

# A new diagnostic for evaluating transport between the tropical upper troposphere and mid-latitudes lower stratosphere in chemistry-transport models

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## Introduction

Horizontal transport from the equator to mid-latitude summer is one of the major routes of stratosphere-troposphere exchange. We classified data by the tropopause height and found observational evidence of the tropical air intrusion at mid-latitudes. As effective methods for diagnosing this transport process in chemistry-transport models, we examined probability density functions in chemical tracer profiles and tracer-tracer correlations in the upper troposphere and lower stratosphere region and compared the results with AIRS and MLS satellite observations. We developed a simple index possibly indicating horizontal exchange between the tropical troposphere and mid-latitude stratosphere and examined its interannual variation.

## Model and Data

- MOZART-3.1 driven by fvGCM 1998
- Calculation of the thermal tropopause height from fvGCM 1998 temperature data (Reichler et al., 2003)
- Satellites: AIRS and MLS

	AIRS v. 5 (2003-2007, 5 years)	MLS v. 2.2 (2004-2008, 5 years)
O3	28 levels (1100 ~ 0.1 hPa)	12 levels (215.4 ~ 3.16 hPa)
CO	12 levels (215.4 ~ 3.16 hPa)	12 levels (215.4 ~ 3.16 hPa)
H2O	14 levels (1049 ~ 60 hPa)	18 levels (215.4 ~ 3.16 hPa)
Tropopause	WMO (1992) criteria	WMO definition, inferred from MLS and GEOS-5 temperatures

- WOUDC ozonesonde (January 1994 to March 2006)

## Thermal-tropopause relative height

- Clear separation in vertical profiles according to the tropopause heights

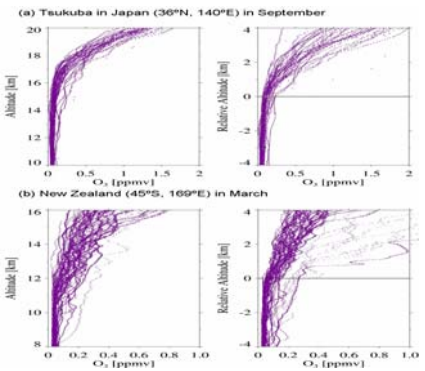


Figure 1. O<sub>3</sub> profiles from WMO ozonesondes observed (a) near Taukuba in Japan (36°N, 140°E) in every September between 1994 and 2005, and (b) in New Zealand (45°S, 169°E) in every March. Profiles are shown in altitude coordinates (left) and in thermal-tropopause relative altitude coordinates (right).

- Finer vertical resolution

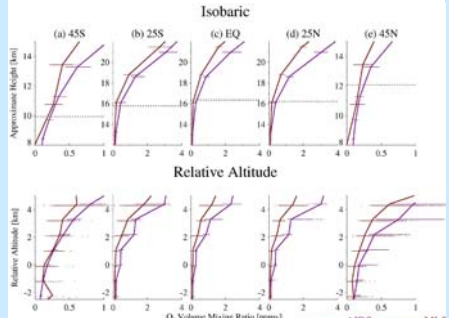


Figure 2. Zonally averaged O<sub>3</sub> vertical profiles from five center latitudes. (a) 45°S, (b) 25°S, (c) equator, (d) 25°N and (e) 45°N in October. Scattered points are values from each observed location whose latitude is closer than 5 degrees from the given center latitude. Upper panel is for isobaric vertical coordinate and lower panel is for relative altitude coordinate. Dotted lines represent zonally averaged tropopause height at each center latitude. Scale height of 70m is used to estimate the approximate height from the pressure.

## Separation of probability density functions

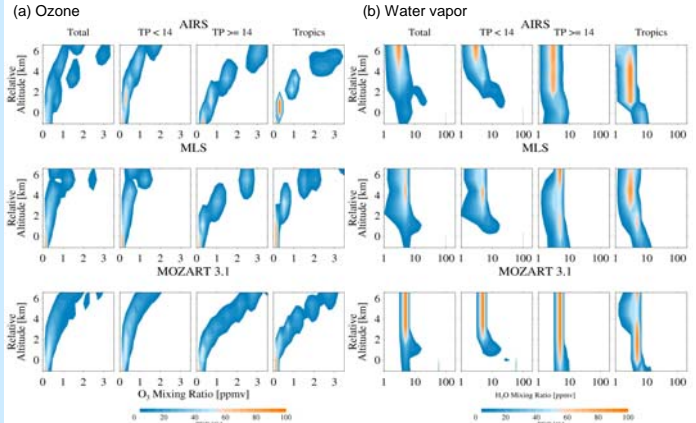


Figure 3. Vertical profiles of (a) O<sub>3</sub> and (b) H<sub>2</sub>O probability density functions (PDFs) in the northern hemisphere mid-latitude region (40°N-50°N) and the tropics (5°S-5°N) from AIRS, MLS and MOZART-3.1 result driven by fvGCM meteorological field in JAS (July, August and September). Relative height to the thermal tropopause is used as a vertical coordinate. Mid-latitude PDFs are sorted by tropopause level 14km. From left to right, total mid-latitude PDF, PDF when the tropopause height (tp hgt) is lower than 14km, PDF when tp hgt is higher than 14km and total PDF in the tropics.

