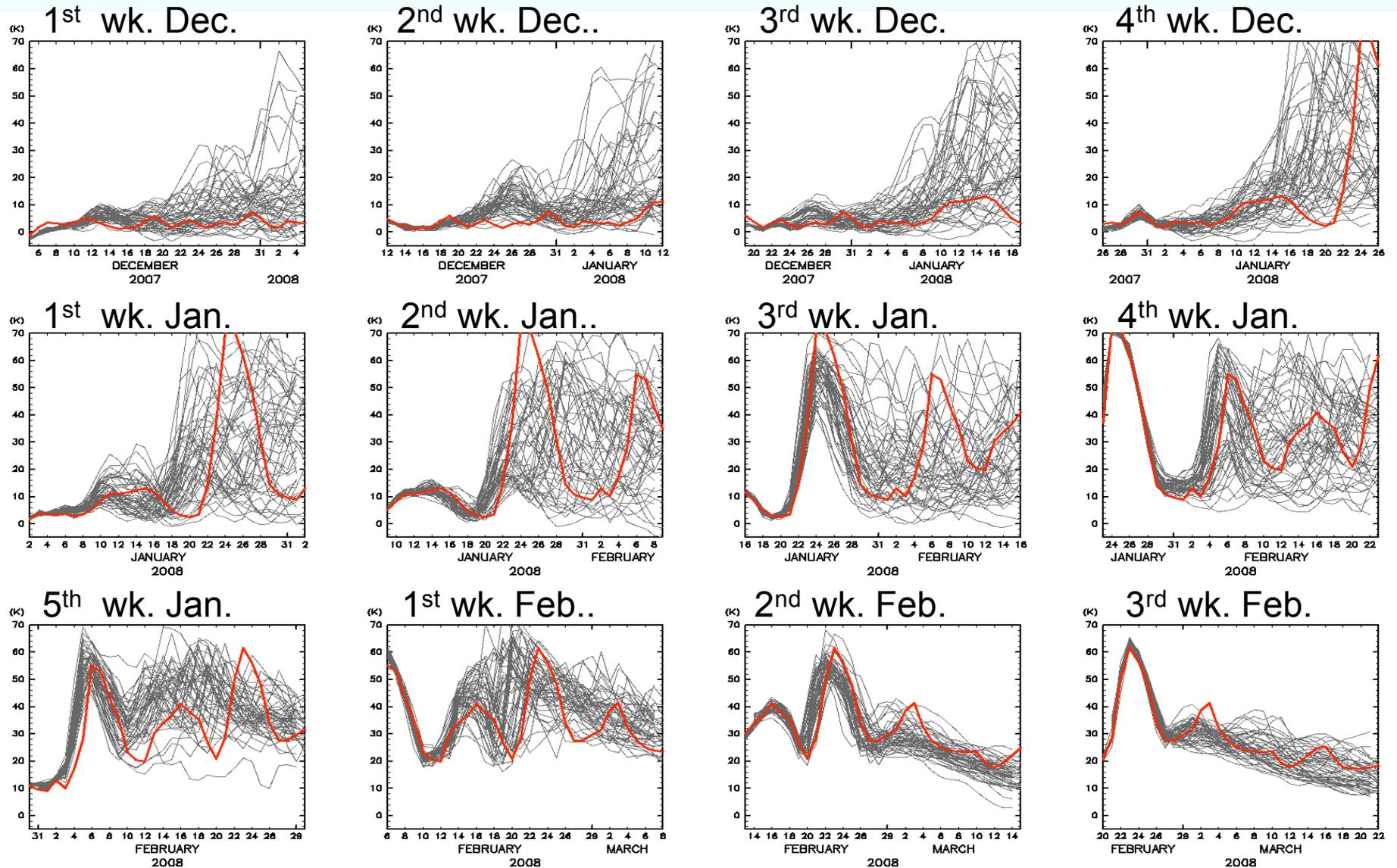
- ❖ a preliminary result on predictability variations in JMA operational **one-month numerical weather predictions (NWP)**
  - global atmospheric model
    - with observed SST anomalies at  $t = 0$
  - full stratosphere
    - $p_{\text{top}} = 0.1$  hPa, 60 layer
  - breeding + time-lagged ensemble forecasts
    - once a week: every Wednesday and Thursday (25+25 = 50 members)

Mt. Rainier



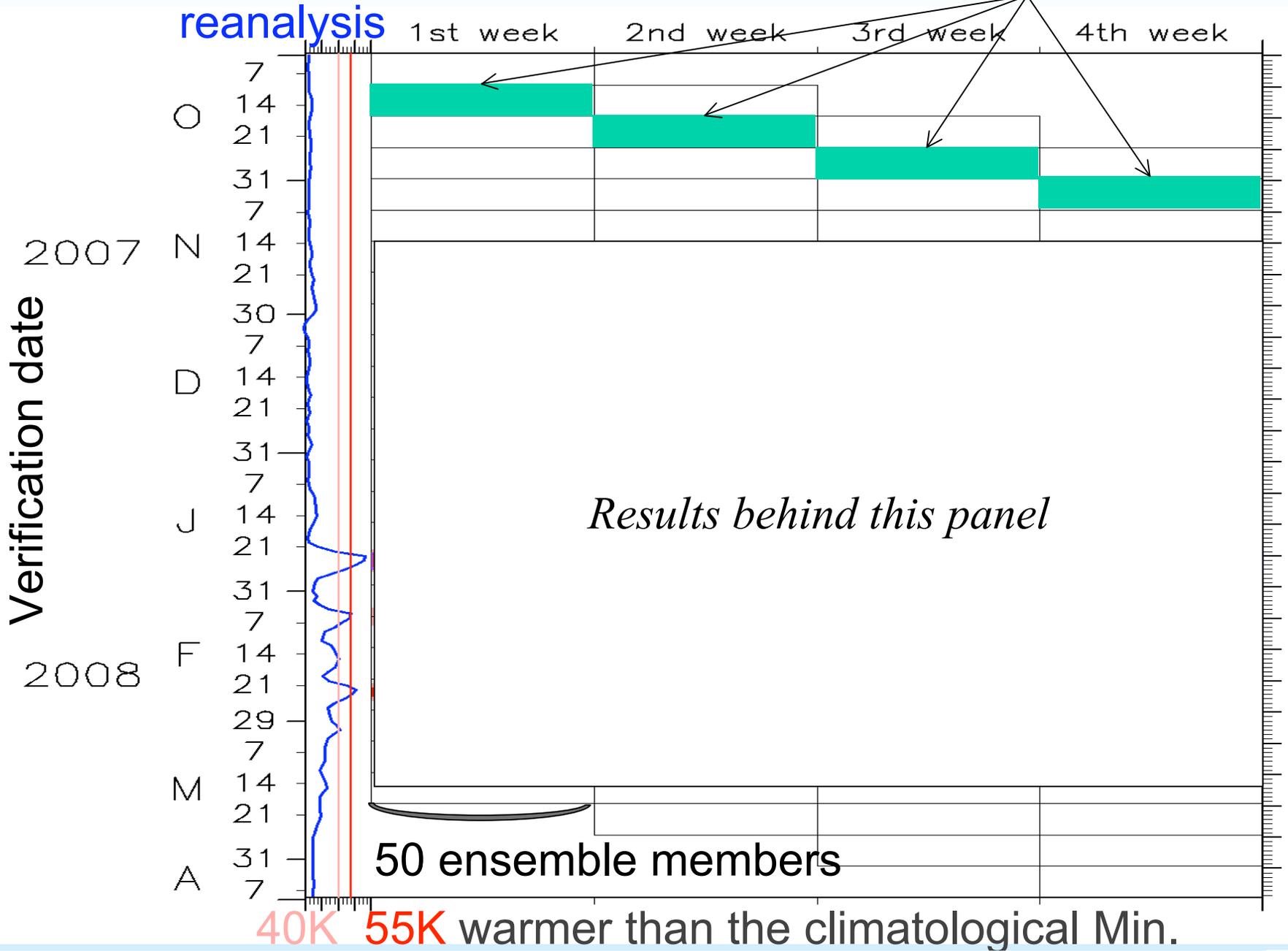
- temperature deviation from the climatological Min. for each calendar day at the North Pole,  $p=10$  hPa in 2007-8 winter

