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LETTER

Evaluating ethane and methane emissions associated with the development of oil and natural gas extraction in North America

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Abstract

Sharp rises in the atmospheric abundance of ethane (C₂H₆) have been detected from 2009 onwards in the Northern Hemisphere as a result of the unprecedented growth in the exploitation of shale gas and tight oil reservoirs in North America. Using time series of C₂H₆ total columns derived from ground-based Fourier transform infrared (FTIR) observations made at five selected Network for the Detection of Atmospheric Composition Change sites, we characterize the recent C₂H₆ evolution and determine growth rates of ~5% yr⁻¹ at mid-latitudes and of ~3% yr⁻¹ at remote sites. Results from CAM-chem simulations with the Hemispheric Transport of Air Pollutants, Phase II bottom-up inventory for anthropogenic emissions are found to greatly underestimate the current C₂H₆ abundances. Doubling global emissions is required to reconcile the simulations and the observations prior to 2009. We further estimate that North American anthropogenic C₂H₆ emissions have increased from 1.6 Tg yr⁻¹ in 2008 to 2.8 Tg yr⁻¹ in 2014, i.e. by 75% over these six years. We also completed a second simulation with new top-down emissions of C₂H₆ from North American oil and gas activities, biofuel consumption and biomass burning, inferred from space-borne observations of methane (CH₄) from Greenhouse Gases Observing SATellite. In this simulation, GEOS-Chem is able to reproduce FTIR measurements at the mid-latitudinal sites, underscoring the impact of the North American oil and gas development on the current C₂H₆ abundance. Finally we estimate that the North American oil and gas emissions of CH₄, a major greenhouse gas, grew from 20 to 35 Tg yr⁻¹ over the period 2008–2014, in association with the recent C₂H₆ rise.

1. Introduction

Ethane (C₂H₆) is a ubiquitous constituent of the Earth's atmosphere, with surface concentrations typically ranging from 500 to 2200 ppt over the Northern Hemisphere (Simpson *et al* 2012). As a result of its relatively short lifetime against oxidation by the main tropospheric oxidant, the hydroxyl radical (OH; Rudolph 1995), C₂H₆ presents a significant seasonal modulation, characterized by a winter maximum and a summer minimum. The main C₂H₆ sources are of

anthropogenic origin (62% from leakage during production and transport of natural gas (NG), 20% from biofuel combustion and 18% from biomass burning; e.g., Xiao *et al* 2008), and this produces a strong latitudinal and interhemispheric gradient (Simpson *et al* 2012).

Ethane plays a number of roles in atmospheric chemistry. As the most abundant non-methane volatile organic compound, its oxidation in the presence of nitrogen oxides (NO_x) can enhance tropospheric ozone (O₃) production, especially in polluted air

masses, and it is a significant source of peroxyacetyl nitrate, the main tropospheric reservoir species for NO_x (Fischer *et al* 2014). Ethane directly impacts the oxidizing capacity of the troposphere by serving as an important sink for OH. It also shares important and concurrent anthropogenic emission sources with methane (CH_4), a major greenhouse gas released to the atmosphere by numerous natural processes and human-related activities. A complete understanding of the atmospheric distribution of C_2H_6 and any trends in the abundance of this species can be used to better constrain the sources of CH_4 , in particular from oil and gas activities.

Until very recently, the abundance of C_2H_6 in the atmosphere has been declining, and this has been mainly attributed to the reduction of fugitive emissions from its fossil fuel sources (Simpson *et al* 2012) as a result of successful implementation of measures aimed at the reduction of atmospheric pollution. Simpson *et al* (2012) showed that global emissions dropped from 14.3 to 11.3 Tg yr^{-1} over the period 1984–2010. Several studies have characterized the atmospheric decrease of C_2H_6 over the last two decades, and found consistent relative trends in the -1 to $-2.7\% \text{ yr}^{-1}$ range, depending on the site and time period (Franco *et al* 2015, and references therein).

There are multiple lines of evidence that the multi-decade decline in C_2H_6 has ended. Using surface measurements of the Photochemical Assessment Monitoring Stations network, Vinciguerra *et al* (2015) showed that C_2H_6 mixing ratios increased after 2010 at locations downwind of active wells in North America. Almost simultaneously, Franco *et al* (2015) reported a $5\% \text{ yr}^{-1}$ increase in the total C_2H_6 column series after 2009 over the remote high-altitude site of the Jungfrauoch in the Swiss Alps, suggesting that emissions associated with hydraulic fracturing and shale gas operations in North America are affecting Europe.

NG production over the US increased by 40% from 1989 to 2013 (Energy Information Administration, <http://eia.gov/dnav/ng/hist/n9010us2a.htm>). The rate of change of NG production increased in 2006, resulting in an annual production of 835 000 million cubic meters by the year 2013 (figure 1(a)). Gas producing wells are located in many regions of the US, the largest number being located in the northeast and the south central portion of the US (figure 1(b); Energy Information Administration, https://eia.gov/dnav/ng/ng_sum_lsum_a_EPG0_xdg_count_a.htm).

Here we use remote-sensing observations from five ground-based Fourier transform infrared (FTIR) instruments over the 2003–2014 time period to characterize the recent C_2H_6 abundance increases over North America that we attribute to the intense development of oil and gas extraction in this region. We show that a significant adjustment of the anthropogenic C_2H_6 emissions (dominated by the oil and gas sector) available from the most current bottom-up emission inventory is needed to allow a correct

reproduction of the C_2H_6 atmospheric levels and recent rise by the CAM-chem model (the Community Atmosphere Model with Chemistry). These findings are confirmed by an independent simulation performed with the chemical transport model GEOS-Chem, implementing spatially resolved top-down emissions of C_2H_6 (Turner *et al* 2015).

