



# A Regional Desert Dust Forecast and Assimilation System

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# Contents

This presentation covers the following areas

- Observation
  - SEVIRI dust optical thickness
  - comparison with MODIS and AERONET
- Dust assimilation
  - preliminary results



# SEVIRI dust optical thickness

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# SEVIRI overview

*Spinning Enhanced Visible and Infrared Imager*

Channel no.		Characteristics of spectral band ( $\mu\text{m}$ )			Main gaseous absorber or window
		$\lambda_{\text{cen}}$	$\lambda_{\text{min}}$	$\lambda_{\text{max}}$	
1	VIS0.6	0.635	0.56	0.71	Window
2	VIS0.8	0.81	0.74	0.88	Window
3	NIR1.6	1.64	1.50	1.78	Window
4	IR3.9	3.90	3.48	4.36	Window
5	WV6.2	6.25	5.35	7.15	Water vapor
6	WV7.3	7.35	6.85	7.85	Water vapor
7	IR8.7	8.70	8.30	9.10	Window
8	IR9.7	9.66	9.38	9.94	Ozone
9	IR10.8	10.80	9.80	11.80	Window
10	IR12.0	12.00	11.00	13.00	Window
11	IR13.4	13.40	12.40	14.40	Carbon dioxide
12	HRV	Broadband (about 0.4 – 1.1)			Window/water vapor

Onboard **MSG**  
 Sampling frequency  
 15 min  
 Spatial resolution (ex HRV)  
 3km @nadir