➤ cf. Kohma, Nishizawa and Yoden (2010, *J. Climate*)



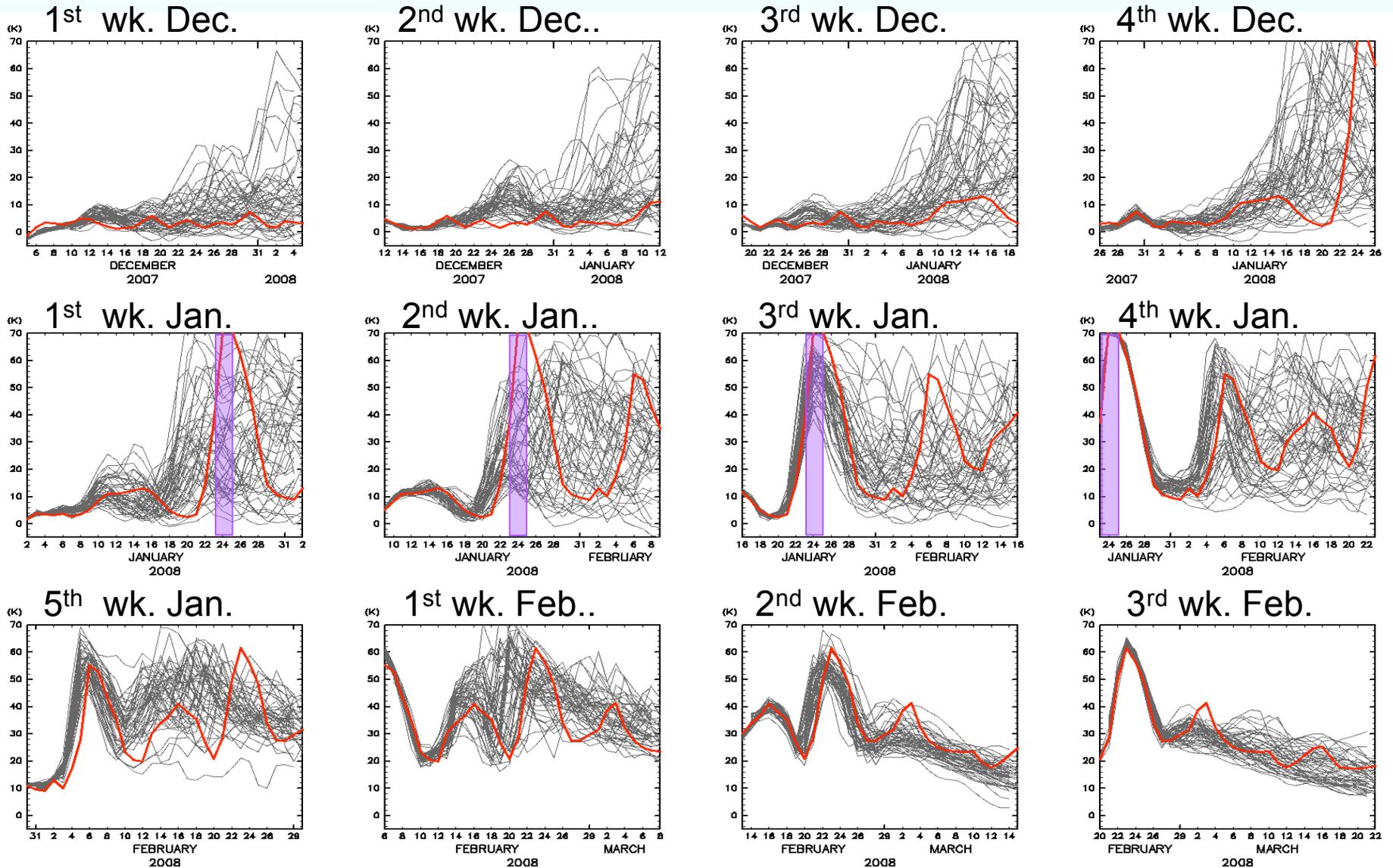
● forecast of extremely warm days

a one-month forecast

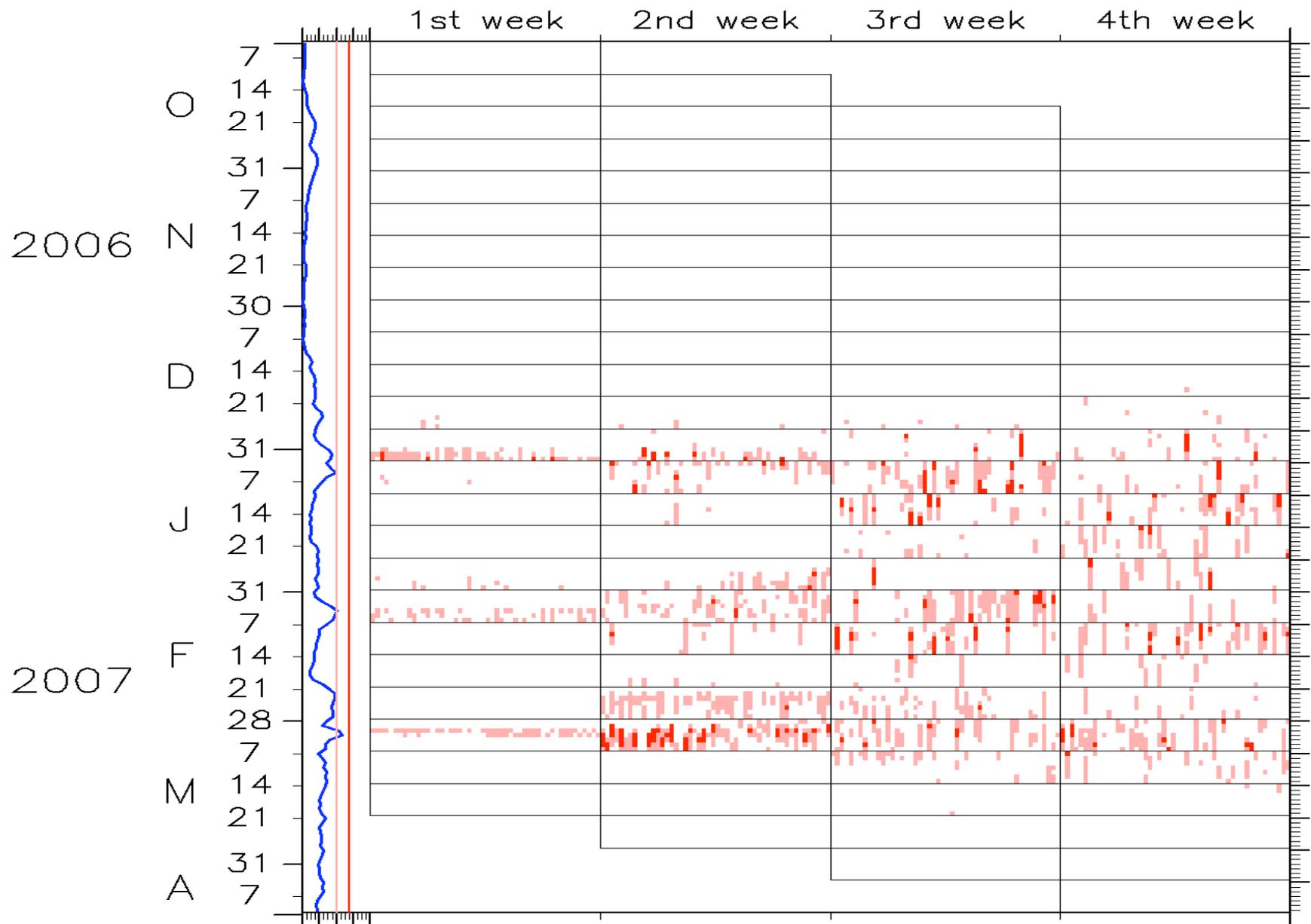


- temperature deviation from the climatological Min. for each calendar day at the North Pole,  $p=10$  hPa in 2007-8 winter