2. Methods

2.1. FTIR and portable atmospheric research interferometric spectrometer for the infrared (PARIS-IR)

High-resolution infrared solar absorption spectra encompassing C_2H_6 absorption features near $3.5 \mu\text{m}$ are regularly recorded under clear-sky conditions at globally distributed ground-based sites within the Network for the Detection of Atmospheric Composition Change (NDACC; see <http://ndacc.org> and www2.acom.ucar.edu/irwg). The present study focuses on observations obtained as part of this monitoring effort in North America, namely at Eureka, Thule and Toronto, and at Mauna Loa. Spectra recorded at Boulder, in operation since 2009 and not affiliated yet to NDACC, have also been included to increase the latitudinal coverage. Table 1 provides information on these sites. To check the ability of lower-resolution, compact and portable instruments to provide meaningful column measurements of C_2H_6 , we further considered observations performed by the portable atmospheric research interferometric spectrometer for the infrared (PARIS-IR; Fu *et al* 2007), an instrument based on the design of the atmospheric chemistry experiment-Fourier transform spectrometer (ACE-FTS) in orbit on-board the SCISAT satellite since August 2003 (Bernath *et al* 2005). PARIS-IR has been operated from Eureka each spring since 2006 and from Toronto in 2009 and 2011, side-by-side with the corresponding affiliated NDACC-FTIR spectrometers.

The C_2H_6 retrieval strategies used here are consistent with those described in Franco *et al* (2015). In brief, two spectral windows encompassing the $^{\text{P}}\text{Q}_3$ and $^{\text{P}}\text{Q}_1$ sub-branches near 2976 and 2983 cm^{-1} were simultaneously fitted, accounting for interferences by water vapor (H_2O), O_3 , methyl chloride (CH_3Cl) and CH_4 . The third optional window near 2986 cm^{-1} , sometimes affected by water vapor, was used at all sites. The adopted spectroscopic parameters correspond to a combination of HITRAN 2008 for most species (Rothman *et al* 2009), with HITRAN 2012 for CH_3Cl (Rothman *et al* 2013), improved parameters for the O_3 lines near 2976.97 cm^{-1} and a pseudo-linelist for C_2H_6 derived from laboratory cross section spectra recorded by Harrison *et al* (2010). Model simulations with the CHemical Atmospheric General Circulation Model for study of atmospheric environment and radiative forcing (CHASER; Sudo *et al* 2002) were used

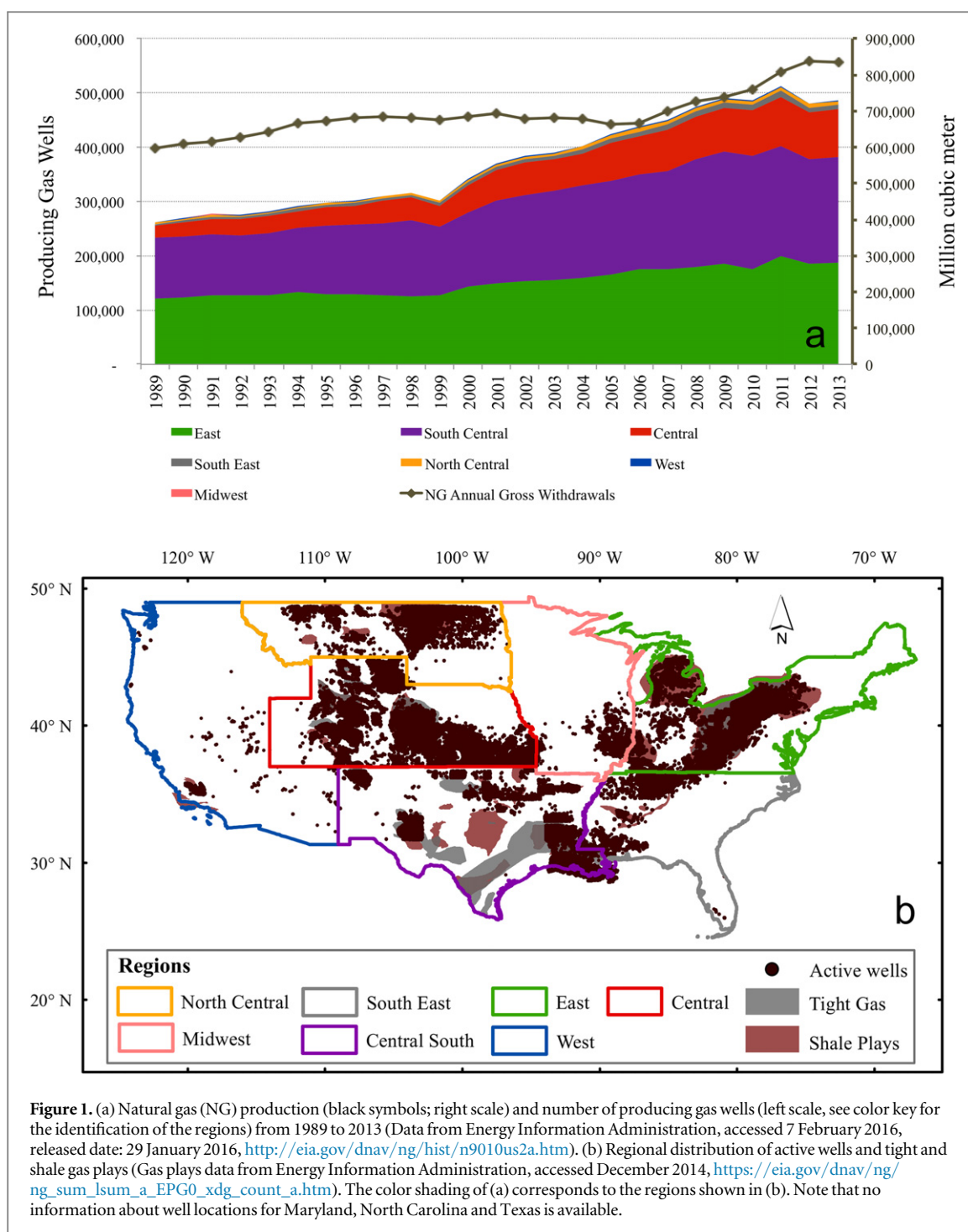


Table 1. Information on the FTIR sites with C_2H_6 measurements presented in this study.