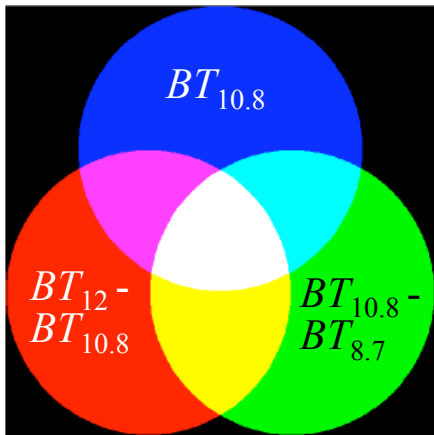
Dust/Aerosol



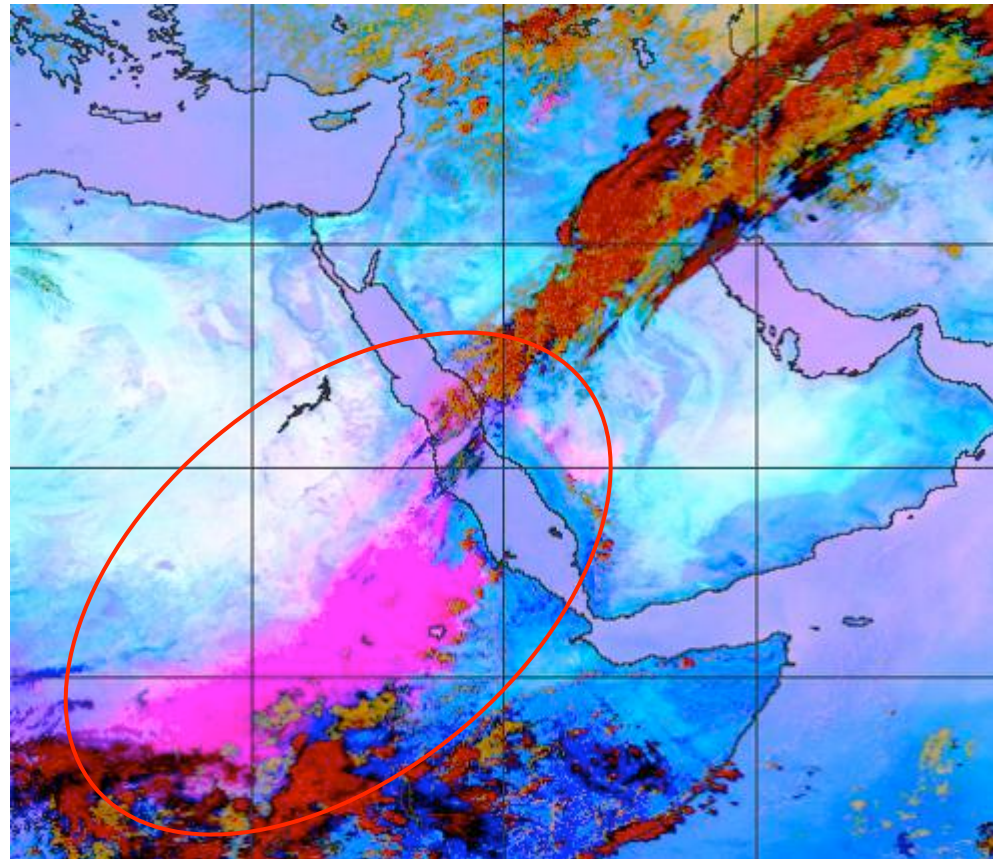


# Dust detection using thermal channels (1)

Dust RGB



Dust  
Convective cells  
anvils  
Low clouds  
Thin mid/low clouds



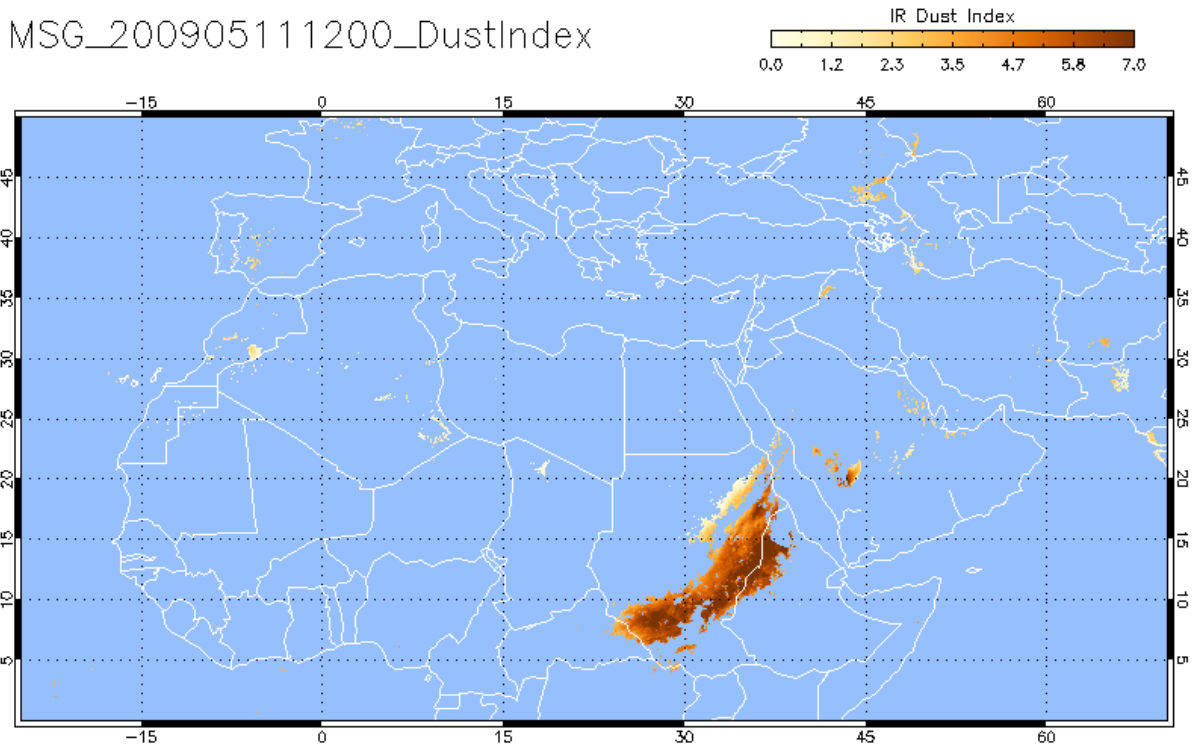
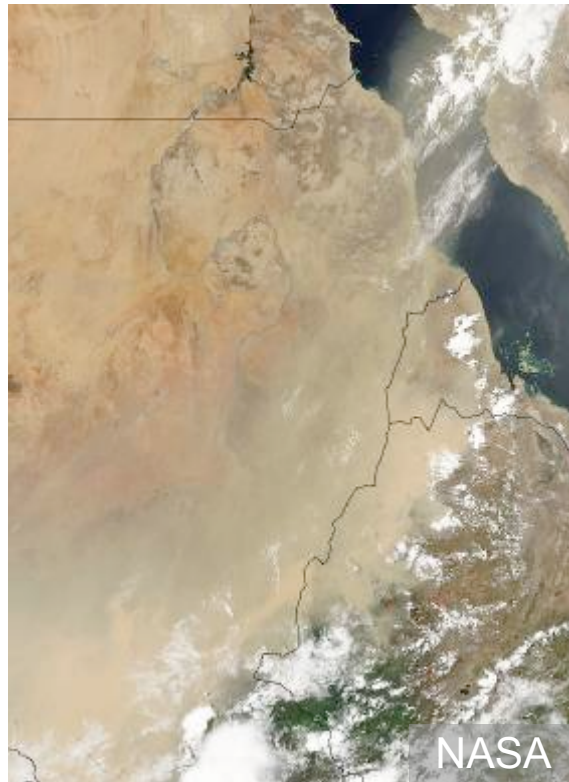
Courtesy: Pete Francis



# Dust detection using thermal channels (2)

IR Dust index

MODIS true colour MSG\_200905111200\_DustIndex



$$DustIndex = (BT_{10.8}^{forecast} - BT_{10.8}^{observed}) - (BT_{12.0}^{forecast} - BT_{12.0}^{observed})$$