➤ cf. Kohma, Nishizawa and Yoden (2010, *J. Climate*)



● another year: 2006-7 winter



## 5. Concluding remarks

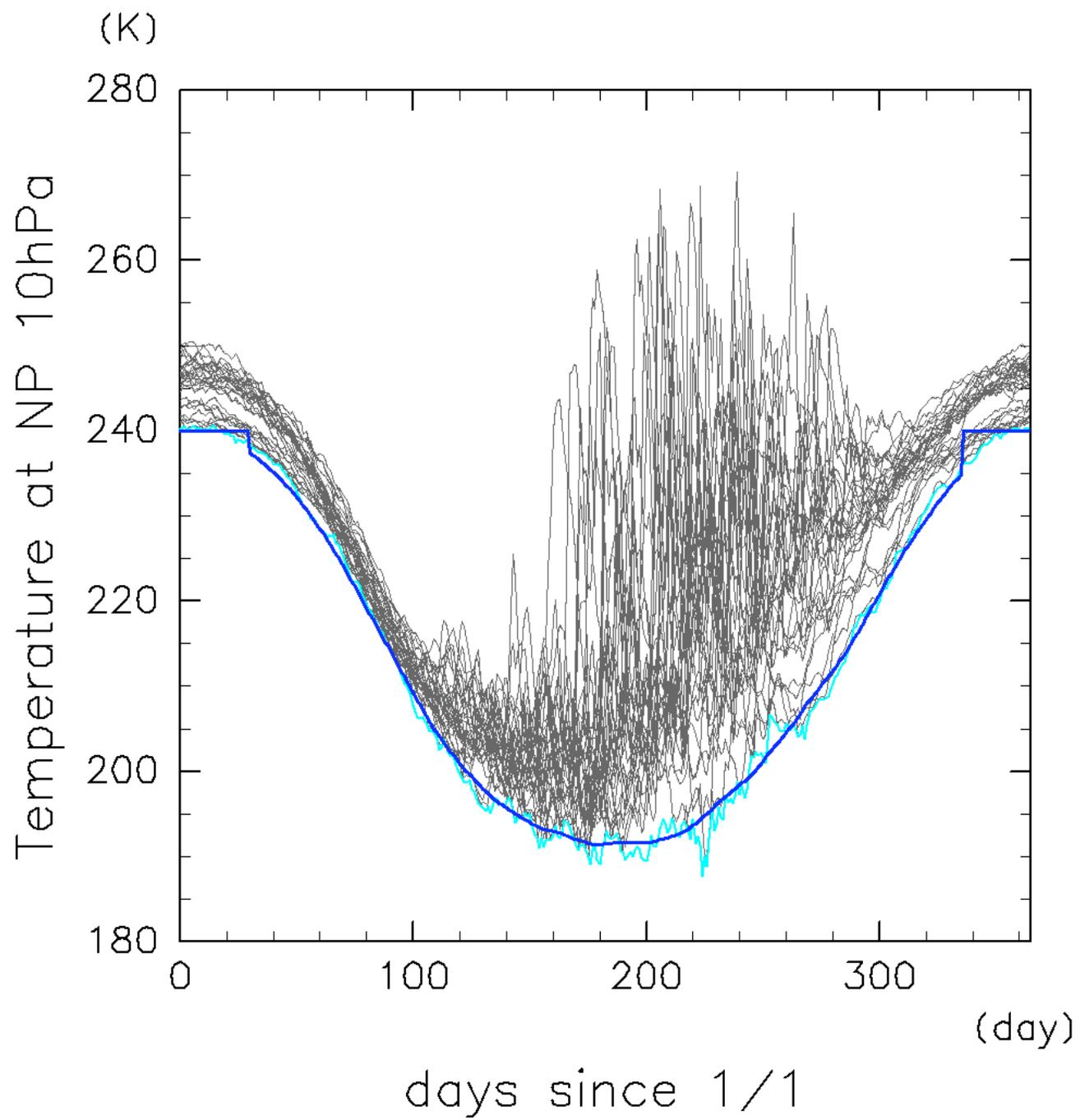
- ❖ **Stratospheric sudden warming** is the most important process to cause intraseasonal and interannual variability in the stratosphere
  - a **highly nonlinear** process: breakdown of the polar vortex
  - mostly (largely) due to **internal dynamics** in planetary scales
  - could be a key process which may amplify a (small) external forcing, such as solar influence, QBO, or else
- ❖ There are some **difficulties in observational studies**
  - **data length** is at most 50 years
  - difficulty in the **separation** of the stratospheric responses to external forcings (solar cycle, QBO, ...) from large internal variations
  - limitation of a cause-result argument for highly nonlinear processes with interactions among dynamics, radiation and/or chemistry with multiple time scales

- ❖ Only **numerical experiments** overcome the difficulties
  - can supply much longer data to obtain **statistical significance**
  - can do the experiments that **control the external forcings**
  - can provide dynamically **consistent and complete data**
- ❖ Advancement in **computing powers** has enabled us to perform **numerical experiments with 3-D MCMs**
  - very long-time integrations to obtain **reliable PDFs**
    - non-Gaussian, bimodal, ...
      - ➔ nonlinear perspectives on climatic variations and trend
    - large sample method is useful for statistical assessment
  - **parameter sweep** experiments to investigate for highly nonlinear processes with combination of external forcings
- ❖ Predictability variations in operational one-month numerical weather predictions look interesting
  - long (3~4 weeks?) lead time for extremely warm days (SSWs)