FTIR site	Location	Reference
Eureka	Canada, NU, 80.05° N, 86.42° W, 610 m a.s.l.	Batchelor <i>et al</i> (2009)
Thule	Greenland, 76.52° N, 68.77° W, 225 m a.s.l.	Hannigan <i>et al</i> (2009)
Jungfrauoch	Swiss Alps, 46.55° N, 7.98° E, 3580 m a.s.l.	Zander <i>et al</i> (2008)
Toronto	Canada, ON, 43.66° N, 79.40° W, 174 m a.s.l.	Wiacek <i>et al</i> (2007)
Boulder	United States, CO, 40.40° N, 102.50° W, 1612m a.s.l.	
Mauna Loa	United States, HI, 19.54° N, 155.57° W, 3396 m a.s.l.	Hannigan <i>et al</i> (2009)

to build the *a priori* covariance matrix and to determine on a site-by-site basis, a scale factor applied to the v6 climatological *a priori* vertical profiles employed in

the NDACC Infrared Working Group that were derived from the Whole Atmosphere Community Climate Model (WACCM; Garcia *et al* 2007, Eyring

et al 2013). The objective characterization of the information content indicated that mean degrees of freedom for signal (DOFS) are in the range of 1.5–1.7 at all sites, providing sensitivity in an altitude range from the ground up to ~ 18 km. The retrieval algorithm and approach adopted for the PARIS-IR low-resolution (0.02 cm^{-1}) observations are consistent with those used for the high-resolution spectra and provide a mean DOFS of approximately 1.0 between the ground and ~ 18 km.

Uncertainties associated with the retrieval of C_2H_6 have been evaluated in numerous previous studies (see Franco *et al* 2015 and references therein), for many different sites with a broad range of atmospheric conditions. All concur in estimating a total systematic error on the total columns close to 6%, dominated by the uncertainty affecting the C_2H_6 line parameters. Reported random errors are in the 2%–4% range for studies making simultaneous use of several windows, with main contribution by the measurement noise.

2.2. CAM-chem

For this study, a simplified version of CAM-chem (Lamarque *et al* 2012), a component of the Community Earth System Model, was used to simulate multi-year distributions of C_2H_6 along with tracers of source region contributions. CAM version 4 (CAM4) with bulk aerosols was used at a 1.9° latitude \times 2.5° longitude resolution, with the global OH distribution specified using output from a previous full chemistry simulation for year 2000 (Tilmes *et al* 2015). The meteorology is specified using Modern-Era Retrospective Analysis for Research and Applications meteorological fields corresponding to each year of simulation. The reaction of C_2H_6 with OH was the only chemical loss included. Since this version of the model does not include complete chemistry, it will be called CAM- C_2H_6 throughout this paper.

The initial anthropogenic C_2H_6 emissions inventory used with CAM- C_2H_6 was developed for the Hemispheric Transport of Air Pollutants, Phase II (HTAP2), and is a composite of regional inventories harmonized to represent 2008 and 2010 (Janssens-Maenhout *et al* 2015). Additional C_2H_6 emissions included in these simulations are biogenic emissions from the online Model of Emissions of Gases and Aerosols from Nature (MEGAN2.1) biogenic emissions model (Guenther *et al* 2012), and fire emissions from the Fire INventory from NCAR (FINNv1.5; Wiedinmyer *et al* 2011).

Daily mean C_2H_6 mixing ratio profiles at the closest pixel to each measurement site were used from the CAM- C_2H_6 output. To ensure a reliable comparison with the ground-based measurements, we imposed the vertical resolution and sensitivity of the FTIR retrievals to the model outputs before computing the total columns: the individual mixing ratio profiles were re-gridded onto the vertical FTIR layer schemes

and convolved by the averaging kernels (AVKs) according to the formalism of Rodgers and Connor (2003). The re-gridding method is a mass-conservative interpolation, preserving the total C_2H_6 mass above the station altitude and ignoring the mass underneath. Seasonally-averaged AVKs were employed for the smoothing.

2.3. GEOS-Chem

GEOS-Chem is a chemical transport model driven by assimilated meteorological fields from the Goddard Earth Observing System of the National Aeronautics and Space Administration (NASA) Global Modelling and Assimilation Office (Bey *et al* 2001). Our analysis was based on a $2^\circ \times 2.5^\circ$ resolution simulation for 2010, with an 18 month spin-up. We used v10-01 with GEOS-5 meteorological fields and the full chemistry (NO_x - O_x -HC-Aer-Br) mechanism. Ethane emission fluxes over the US and Canada were based on CH_4 fluxes derived from the Greenhouse Gases Observing SATellite (GOSAT) at a $50\text{ km} \times 50\text{ km}$ spatial resolution (Turner *et al* 2015). The CH_4 inversions were calculated using a state vector optimally defined with radial basis functions to enable analytical inversion with full error characterization while minimizing aggregation error. A full description of the uncertainty and error characterization of this method for calculating CH_4 emissions is presented in Turner and Jacob (2015).