# Quantification of AOD over land

- Based on the IDDI method of [Legrand \*et al.\* \(2001\)](#)
- relies on dust's strong absorption in the 10.8 $\mu$ m channel
- uses the dust optical properties described in [Volz \(1973\)](#).
- AOD and the change in  $BT_{10.8}$  can be expressed as ([Brindley and Russell, 2009](#))

$$AOD_{550} = c_1(\theta_v) \cdot c_2(\theta_v)^{\Delta BT_{10.8}} + c_3(\theta_v)$$

$c_{1,2,3}$  fitted coefficients as a function of viewing zenith angle  $\theta_v$

**Simulation criteria** AOD: 0-5, Surface type: (semi) arid,  
TCWV: 2-20kg/m<sup>2</sup>, Ts: 300-330K,  $\theta_v$ : 0-70°



# Quantification of AOD over land – application to SEVIRI

- Met Office operational cloud mask with dust-flags was used to screen out cloud-only pixels
- A 16-day window is considered to obtain a reference (clear-sky)  $BT_{10.8}$ 
  - corrections were applied to account for the change due to the change in meteorology
- $\Delta BT_{10.8}$  was then obtained by subtracting observed BT from the reference BT





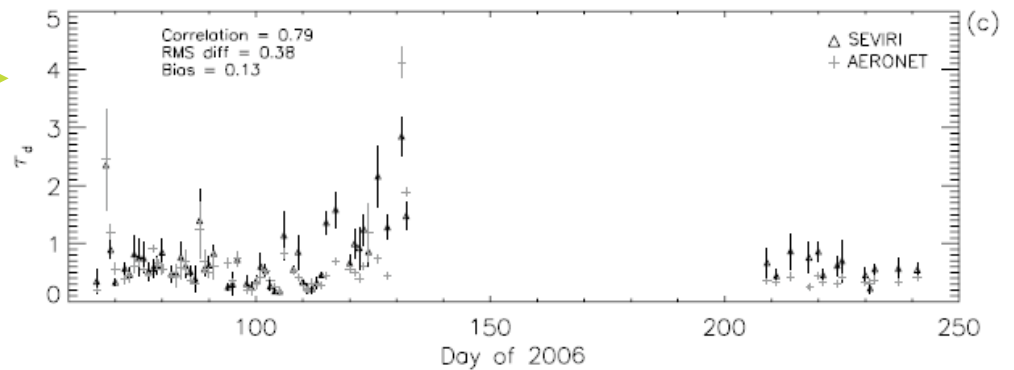
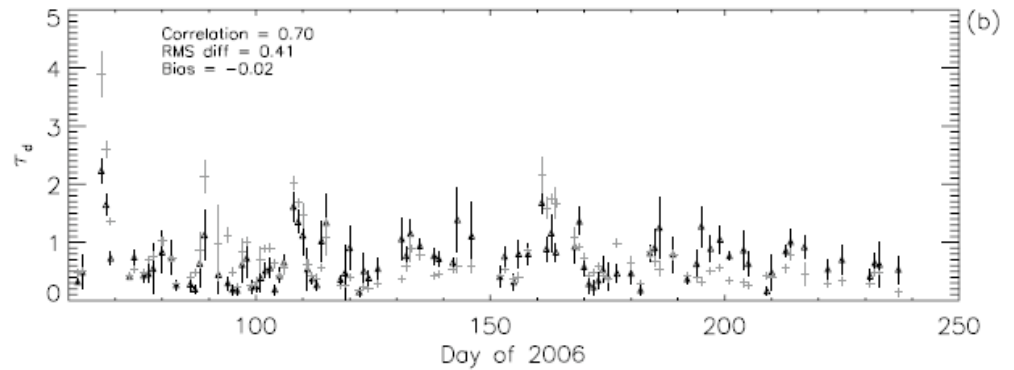
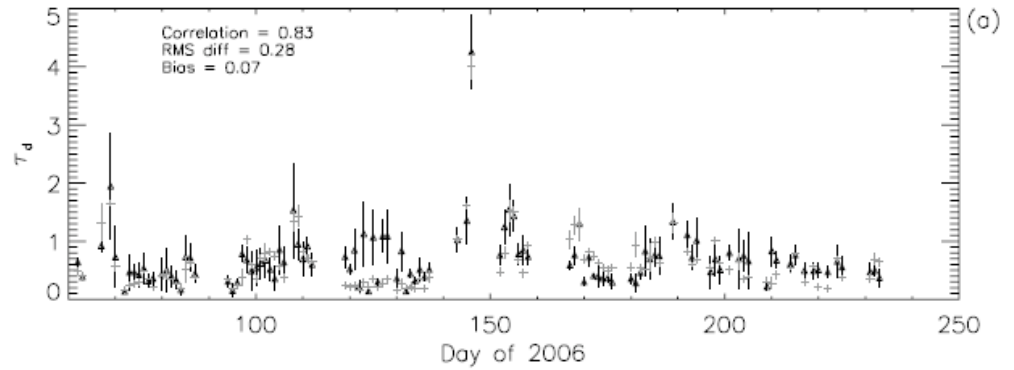
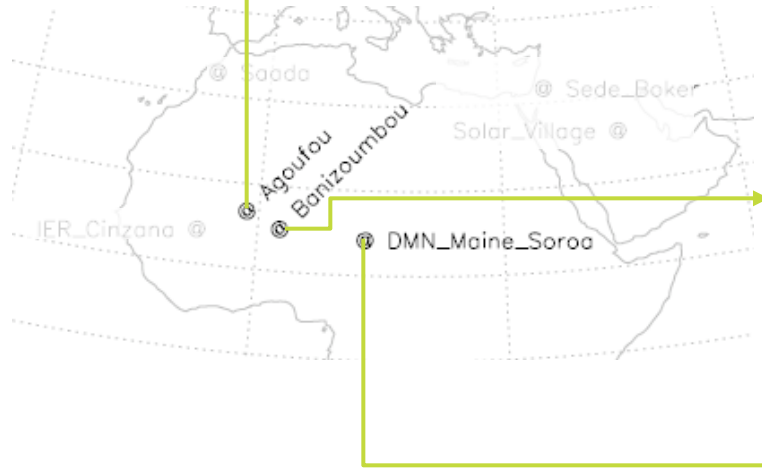
# Limitations



