Antarctic Bottom Water formation and variability of Antarctic water masses

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Courtesy of Nick Owens, BAS

Formation of Antarctic Bottom Water : where and how

Changes to the source water for Antarctic Bottom Water

What might we predict will happen to Antarctic Bottom Water?

How is Antarctic Bottom Water changing?





The Southern Ocean is the source of cold dense water at the bottom of most of the world ocean.



WOCE Southern Ocean Atlas



Elements of the Southern Ocean overturning circulation



Antarctic Circumpolar Current

Transport across 30°S from Observations and Climate models



Slide courtesy of Stephanie Downes, Princeton

Recipe for Antarctic Bottom Water

Ingredients:

- A ready supply of source water ('Warm Deep Water')
- A broad continental shelf
- A cold atmosphere
- An abundance of ice shelves

Instructions:

- •Take Warm Deep Water
- Freshen it by adding ice melt
- Transport it westward to regions of wide continental shelf
- Cool it by latent and sensible heat fluxes in polynyas
- · Season to taste with ice shelf meltwater
- Keep cooling until sea ice forms and releases brine to make it salty again
- When it's dense enough to spill off the shelf by itself, it's ready!

But.....many climate models form Antarctic Bottom Water by deep convection in the open ocean!



Riffenburgh, 269-272.



All these processes are too small-scale to be represented in climate models and so must be parameterised.



Properties of 28.27 kg m⁻³ neutral density layer: Antarctic Bottom Water

WOCE Southern Ocean atlas

Synoptic Antarctic Shelf-Slope Interactions (SASSI)

Project for the International Polar Year coordinated by iAnZone

Hydrographic sections ('hedgehog')

Moorings on continental shelf and slope

Quasi-circumpolar

We do now have datasets to compare with climate models.



Warming at 900 m depth in the Southern Ocean



Historical data, prior to 1990



Freely drifting floats, after 1990

Floats minus hydrography, showing warming in Antarctic Circumpolar Current

0.0

30°N

150 W

-0.5

С

0°n

120°W

-1.0

M.06

30°E

150°E

0.5

60°E

1205

°C

1.0

90°Е

Gille (2002) Science.





Some Speculation:

Warmer source water (Warm Deep Water) might lead to:

(i)Overspilling water entraining warmer-than-usual water above it, so exported water might be warmer and less dense?

(ii)Warmer waters penetrating onto the continental shelves and melting more ice?

So shelf water is fresher, so is the water spilling off the shelf colder to compensate, keeping the same density? Or is the outflow water less dense?



Impact of changing atmospheric cyclonic forcing on Weddell Gyre intensity and the consequences for the export of AABW across the South Scotia Ridge.

Steepening (A) or slumping (B) of isopycnal surfaces in the Weddell Sea in response to gyre spinup/down.

Stronger Weddell Gyre leads to less dense (warmer) WSDW crossing the ridge.

Weaker Weddell Gyre leads to denser (colder) WSDW exported.



The Weddell gyre circulation and the area of observations carried out during WECCON

Water mass and temperature distributions at the Greenwich meridian



Mean Pot.Temperature in WDW Mean Salinity in WDW 0.4 34.683 0.39 a 34.682 b 0.38 0.37 0.36 0.35 0.35 **WDW** temperature 34.681 Salinity increased until 34.68 1996 and 34.679 34.678 0.34 decreased until 0.33 34.677 2005 and is 1985 1990 1995 2000 2005 1985 1990 1995 2000 2005 Time [Years] Time [Years] increasing since. Mean Pot.Temperature in WSDW Mean Salinity in WSDW -0.37 34.662 C d -0.375 34.661 pot.Temperature [°C] Salinity Courtesy of -0.38 34.66 Eberhard -0.385 34.659 **WSBW** Fahrhach, -0.39 34.658 1985 1990 1995 2000 2005 1985 1990 1995 2000 2005 temperature is Time [Years] Time [Years] AWI increasing since Mean Salinity in WSBW Mean Pot.Temperature in WSBW -0.765 34.652 1992. -0.77 е 34.651 pot.Temperature [°C] -0.775 34.65 **Atiuiles** 34.649 34.648 -0.78 -0.785 -0.79 34.647 **Potential** -0.795 34.646 **Salinity** -0.8 34.645 temperature 1990 2000 1985 2005 1985 1995 2005 1990 1995 2000 Time [Years] Time [Years]

Mean temperature and salinity in the deep water masses of the Weddell gyre at the Greenwich meridian

Long term freshening of Adelie Land Bottom Water : the variety of Antarctic Bottom Water found in the Australian-Antarctic Basin.



Aoki et al., 2005, GRL.



What is causing any freshening and cooling?





Causes significant zonal mean ocean temperature anomalies (°C).

Heat that is normally vented to the atmosphere is trapped in the ocean.

Strength of the Antarctic overturning circulation over 100 year run (without freshwater dump, with freshwater dump)

Do we care about increased freshwater? Climate models suggest we should.

Effects of perturbation experiment in HadCM3, adding surface freshwater layer that prevents convection.

(Richardson et al., GRL, 2005)





Surface air temperature differences, 6-10 years after dumping freshwater around Antarctica.

Richardson et al., GRL, 2005

Climate impacts

Anomalies in Antarctic sea ice thickness (m), 6-10 years after dumping freshwater around Antarctica. Dashed contour shows maximum extent in model normally.





Conclusions

The processes driving Antarctic Bottom Water are complex, and generally too local to be represented in climate models – we need better parameterisations.

The source waters for Antarctic Bottom Water seem to have warmed in recent decades, and the waters on the Antarctic continental shelf have freshened.

Close to Antarctica, Antarctic Bottom Water has in many places cooled and/or freshened.

Until we better understand the processes, it is difficult to predict what will happen to the Antarctic Bottom Water overturning cell.