



Influence of the Mt. Pinatubo Eruption on the Quasi-Biennial Oscillation

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- Naujokat (1986)
 - Mt. Agung, March 1963: unusually long westerly phase of 18 months, but also the following cycle has a long westerly phase.
 - El Chichon, April 1982: relatively short westerly phase of 10 months.
 - ✤ Both cases: E jet at 25 hPa above terminating westerly phase
- Seol and Yamazaki (1998) Intensified tropical upward mass flux at 100 hPa after the Mt. Pinatubo eruption in June 1991.
- Deushi and Shibata (P20/Friday 13.00)
 Effects of the eruption of Mt. Pinatubo on the quasibiennial oscillation as revealed with MRI chemistryclimate model







1. Zonally averaged sulfate aerosol optical depth (AOD) at 0.55 µm from June 1991 to May 1993 (Stenchkov, priv. comm.)

AOD=~0.2 from 09/91 - 04/93

2. QBO in zonal wind at Singapore from January 1991 to December 1993 (Naujokat, priv. comm.)

High tropical *aerosol loading* concurrent with lifting and strengthening of the westerly phase in winter 1991/1992.



How different was this cycle from the average cycle?

30

20

-20

16

14

12

10





 Shaded: Composite average of 11 QBO cycles with easterlies starting at 45 hPa in

May/June/July of 1953-2005

Red cont.: Cycle of 1991-1993

is delayed by appr. 4 – 6 months compared to composite

4. Shaded: Composite standard deviation $\sigma(U)$

Black cont .: Difference between

u(1991-1993) and U

(ct. int.=8m/s)

Violet cont.: $|Diff| = 2\sigma$

The QBO cycle after June 1991 is an outlier wrt. the available data



Hypothesis



- Mt. Pinatubo eruption \rightarrow SO₂ in stratosphere
 - > sulfate aerosols forming in the tropical stratosphere
 - Radiative heating in near IR and thermal IR in tropics
 - Increased tropical upwelling / stronger BDC
 - > Upward advection of QBO jets
 - Delayed QBO evolution
- Secondary effects:
 - Modified wave sources in troposphere? (suppressed in this study)
- Removal of stratopsheric sulfate aerosol
 - Return to normal upwelling and QBO evolution





Experimental design



- Compare 2 ensembles of QBO simulations
 - Ctrl: No volcanic forcing
 - Aer: Stratosph. sulfate aerosols following the Mt. Pinatubo eruption is prescribed by optical properties in SW and LW
 - Aer branches off from Ctrl in June 1991
 - Differences between simulations Aer and Ctrl are related to aerosol forcing in Aer, or internal variability
 - ✤ Use ensembles to increase signal to noise ratio, N=10
- Boundary conditions for SST+sea ice
 - Climatology of AMIP2 boundary conditions of 1979-1996
 - > No Mt. Pinatubo or QBO signals in SST
- Initial conditions
 - Selected from a control run to resemble QBO phase of June 1991
 - Generate ensemble members by parameter modifications in 1st month
- Model:
 - MAECHAM5 at res. T42 L90 (Giorgetta et al., 2006)



Ensemble mean QBO in Ctrl and Aer





Shaded cont.: ensemble mean U

Red cont.: ensemble std.dev. in U

Violet cont.: t test at 95% sign.

Westerly phase lasts 4.5 months longer in AER than in CTR

Lifting and strengthening



ZMAN



Tendencies dU/dt in August-October 1991





Left: Difference in U grows from July to Nov.

Middle: tropical upwelling stronger in AER

Right: dU/dt|w* is more westerly in *AER* (dU/dt|divEP and dU/dt|vs change much less)





Conclusions



- The QBO cycle following the Mt. Pinatubo eruption was slower than normal
 - The westerly lasted 4 6 months longer than on average
 - The westerly jet was lifted and strengthened
 - The following easterly phase was delayed
- Model simulations reproduce these observed features inn response to the prescribed Mt. Pinatubo aerosol focing.
- Mechanism:
 - Volcanic aerosol radiative heating in tropical stratosphere
 - Stronger upwelling in tropical stratosphere
 - Increased upward advection of westerly wind above westerly jet
 - Delayed downward propagation of easterly jet
 - Prolonged westerly jet in lower stratosphere





- What happened in years following tropical volcanic eruptions larger than that of Mt. Pinatubo, for instance after Krakatau or Toba or the eruption of 1258 of unknown origin?
- Are these results relevant for future climate change or strat. geo-engineering schemes?
 - Common: expected changes in tropical upwelling
 - Differences:
 - Climate change: rad. forcing + wave mean-flow inteaction
 - ◆ Strat. sulfate aerosol schemes: imitate volcanic effects
 - Both: longer time scale





END