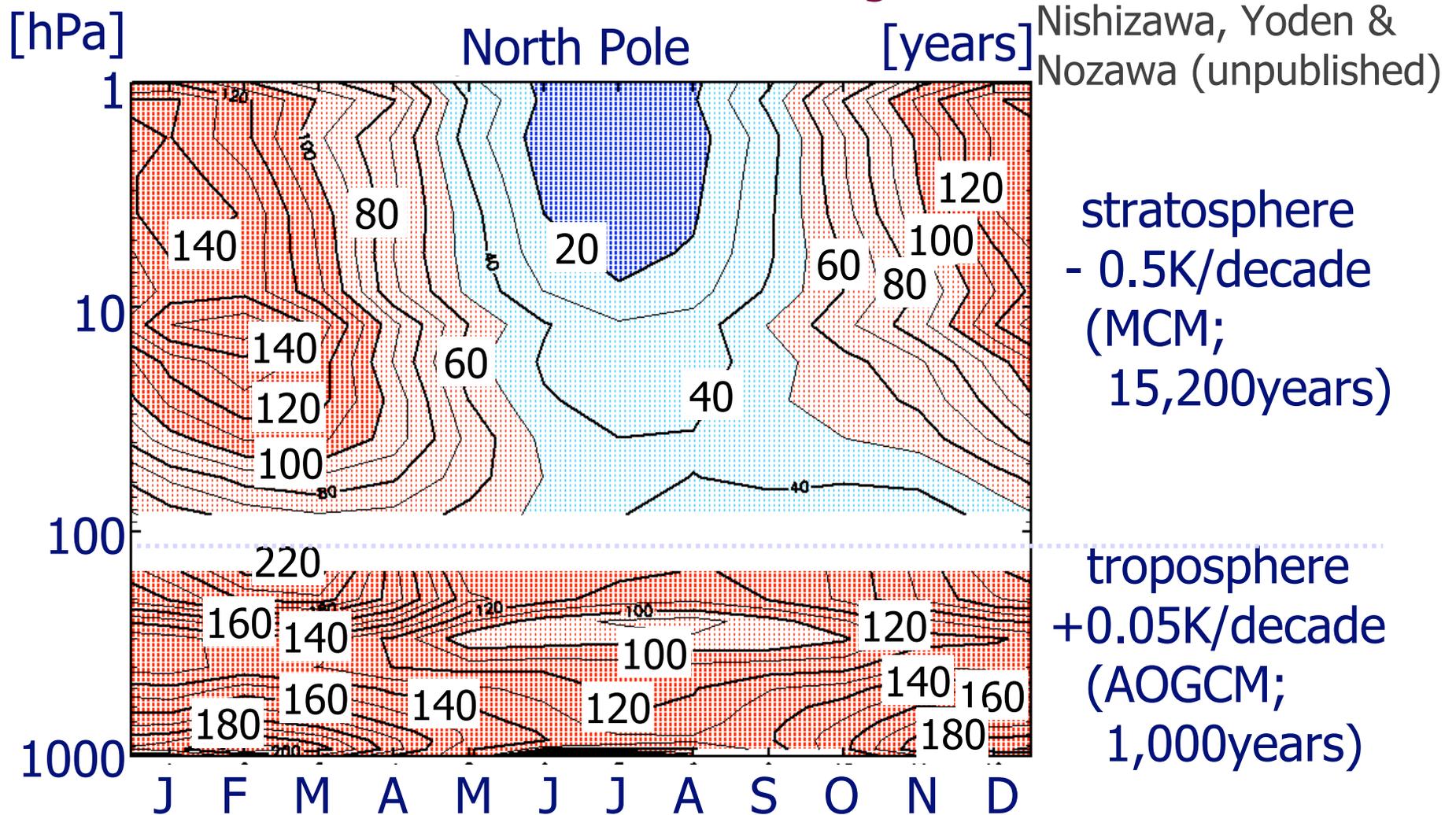
**Thank you !**

June 14, 2010 Mt. Rainier



❖ seasonally dependent detectability in the troposphere

- a natural variability run of AOGCM for 1,000 years
- necessary data length [years] to detect a linear trend of  $+0.05\text{K/decade}$  with 90% statistical significance

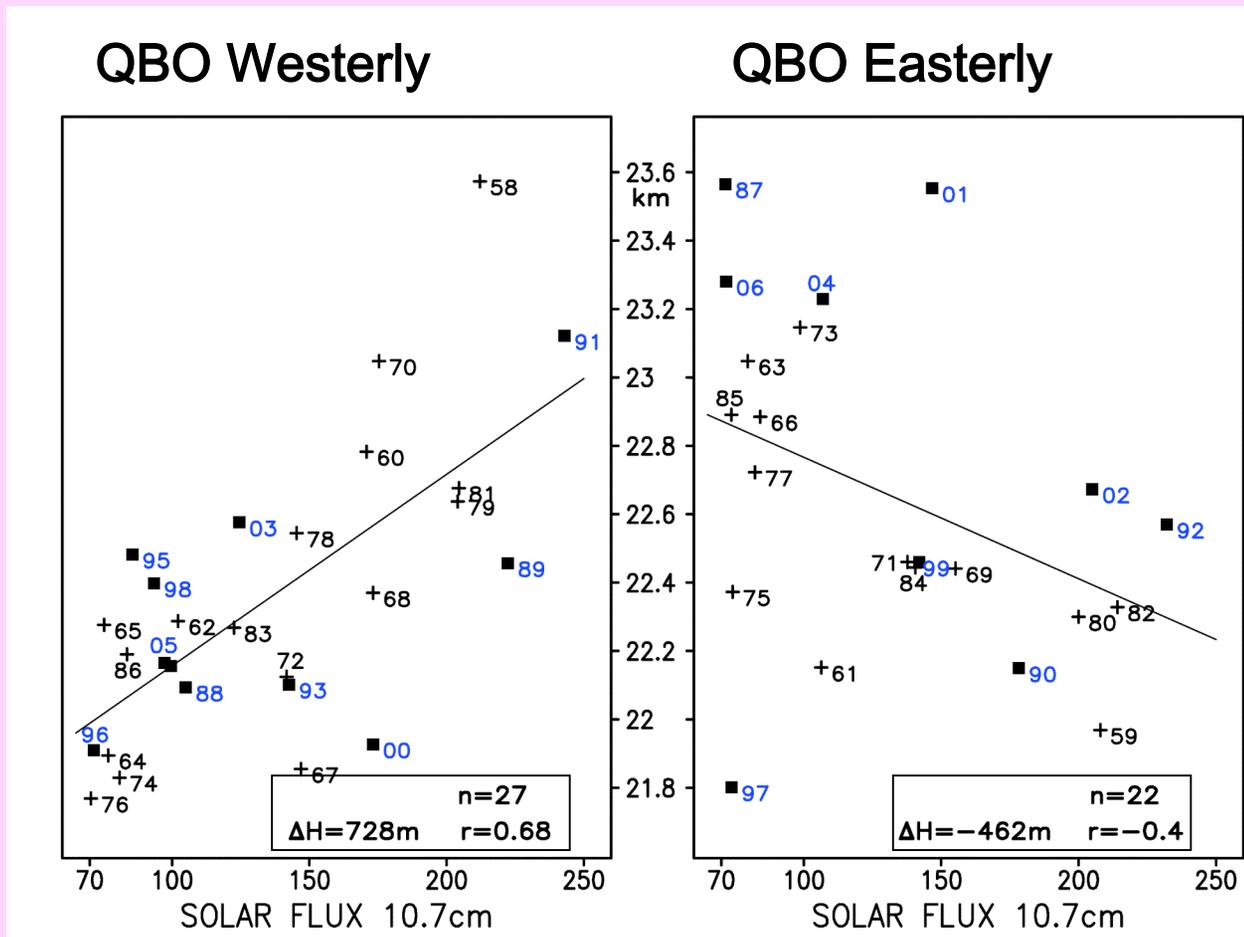


# 3. Solar effect in the presence of QBO

## ❖ Motivations

- Labitzke (1987, 2006)

- Correlations between 30-hPa heights and the solar flux of 10.7cm
- 1958-2006 (49 years; NCEP/NCAR RA), (20 more years, in blue)