- relies on accurate discrimination of dust from cloud
- some level of background dust loading is always present in reality
- signals from the dust layer are difficult to separate when very close to the ground
- may have a reverse effect on BT during non-sunlit period



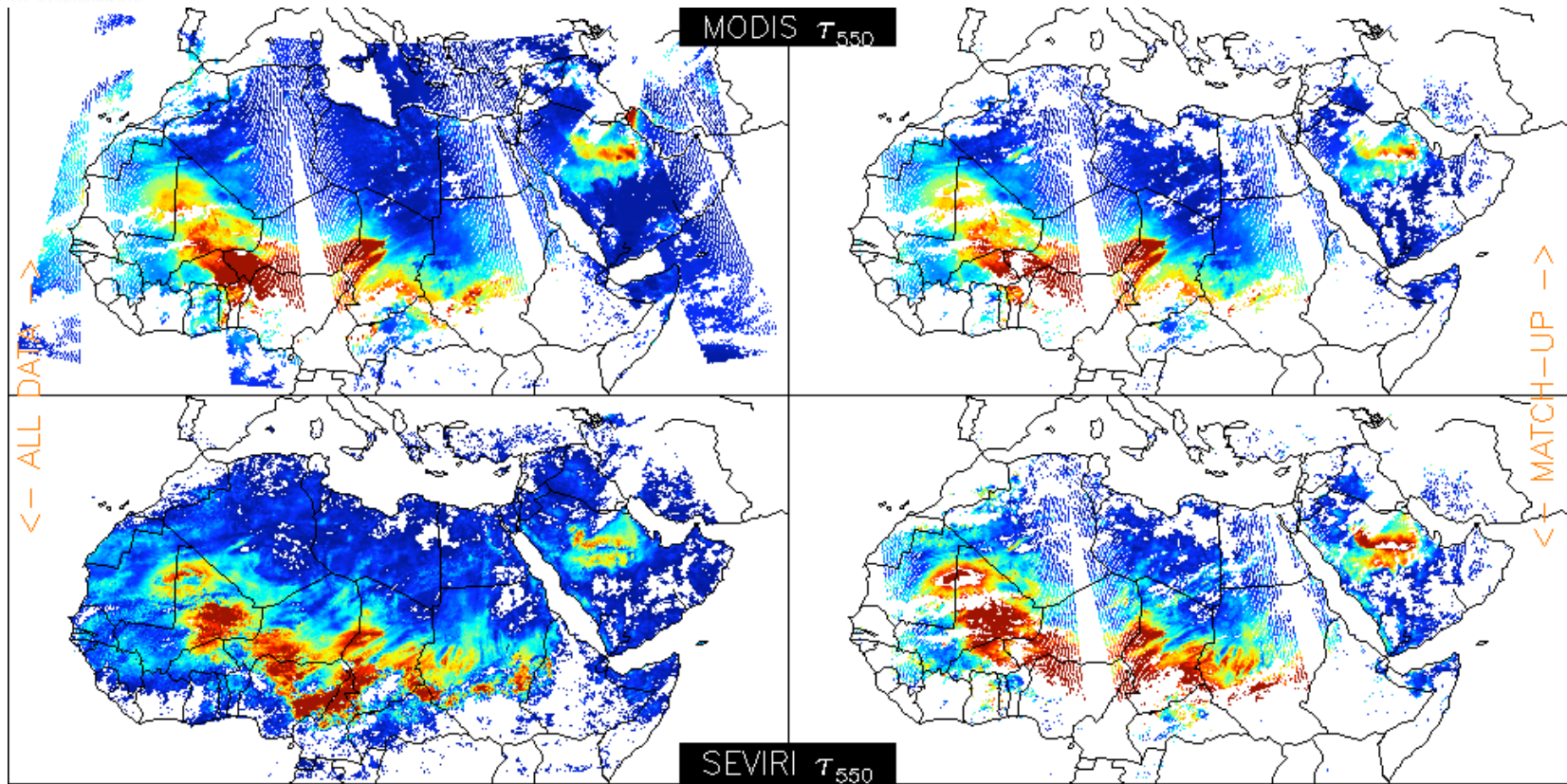
# Comparison with AERONET



Comparison with data from 7 AERONET locations show an overall rms difference of 0.37 and bias 0.01