In order to attribute nationwide percentages of anthropogenic CH_4 emissions, Turner *et al* (2015) used prior information on source locations. As a result, they estimated CH_4 fluxes for 12 anthropogenic categories. In this study, we only considered the categories and thus locations relevant to C_2H_6 , and these include: (1) oil and NG, (2) biofuel, and (3) biomass burning. For these categories, we estimated C_2H_6 emissions using $\text{C}_2\text{H}_6/\text{CH}_4$ emission ratios from the literature and implemented them into the model using the Harvard-NASA Emissions Component v1.1.005 (Keller *et al* 2014). To estimate oil and gas C_2H_6 emission fluxes, we used a single $\text{C}_2\text{H}_6/\text{CH}_4$ emission ratio based on Katzenstein *et al* (2003). To derive biomass burning and biofuel C_2H_6 emissions, we considered biofuel consumption as residential biomass burning; and we used the $\text{C}_2\text{H}_6/\text{CH}_4$ emission ratios for temperate forests from Akagi *et al* (2011). The use of relatively few emission ratios is a large source of uncertainty in this analysis on top of errors associated with the GOSAT based CH_4 fluxes. However, there are relatively few emission ratios available in the published literature, and our approach offers a simple and reproducible implementation in the model. Consistent with the CAM- C_2H_6 outputs, we used the GEOS-Chem mixing ratio profiles accounting for the FTIR vertical resolution and sensitivity.

3. Results

3.1. Observed ethane increase

We investigated C_2H_6 time series derived from the ground-based observations to determine the current abundances of C_2H_6 as well as its recent evolution over the North American FTIR sites. The FTIR daily mean total columns from January 2003 through December 2014 are shown as gray circles in figure 2, while the light blue triangles correspond to the daily-averaged PARIS-IR measurements at Eureka and Toronto. On each frame, the right y -axis scale converts the total columns into approximate surface mole fractions. The scale conversion was determined for each site according to the linear regression between the total column abundance and the C_2H_6 mole fraction in the lowermost layer of the FTIR retrieval scheme. These regressions are characterized by a coefficient of determination (R^2) of at least 0.8.

Differences in C_2H_6 abundance present in figure 2 are largely a function of site latitude (Simpson *et al* 2012) and altitude (Gardiner *et al* 2008, Angelbratt *et al* 2011); the bulk of the C_2H_6 molecules are found in the lowermost tropospheric layers. The mean column for each site ($\pm 1\sigma$) is as follows: 1.7 ± 0.5 (Eureka; 610 m a.s.l.), 1.8 ± 0.6 (Thule; 225 m a.s.l.), 2.1 ± 0.4 (Toronto; 174 m a.s.l.), 1.6 ± 0.5 (Boulder; 1612 m a.s.l.) and 0.6 ± 0.2 molec cm^{-2} (Mauna Loa; 3396 m a.s.l.). There are also seasonal and interannual differences in the C_2H_6 burden between the sites. Mainly driven by the rate of reaction with OH, the typical seasonal peak-to-peak amplitude for C_2H_6 is on the order of 50% at Boulder and Toronto, and it increases to $\sim 80\%$ at the remote sites. It is worth noting the good agreement between the PARIS-IR and FTIR total columns above both Eureka and Toronto (figures 2(a) and (c)).

We characterize the recent atmospheric evolution of C_2H_6 , considering the daily means, with a statistical bootstrap resampling tool that combines a linear function and a Fourier series accounting for the intra-annual variability of the data sets (Gardiner *et al* 2008). A third-order Fourier series is used to adjust the seasonal variations, except at the Arctic sites, Eureka and Thule, where a second-order fit is appropriate to capture the seasonality over the spring to fall time period (no observations during the polar night). As in Franco *et al* (2015), we consider separately the daily mean subsets before 1 January 2009 (when enough observations are available), and after 1 January 2009. The fitted seasonality is drawn as a blue curve in each frame of figure 2, while the green and red solid lines designate the corresponding linear regressions. At Eureka (figure 2(a)), a linear regression (red dashed line) is calculated solely on the basis of the 2009–2014 PARIS-IR campaign measurements. Figure 3 shows the relative annual C_2H_6 rates of change, using either the January 2003 or January 2009 computed total columns as reference. The available annual rates of C_2H_6 over

both the 2003–2008 and 2009–2014 time periods, along with the 2σ confidence level, for the five FTIR sites are plotted in blue. The equivalent annual rates derived from the Jungfraujoch (table 1) time series investigated in Franco *et al* (2015) are added for comparison.

From 2003 to 2008, we observe a slow decline of the C_2H_6 total columns between -1.0 and -1.5% yr^{-1} , with consistent rates within the different latitudes (figure 3(a)). The Mauna Loa rate of change is not significant at the 2σ confidence level ($0.9 \pm 1.1\%$ yr^{-1}). However, when using the dataset from inception, we compute a decrease of $-1.0 \pm 0.3\%$ yr^{-1} over 1995–2008. These results agree with the global decline of the atmospheric C_2H_6 burden from the mid-1980s up to the end of the 2000s, which was widely reported in the scientific literature (e.g., Gardiner *et al* 2008, Angelbratt *et al* 2011, Franco *et al* 2015, and references therein). It was attributed to reduced fugitive fossil fuel emissions that were thought to also contribute significantly to the slowing growth rate of atmospheric CH_4 since the mid-1980s (Aydin *et al* 2011, Simpson *et al* 2012).