	W	E
MA	W	C
MI	C	W

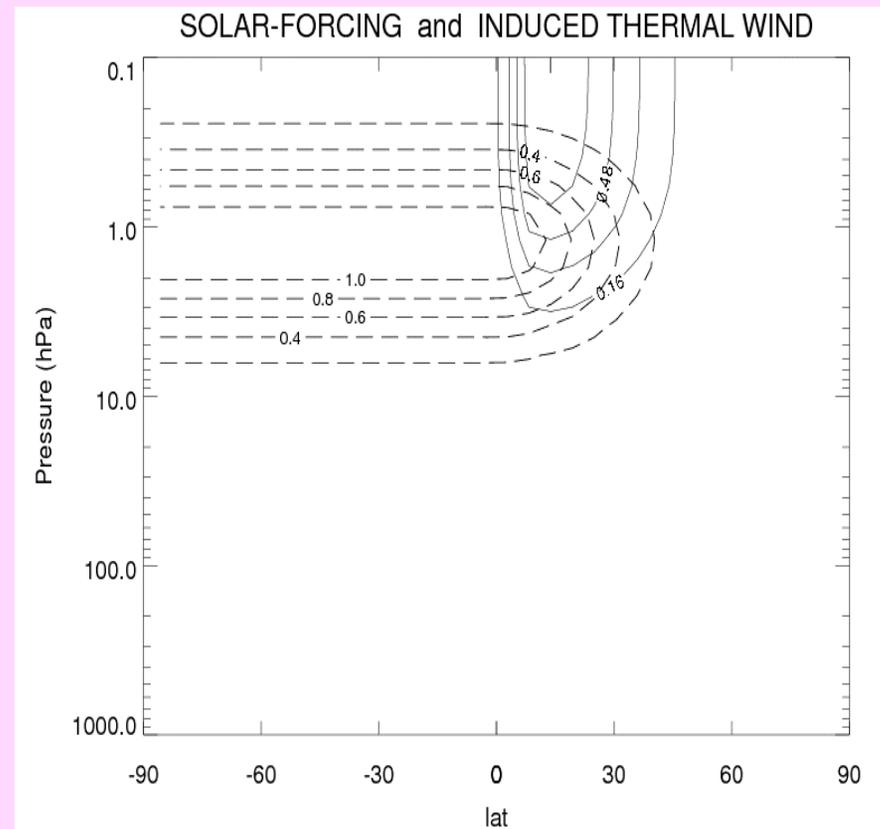
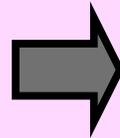
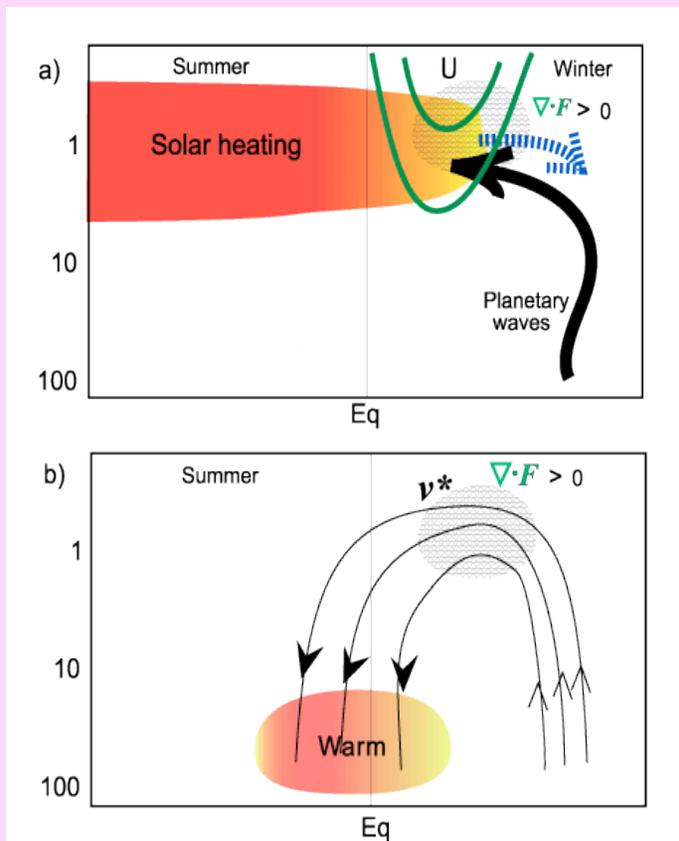
## ❖ Experimental design

- equatorial QBO

- identical to **Naito and Yoden(2006)**
- WWWW and EEEE

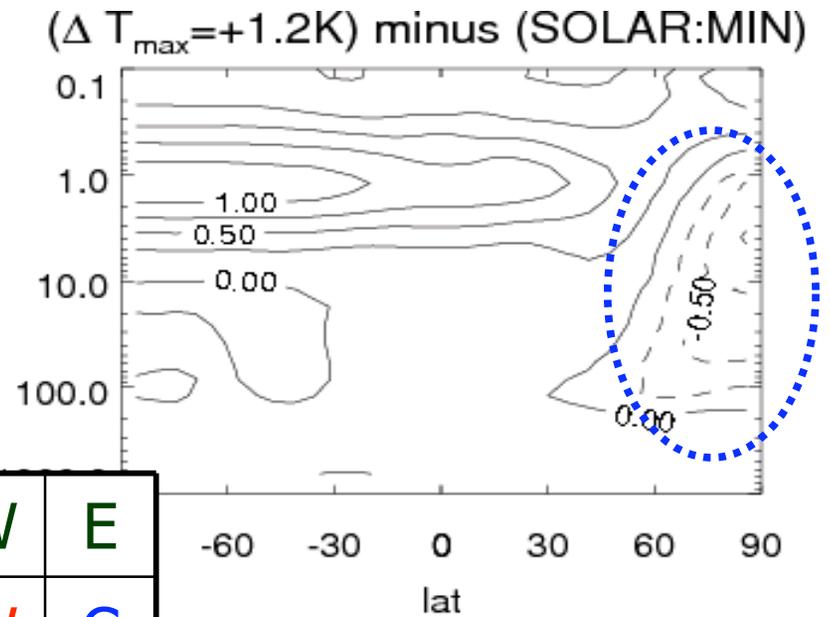
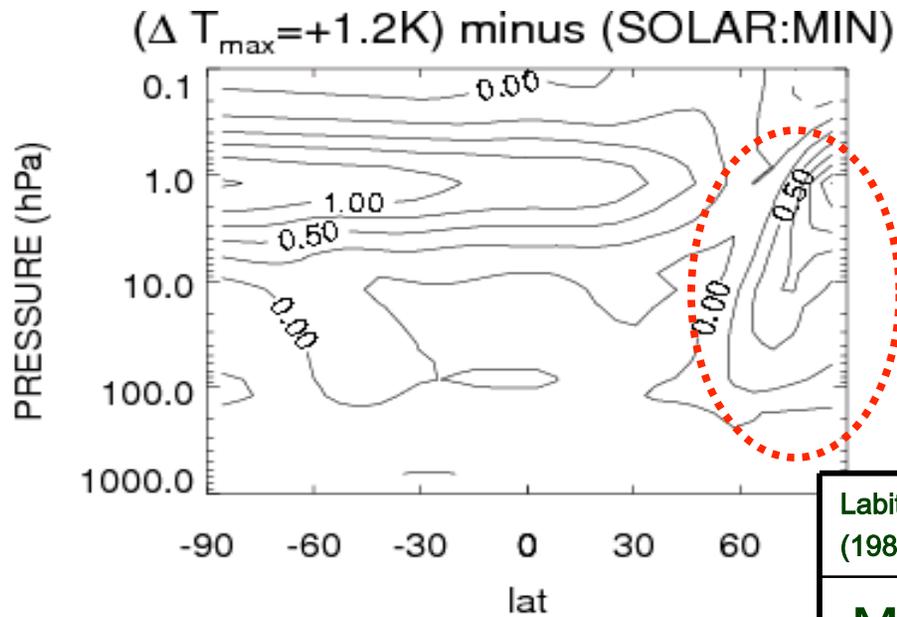
- solar heating

- Kodera and Kuroda(2002)