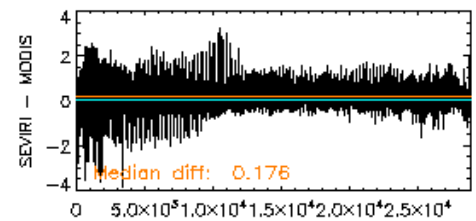
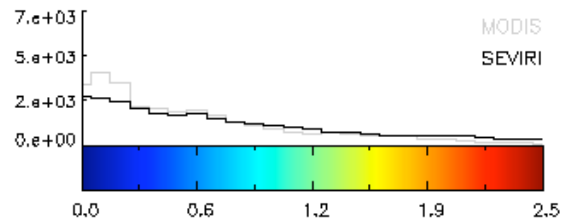
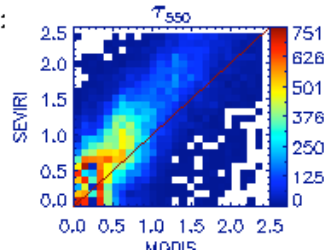


# Comparison with Aqua/MODIS daily, near real time



19032010 Stats:

Slope 1.07  
 Intcpt 0.16  
 bias 0.21  
 rms 0.57  
 R 0.81  
 N 28917



Median diff: 0.176



# Summary

- semi-operational hourly SEVIRI  $\tau_{550}$  at near-real-time with a pre-assessed rms  $\pm 0.37$
- capable to detect and quantify major dust events

## Scope for improvement

- dust detection (very thick layers) needs further improvement
- address the impact of dust layer height on the retrieval



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# Dust assimilation

K. Ngan, Y. Pradhan, M. Brooks, D. Walters & D. Jackson

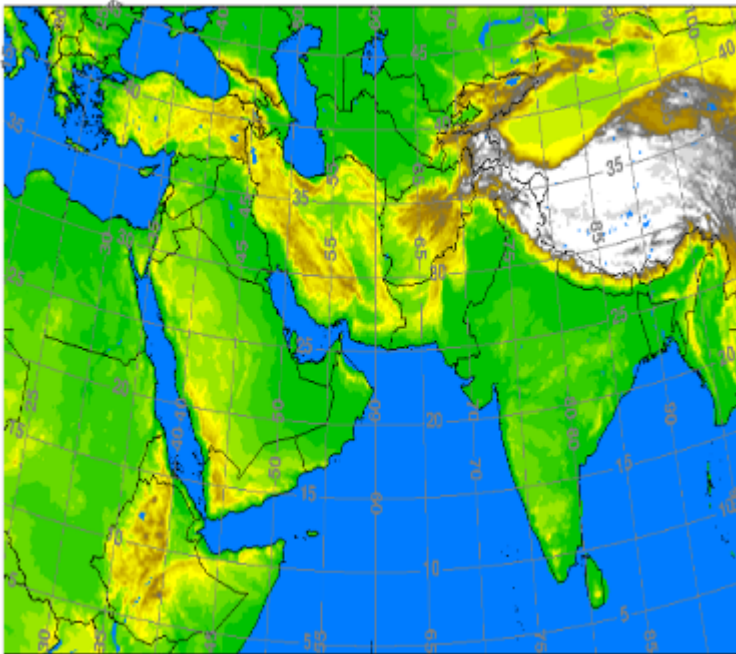


# Aerosol / dust forecasting

- Numerical Model
- Dust scheme
- Observations
- Assimilation trials



# South Asia Model (SAM)

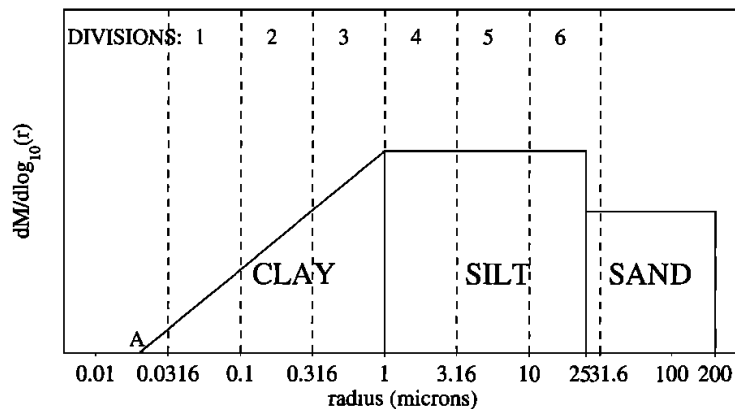


- Limited area NWP configuration of MetUM introduced in Apr 2008
- Grid 648x400x70 (Mar 2010)
- Horizontal Resolution 12 km
- Vertical lid 80 km
- LBCs and covariance statistics are derived from the global model



# Dust scheme

Since dust particles vary widely in size, a **discrete representation** is necessary. The Met Office scheme (Woodward 2001) uses 6 bins.