However, as at Jungfraujoch (Franco *et al* 2015), the FTIR measurements over the North American sites show a reversal of the long term decline in C_2H_6 around 2009, followed by a sharp and significant (at the 2σ confidence level) increase at all FTIR sites (figure 3(b)). The observed rates of change are very consistent and vary between $\sim 3\%$ yr^{-1} at the remote stations and $\sim 5\%$ yr^{-1} at mid-latitudes. The increase in C_2H_6 deduced from the PARIS-IR and FTIR observations at Eureka are very consistent as well. We hypothesize that this recent increase is the result of the recent massive growth in the exploitation of shale gas and tight oil reservoirs in North America, where the drilling productivity began to grow rapidly after 2009 (Schneising *et al* 2014, Warneke *et al* 2014, Vinciguerra *et al* 2015). Of the observation sites presented here, Boulder is located closest to a highly productive oil and gas basin, the Denver-Julesberg Basin. We suggest that is why the Boulder data show a large variability in the C_2H_6 columns ($5.0 \pm 4.5\%$ yr^{-1}), influenced by local air mass circulation. Similar elevated mixing ratios of C_2H_6 have been observed at the Boulder Atmospheric Observatory by Gilman *et al* (2013).

3.2. Results from HTAP2 emissions inventory

We compared C_2H_6 simulations produced with CAM- C_2H_6 for the period 2003–2014, using the bottom-up anthropogenic inventory HTAP2, with the remote-sensing measurements. The C_2H_6 total columns derived from this simulation, represented by the dashed curve in figure 2, greatly underestimate the FTIR observations at all sites in terms of C_2H_6 abundance and seasonal amplitude, with a low bias that is far larger than the typical systematic uncertainty

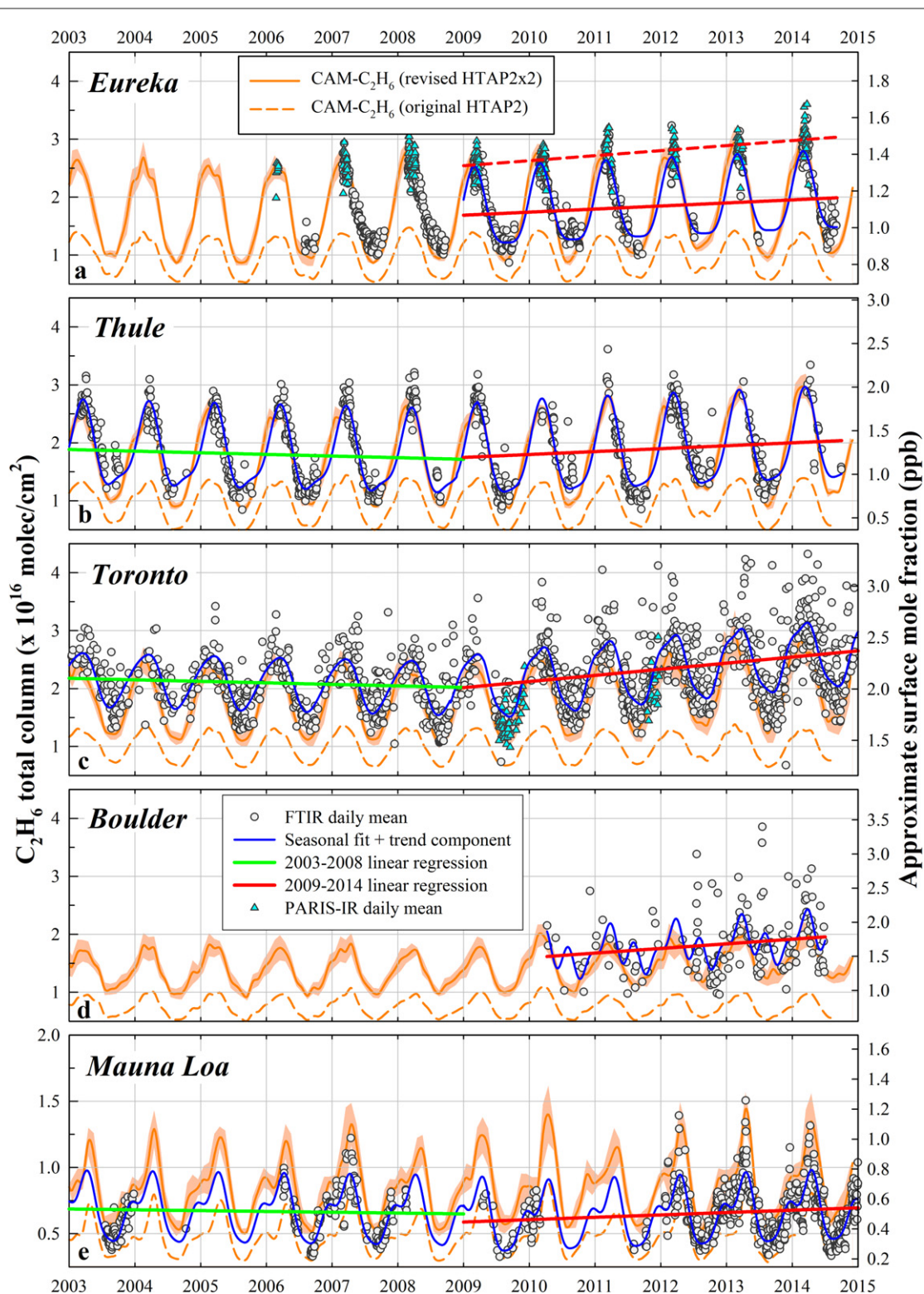


Figure 2. Daily mean C_2H_6 total columns derived from the FTIR (gray circles) and PARIS-IR (light blue triangles) observations performed between January 2003 and December 2014. The right y-axis scale converts the total columns into approximate surface mole fraction (see text for details). The blue curve visualizes the function (including seasonal modulation and trend component) fitted to all daily FTIR means over the periods 2003–2008 and 2009–2014, using a bootstrap tool (Gardiner *et al* 2008). The green and red lines are the associated linear regressions (as solid line for FTIR and dashed line for PARIS-IR). The dashed and solid orange curves are the monthly mean C_2H_6 total columns simulated by CAM- C_2H_6 , implementing the original HTAP2 and revised HTAP2x2 (with increasing North American emissions) inventories, respectively. The shaded area corresponds to the 1σ standard deviation.

of 6% affecting the total columns. Furthermore, the model does not reproduce the recent C_2H_6 increase.