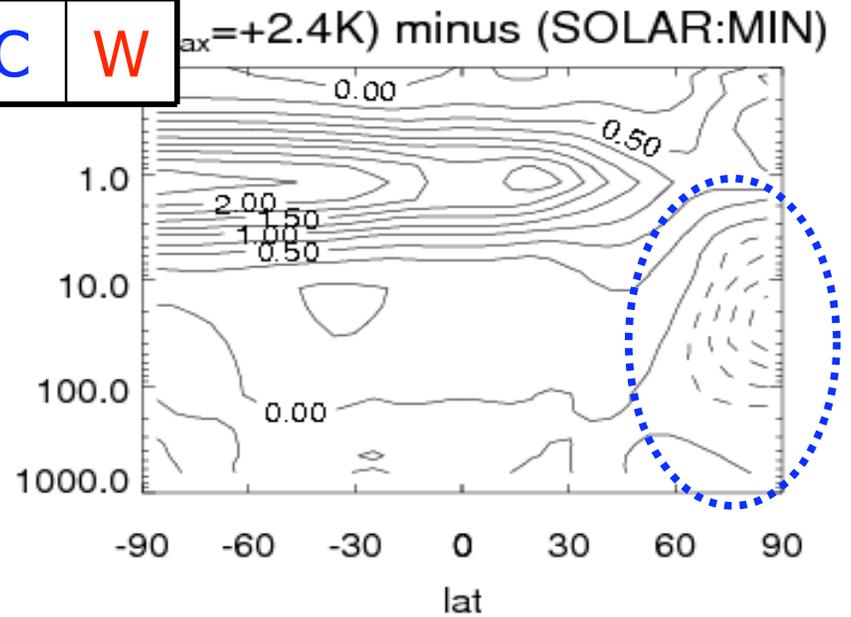
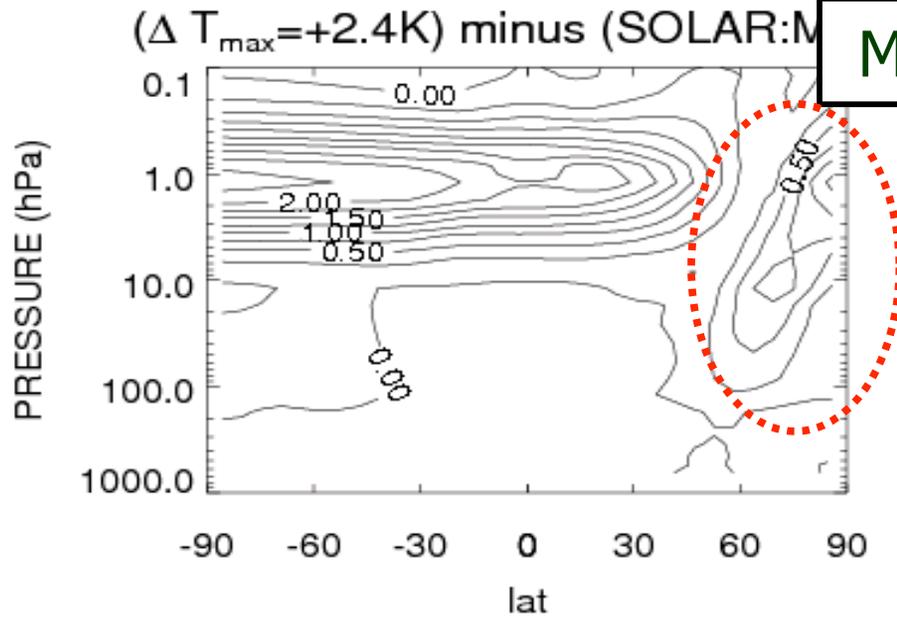