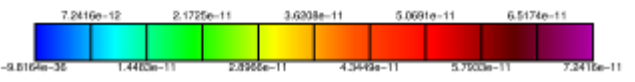
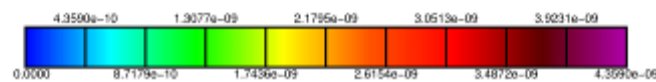
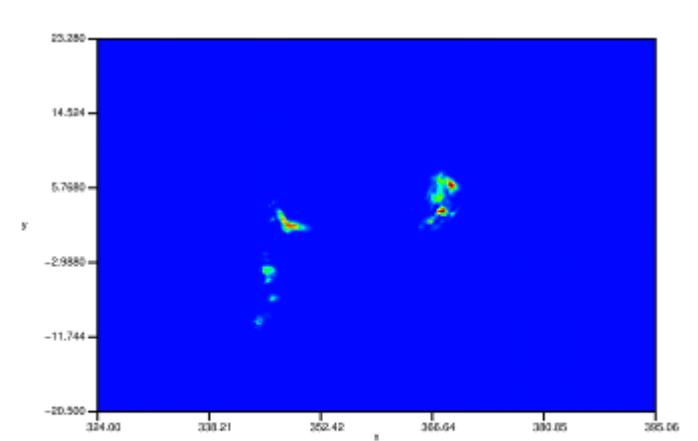
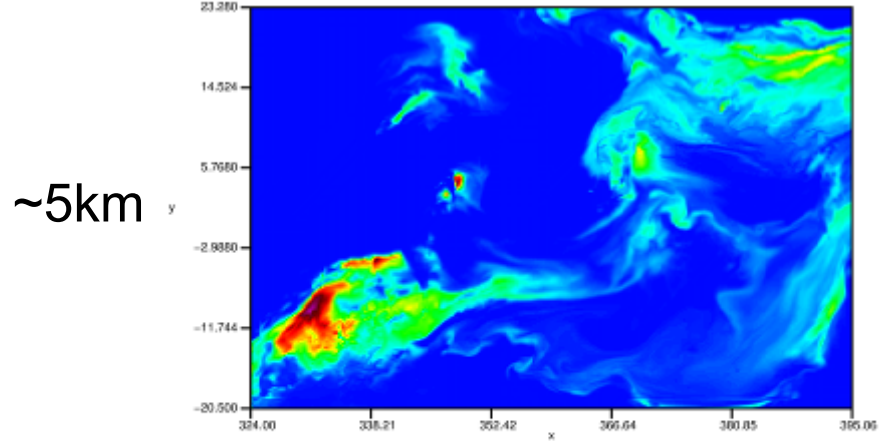
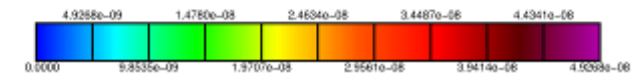
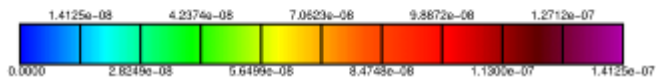
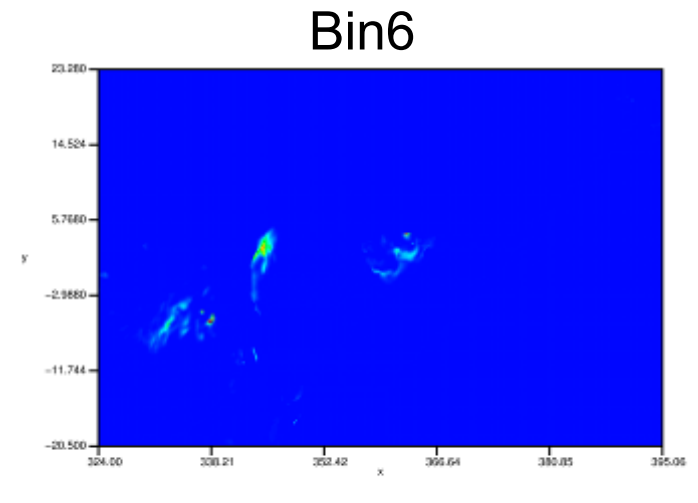
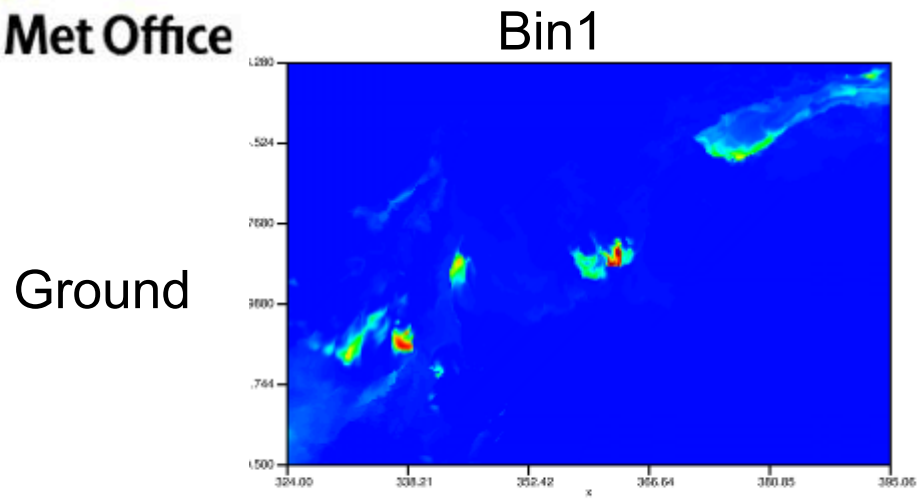
Division number	1	2	3	4	5	6
Lower bound ( $\mu\text{m}$ )	0.0316	0.1	0.316	1.0	3.16	10.0
Upper bound ( $\mu\text{m}$ )	0.1	0.316	1.0	3.16	10.0	31.6