Based on comparisons between the model results and the FTIR measurements, we estimate that the

global HTAP2 anthropogenic C_2H_6 emissions must be doubled to simulate the observed C_2H_6 abundance prior to 2009. We refer to this quantity of C_2H_6 emissions as HTAP2x2. An additional increase in C_2H_6

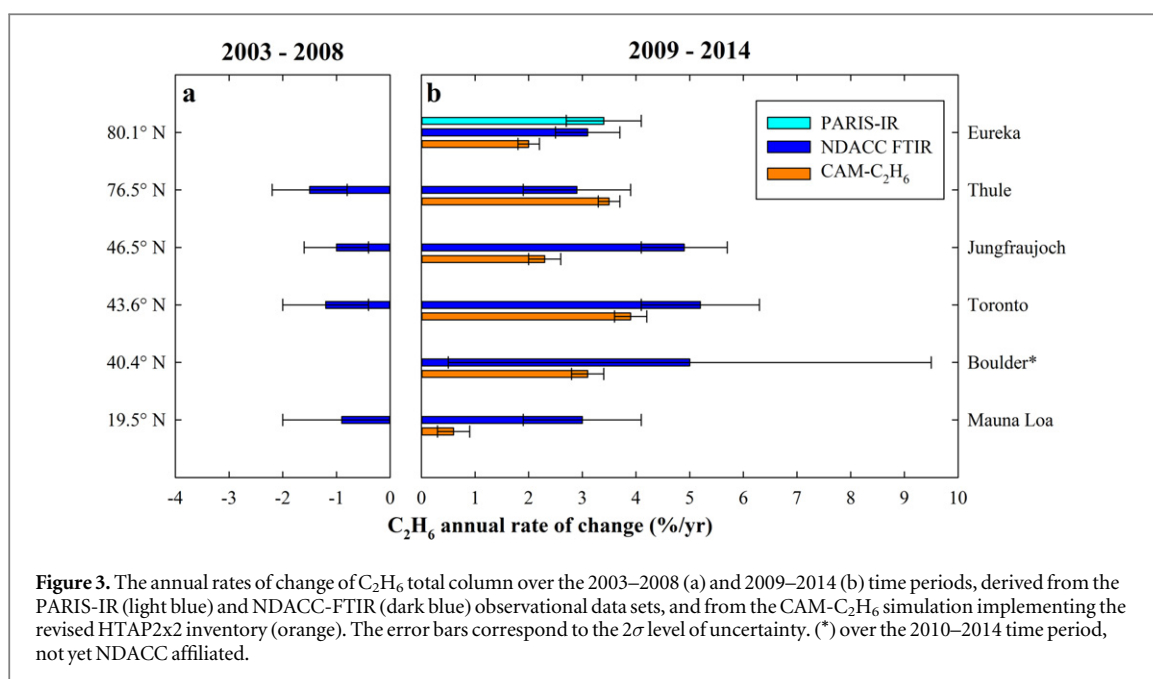


Figure 3. The annual rates of change of C_2H_6 total column over the 2003–2008 (a) and 2009–2014 (b) time periods, derived from the PARIS-IR (light blue) and NDACC-FTIR (dark blue) observational data sets, and from the CAM- C_2H_6 simulation implementing the revised HTAP2x2 inventory (orange). The error bars correspond to the 2σ level of uncertainty. (*) over the 2010–2014 time period, not yet NDACC affiliated.