## QBO Westerly

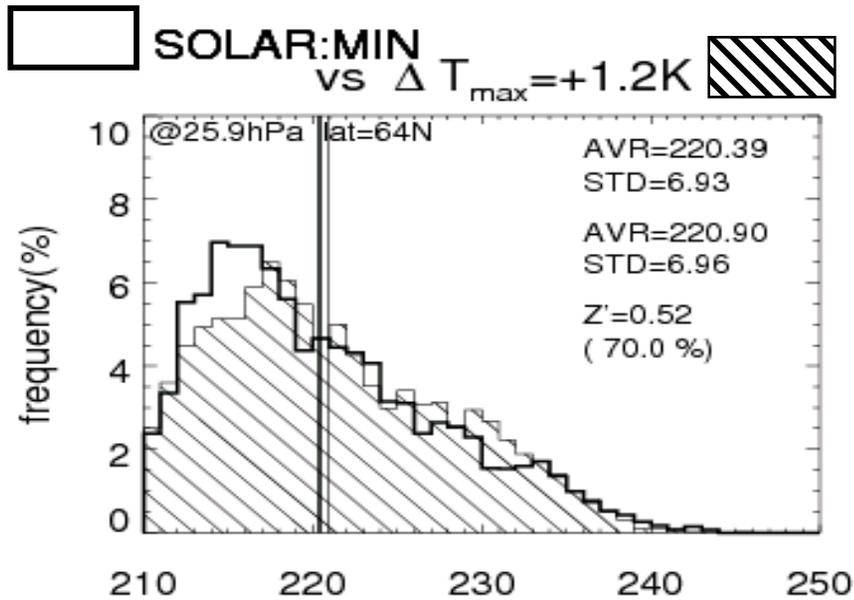
## QBO Easterly



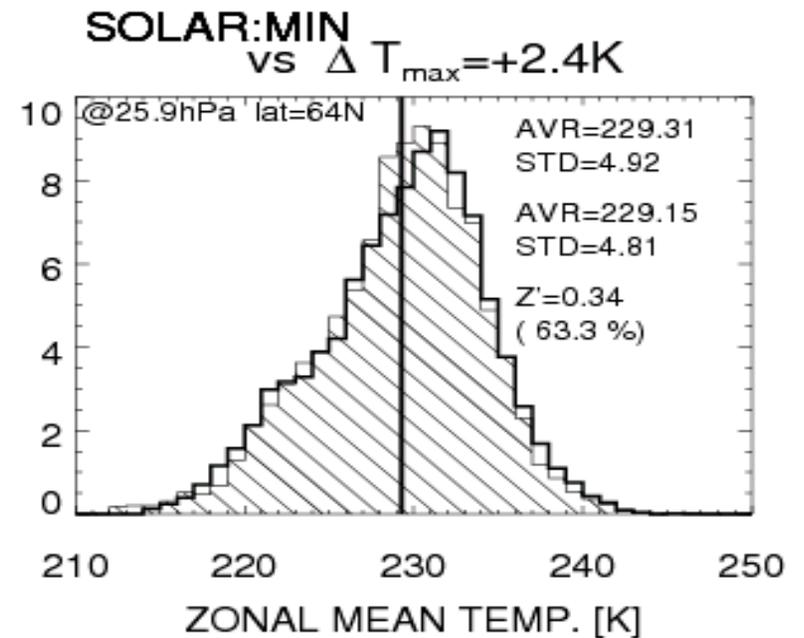
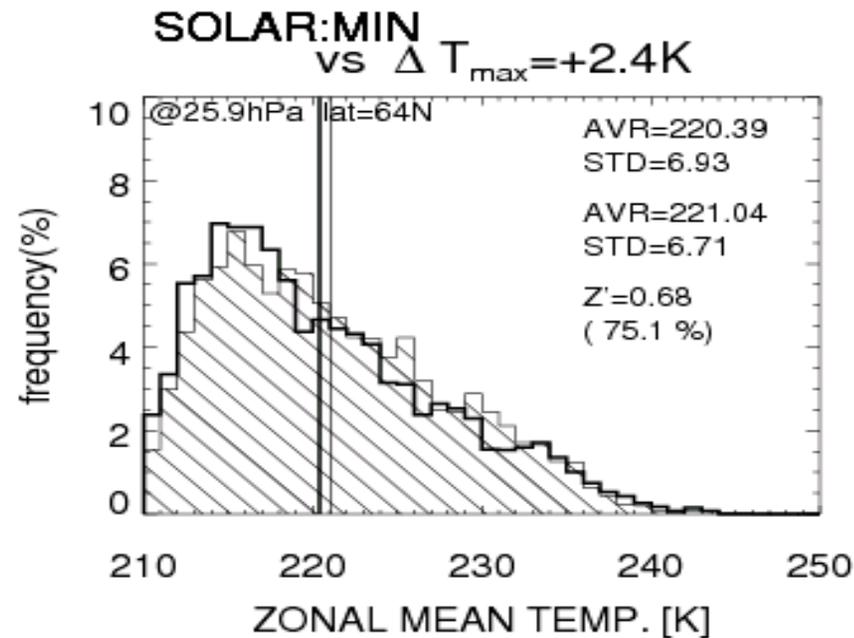
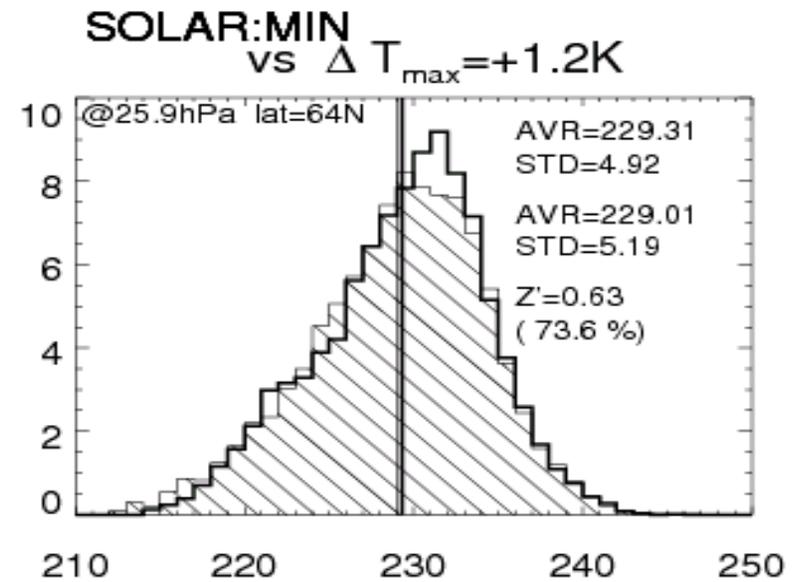
Labitzke (1987)	W	E
MA	W	C
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## QBO Westerly

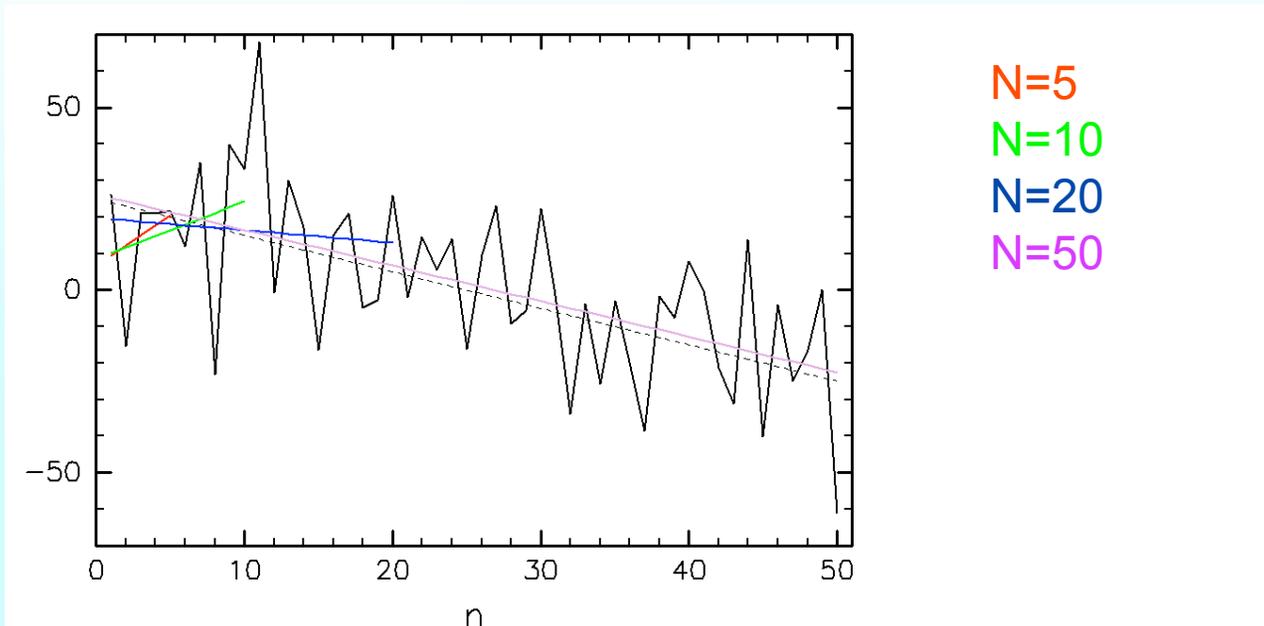


## QBO Easterly



## ❖ Spurious trend may exist in a finite-length dataset

- natural variability



- long period variations of external forcing

- periodic forcing: solar 11-year cycle
- intermittent forcing: volcanic eruptions

- gap in quality of data

- change in observation method: Start of satellite obs., ...

## ❖ Previous studies

- Standard deviation of the spurious trend

- Tiao et al. (1990)
- Weatherhead et al. (1998)