O<sub>3</sub> and H<sub>2</sub>O PDFs during July, August and September are calculated and displayed in Figure 3. Vertical transition near the tropopause are clearly shown near the tropopause level. Due to the discordance between the real atmosphere and model meteorological fields, the total PDFs from MOZART-3.1 can be different from satellite observations. More importantly, when comparing the total PDFs, both of satellites and MOZART-3.1 clearly shows broader distribution with multiple peaks above 3 km level for O<sub>3</sub> and below 2 km level for H<sub>2</sub>O. The bifurcation of PDFs are related with various tropopause heights. For O<sub>3</sub> and H<sub>2</sub>O, PDFs from the higher tropopause case show similar features to the tropical PDFs. Also the PDFs for the higher tropopause cases have a less distinctive annual cycle than the total PDFs.

## Tracer-tracer correlations

The tracer-tracer correlation can be used to identify the transition region between tropospheric and stratospheric air masses (Hoor et al., 2002; Pan et al., 2004, 2007) and it varies based on its origin. Figure 4 shows MLS and MOZART O<sub>3</sub>-CO and O<sub>3</sub>-H<sub>2</sub>O relationships in the NH mid-latitude and tropics during July to September. The transition characteristics of both species form the shape of correlation map. Mid-latitude O<sub>3</sub>-CO correlation for high tropopause cases are similar to those in the tropics. This means that higher tropopause occurrences at mid-latitudes are related to air transport from the tropics.

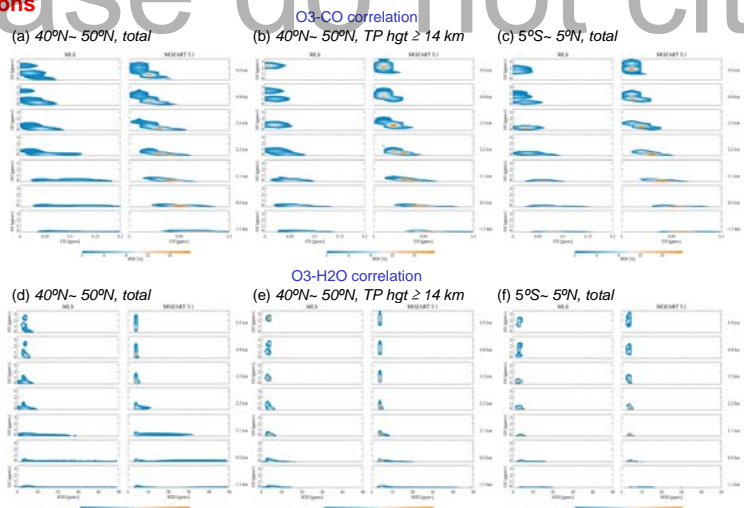


Figure 4. Two dimensional probability function showing the distribution of CO and O<sub>3</sub> [(a) - (c)] for MLS and MOZART-3.1 result at relative altitude surface from -1.1 km to 5.5 km (from bottom to top) (a) in the NH mid-latitude (40°N-50°N), (b) for higher tp hgt (>=14km) and (c) in the tropics from July through September, (d) - (f) are same as (a) - (c) but for H<sub>2</sub>O and O<sub>3</sub>.

## Frequency of the higher tropopause as an index

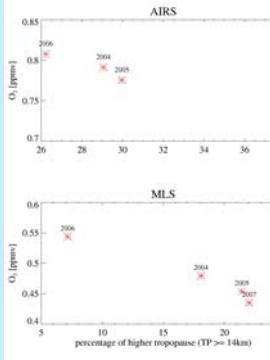


Figure 5. Scatter plot of higher tropopause index and zonal mean O<sub>3</sub> at 100hPa from AIRS (upper) and MLS (lower) in 45°N in September.

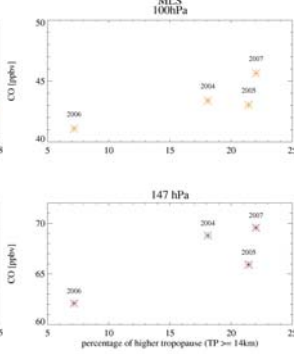


Figure 6. Scatter plot of higher tropopause index and zonal mean CO at 100hPa from AIRS (upper) and 147 hPa from MLS in 45°N in September.

For O<sub>3</sub> and CO at mid-latitudes, there are actual distinguishable changes in vertical profiles not only in the relative altitude but also in the isobaric coordinate, by the changes of tropopause heights. We made simple index indicating the tropical air intrusion in 40° ~ 50°N each month.

$$\text{Index (\%)} = \frac{\text{number of higher tropopause } (\geq 14\text{km}) \text{ observed}}{\text{number of total observation}} \times 100$$

This simple index is related to observed O<sub>3</sub> and CO in the UT/LS region. When there are more high tropopause occurrences at mid-latitudes, there is a decrease in O<sub>3</sub> and an increase in CO at around 100hPa. However, our demonstration is limited by the amount of data. We are conducting modeling studies.

## Acknowledgements

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