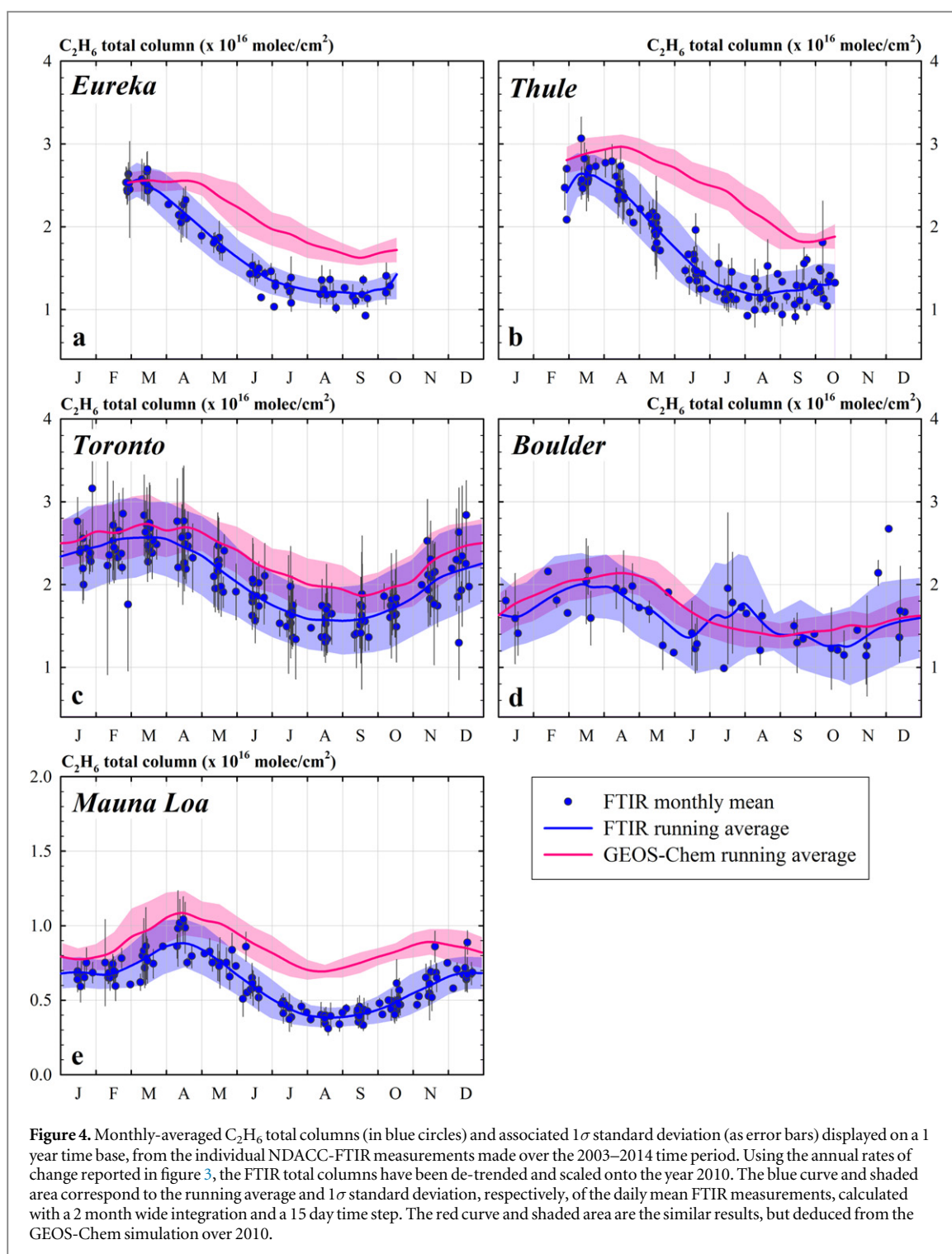
Table 2. Total annual emissions of C_2H_6 (in $Tg\ yr^{-1}$) used for the CAM- C_2H_6 (with the original HTAP2 and revised HTAP2x2 inventories) and GEOS-Chem v10-01 simulations, according to regional/sector sources. The original HTAP2 anthropogenic emissions for 2008 are used for all years, while in the revised HTAP2x2 the global anthropogenic emissions for 2008 were doubled for all years prior to 2009, with increasing North American emissions after 2009.

Region—sector	CAM- C_2H_6 (revised HTAP2x2)				GEOS-Chem
	CAM- C_2H_6 (original HTAP2) 2008–2014	2008	2010	2014	2010
Globe—all sectors	9.7–10.2	17.3	17.9	18.7	13.2
Globe—anthropogenic	7.5	15.0	15.3	16.2	10.5
Globe—biomass burning	1.8–2.3	1.9	2.2	2.2	2.7
Globe—biogenic	0.4	0.4	0.4	0.4	Not included
North America—anthropogenic	0.8	1.6	1.9	2.8	1.8

emissions beyond HTAP2x2 is required to simulate the recent rise from 2009 to 2014. Given that C_2H_6 emissions related to the oil and gas sector represent up to 80% of the total anthropogenic C_2H_6 emissions over North America, and if we assume that the missing emissions during this period resulted from the recent increase in oil and gas extraction in North America, we can use the model to estimate the increase in C_2H_6 emissions needed to match the observations. To match the 2014 observations at the continental sites, the North American annual contribution has to almost double over the period 2009–2014. This corresponds to a 1.2 Tg increase of the North American annual C_2H_6 emissions in 2014 over the 2008 emission rates. An overview of the total C_2H_6 emission rates used for the CAM- C_2H_6 simulations is provided in table 2.

The solid orange line in figure 2 signifies the C_2H_6 total columns simulated by CAM- C_2H_6 with the revised HTAP2x2 inventory and North American emission increases from 2009 onwards. Similar to the FTIR time series, the annual rates of change relative to January 2009 derived from this new CAM- C_2H_6

simulation are calculated and reported in figure 3. The results show a better agreement with the remote-sensing time series for both the abundance of C_2H_6 and its seasonal amplitude, especially at the northern sites. The mean daily biases of the model to the FTIR fitted seasonality over the period 2009–2014 are as follows: $-4 \pm 16\%$ (Eureka), $-4 \pm 14\%$ (Thule), $-18 \pm 13\%$ (Toronto), $-15 \pm 15\%$ (Boulder) and $+37 \pm 18\%$ (Mauna Loa). Globally, the revised HTAP2x2 inventory allows the model to simulate significant C_2H_6 increases between ~ 2.5 and $\sim 4\%\ yr^{-1}$ from 2009 onwards, although the model still slightly underestimates the rates of change observed at mid-latitudes. Similar findings are deduced from the comparison between CAM- C_2H_6 and FTIR at Jungfrauoch. These results support the hypothesis that the enhanced North American oil and gas extraction largely contributes to the recent C_2H_6 increase detected, as assumed previously by Franco *et al* (2015). For the remote non-continental Mauna Loa site, the CAM- C_2H_6 results overestimate the FTIR data (figure 2(e)), suggesting that the emissions in Asia (more directly



upwind of Mauna Loa) should not be increased as much as 2 times the original HTAP2 inventory.