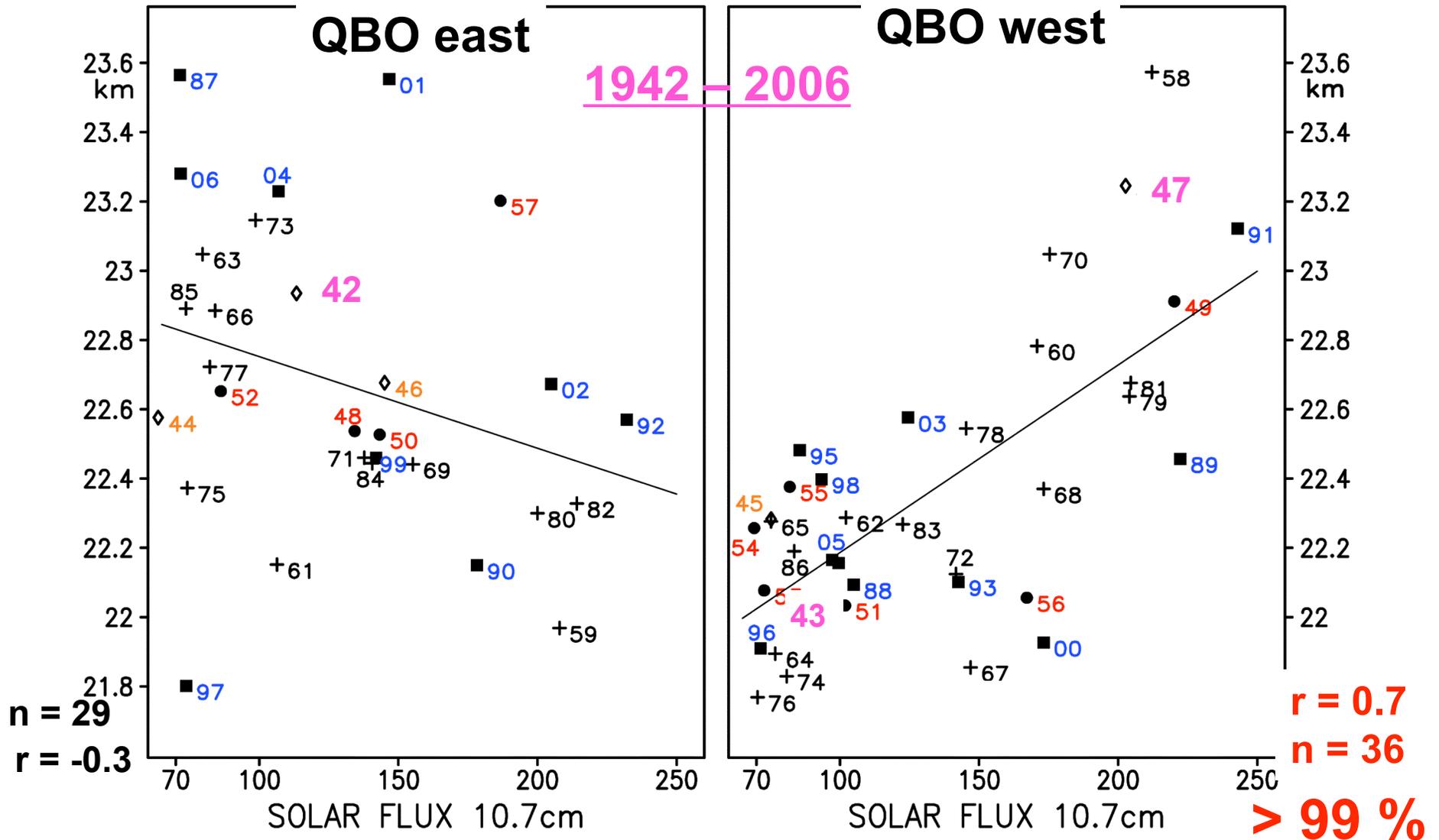
- Student's  $t$ -test for statistical significance of estimated trend

- assumption: the spurious trend has a normal distribution

## ❖ The PDF of the spurious trend depends on the PDF of natural variability

- Some atmospheric natural variations have a non-normal distribution

# North Pole February 30-hPa Heights NCEP/NCAR + REC

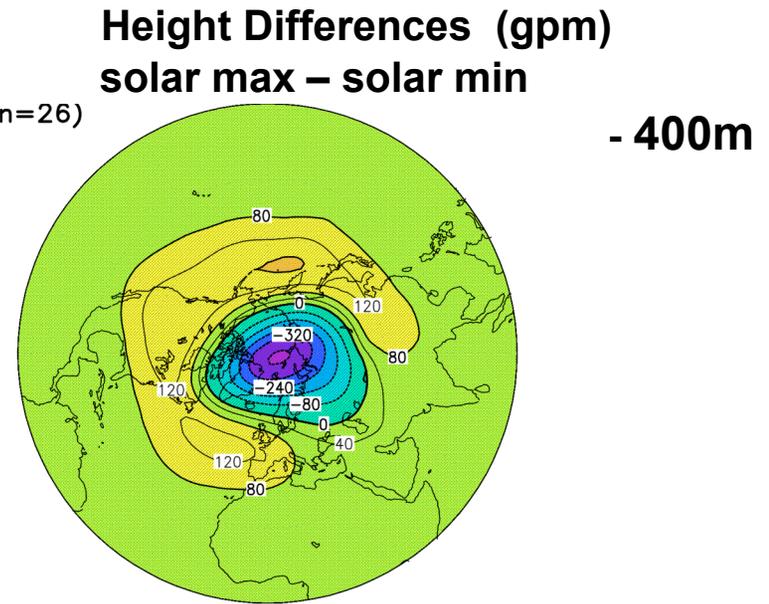
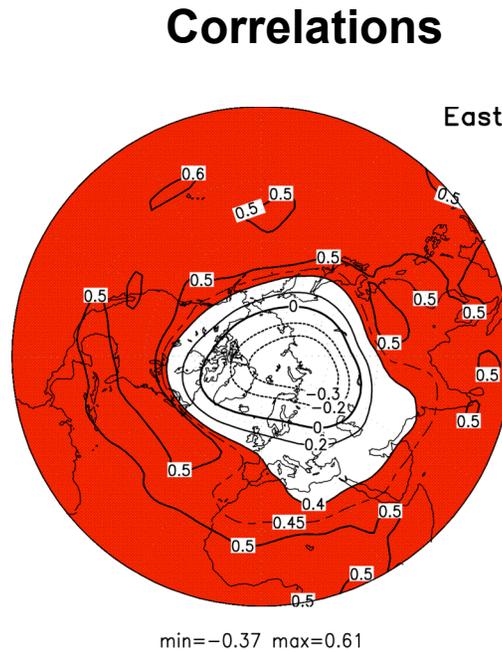


Correlations between 30-hPa Heights and the Solar Flux

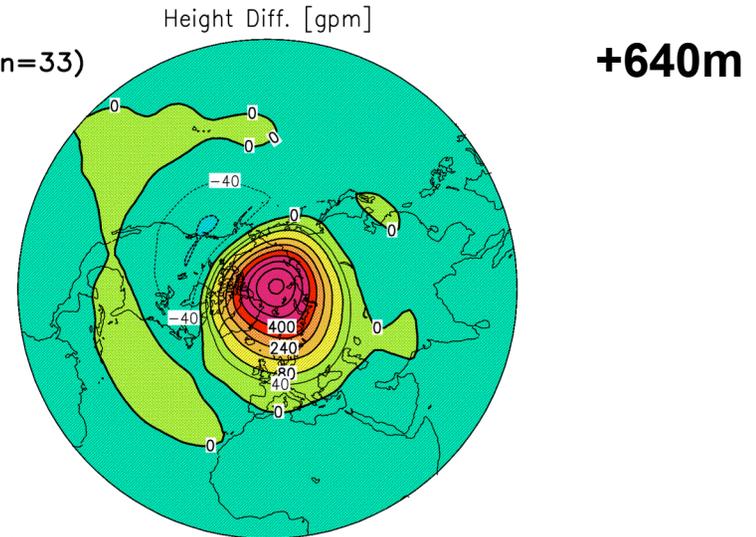
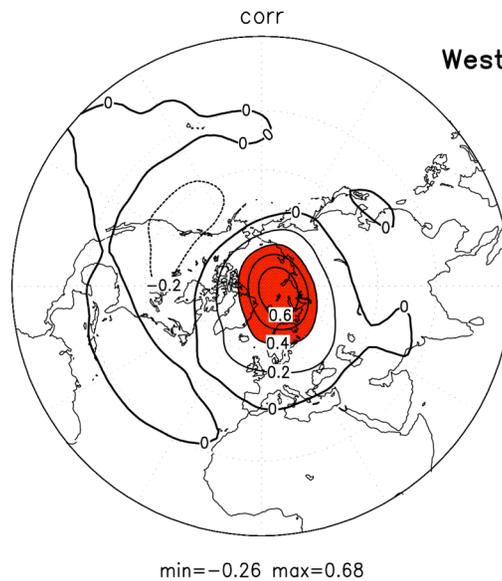
3) 16 years back: 1948 – 1957 (10 y, red); 1942 – 1947 (6 y, orange, REC )

# 30-hPa Heights February (~ 22 – 24 km)

**East**  
 $r_{\text{max}} = 0.61$   
 $n = 26$   
 95%



**West**  
 $r_{\text{max}} = 0.68$   
 $n = 33$   
 99%



(Labitzke et al., 2006)

February 1948 – 2006, NCEP/NCAR, n = 59 years<sup>36</sup>