- Emission is parameterised using density and relative mass of each size bin.
- Both wet and dry deposition are included.
- Direct radiative effect is included





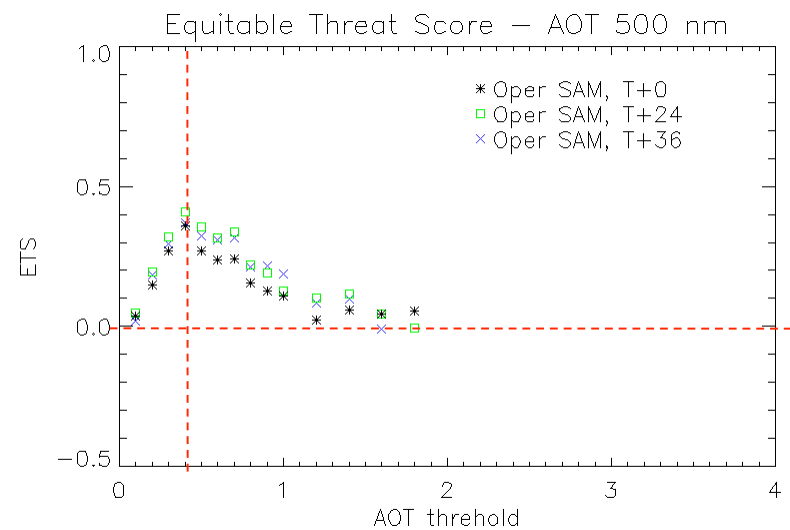
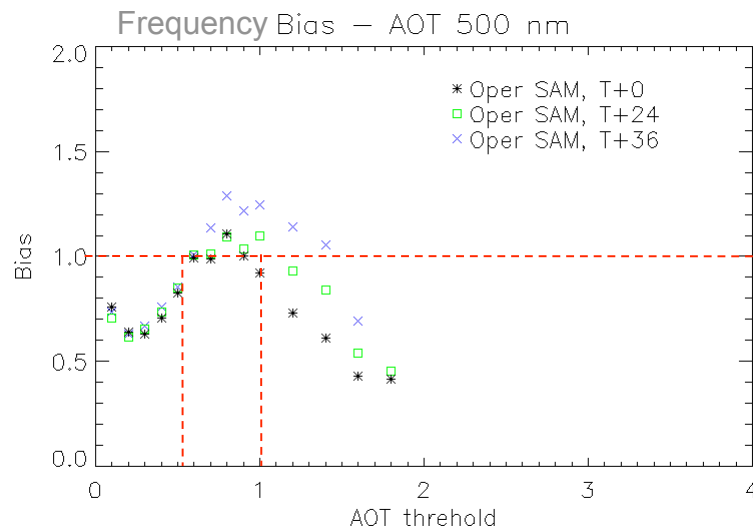
# Model fields (dust mixing ratio)





# SAM forecast skill

March 2009-March 2010 (covering 10 AERONET locations over Middle-east and N. Africa)



## SAM forecasts against AERONET (no assimilation)

- under-predict AOD at low thresholds (probably due to the absence of non-dust aerosols)
- relatively unbiased for  $0.5 < \text{AOD} < 1$
- most difficult to predict for larger AOD events