3.3. Results from GOSAT-based inventory

In this section, we compare the results of GEOS-Chem simulations performed with a new top-down inventory of anthropogenic C_2H_6 emissions, based on CH_4 measurements from GOSAT, to the FTIR observations. Since this new inventory is only for 2010, we compare the simulated C_2H_6 columns to the FTIR measurements for this year only (figure 4). In order to

provide more robust statistics for the representation of the C_2H_6 seasonal cycle at each site, the available daily mean columns from the 2003–2014 FTIR time series are de-trended and scaled to the year 2010 according to the annual rates of change reported in figure 3. In figure 4, the FTIR monthly means are represented as blue circles along with 1σ standard deviation as error bars, while the blue curve and the shaded area correspond to the running average of the daily mean total columns (with a 2 month wide integration and a 15 day time step) and the associated 1σ standard

deviation, respectively. Similarly, the red curve and shaded area designate the GEOS-Chem results. The total C_2H_6 emissions provided from the top-down inventory for the year 2010 are reported in table 2.

GEOS-Chem properly simulates the late winter maximum abundance of C_2H_6 (figure 4) and reproduces the FTIR total columns found at the two mid-latitude sites, Toronto and Boulder (figures 4(c) and (d)). Both stations are located in the vicinity of intense oil and NG basins: several studies have shown a volatile organic compound source signature associated with the oil and gas industry near Boulder (Pétron *et al* 2012, Gilman *et al* 2013, Swarthout *et al* 2013, Thompson *et al* 2014), while Toronto is located north of one of the most active basins in the US, the Marcellus shale gas play. Due to its relatively coarse horizontal resolution, the model dilutes local enhancements in C_2H_6 and hence does not simulate the secondary maximum observed at Boulder during summertime (which is likely due to local air mass circulations).

GEOS-Chem shows a high-bias for the summertime C_2H_6 total columns at most sites, especially the remote sites (Eureka, Thule and Mauna Loa). The levels of OH, the main tropospheric sink of C_2H_6 , are globally lower by $\sim 10\%$ – 15% in GEOS-Chem v10-01 with respect to previous model versions. This leads to less C_2H_6 oxidation at the time of OH maximum and to an underestimation of the seasonal cycle. This is particularly relevant for the polar sites, where the lower insolation dampens further the photochemical oxidation, causing C_2H_6 to accumulate in summer. We cannot rule out a contribution from transport; however, direct transport of pollutants into the arctic is less frequent during summer (Klonecki *et al* 2003).

The assumptions with regard to the C_2H_6/CH_4 ratios could also induce uncertainties in the inferred C_2H_6 emissions and could contribute to the biases observed between GEOS-Chem and some FTIR results. Indeed, such ratios largely vary in space and time, e.g., according to the types of facilities and the composition of the NG (Mitchel *et al* 2015, Subramanian *et al* 2015).

4. Discussion and conclusions

We used NDACC-FTIR and PARIS-IR time series from five sites to characterize the recent changes in the abundance of C_2H_6 over North America, and we determined significant annual growth rates of 3% – 5% yr^{-1} over 2009–2014, depending on the latitude. These results corroborate the increase observed in the European remote atmosphere and attributed to the development of oil and NG extraction in North America (Franco *et al* 2015). Using customized simulations of CAM-chem, we evaluated the magnitude of the anthropogenic C_2H_6 emissions required to simulate the observed abundance of C_2H_6 prior to

2009 and its subsequent rise. First the simulations indicate that a doubling of the global HTAP2 anthropogenic emissions is necessary to match the C_2H_6 measurements prior to 2009. Furthermore, an additional increase of the North American anthropogenic C_2H_6 emissions—dominated by up to 80% by emissions from the oil and gas sector—from 1.6 Tg yr^{-1} in 2008 to 2.8 Tg yr^{-1} in 2014, i.e. by 75%, is needed to capture the recent C_2H_6 atmospheric rise. In an upcoming analysis, it is our aim to refine the source attribution and identification of the geographical origin of the missing emissions with the help of C_2H_6 -tagged simulations and more than twenty NDACC-FTIR time series from various locations around the globe.

Simulations of GEOS-Chem v10-01 were run using C_2H_6 emissions over North America derived from space-borne CH_4 observations from GOSAT (Turner *et al* 2015). By using a single C_2H_6/CH_4 emission ratio for each source category, we estimated C_2H_6 emissions from oil and gas activities, biofuel consumption and biomass burning. Results from GEOS-Chem implementing this new top-down inventory agreed well with the FTIR measurements at the mid-latitude sites, close to regions with high drilling productivity, but overestimated the summertime minimum at remote sites. These independent results contribute to underscoring the impact of large emissions associated with hydraulic fracturing and shale gas operations in North America on the enhanced C_2H_6 abundance in the atmosphere. It is worth noting the good agreement between the emission rates derived for the year 2010 from the CAM- C_2H_6 and GEOS-Chem model investigations (table 2), with the top-down approach offering the opportunity to allocate the C_2H_6 emissions on the basis of measurements, for a more accurate representation of recent changes in the atmospheric composition.

Ethane shares important anthropogenic sources with CH_4 , the second most important anthropogenic greenhouse gas, of which the growth rate in the atmosphere has gone on the rise again since the mid-2000s (e.g., Kirschke *et al* 2013, Nisbet *et al* 2014). Using C_2H_6 as a proxy can help to constrain the budget of CH_4 for the emissions originating from the recent massive growth in the North American oil and gas industries. If we apply the C_2H_6/CH_4 ratios used to derive the C_2H_6 emissions from the GOSAT measurements, we estimate that the CH_4 emissions from oil and gas activities in North America have increased from 20 Tg yr^{-1} in 2008 to 35 Tg yr^{-1} in 2014, in association with the recent rise of the annual C_2H_6 emissions.

Pursuing atmospheric monitoring activities is of primary importance for evaluating the impacts of the exploitation of shale gas and tight oil reservoirs on greenhouse gas emissions and air quality degradation, especially now that many countries worldwide, e.g., in Europe, are investigating the possibility of starting

shale gas exploitation within their own territory in the near future.

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