# DA approach

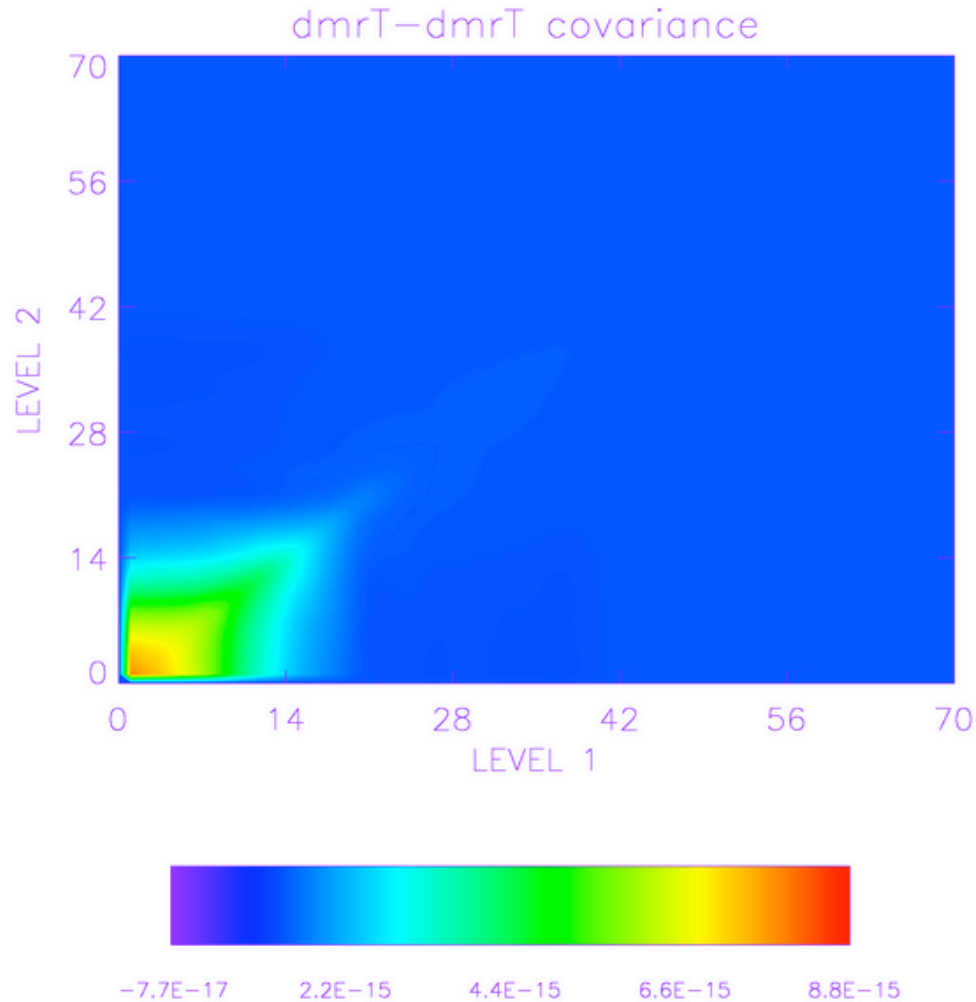
Work within framework of existing incremental 3D-Var system.

## Key elements

- Background statistics obtained via the NMC method
- Observation variable is **aerosol optical depth**
  - Obs err matrix (**R**) is assumed to be diagonal (rms error =0.37 from comparisons of SEVIRI with AERONET)
- Control variable is **total dust mixing ratio** (after [Benedetti et al. 2009](#)). Total dust is obtained by summing over all dust bins.



# Background error covariance

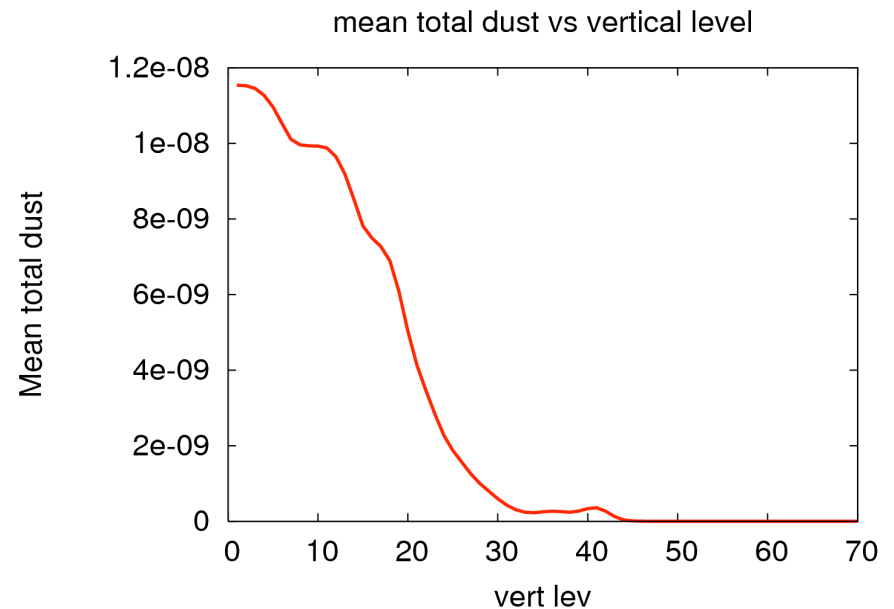
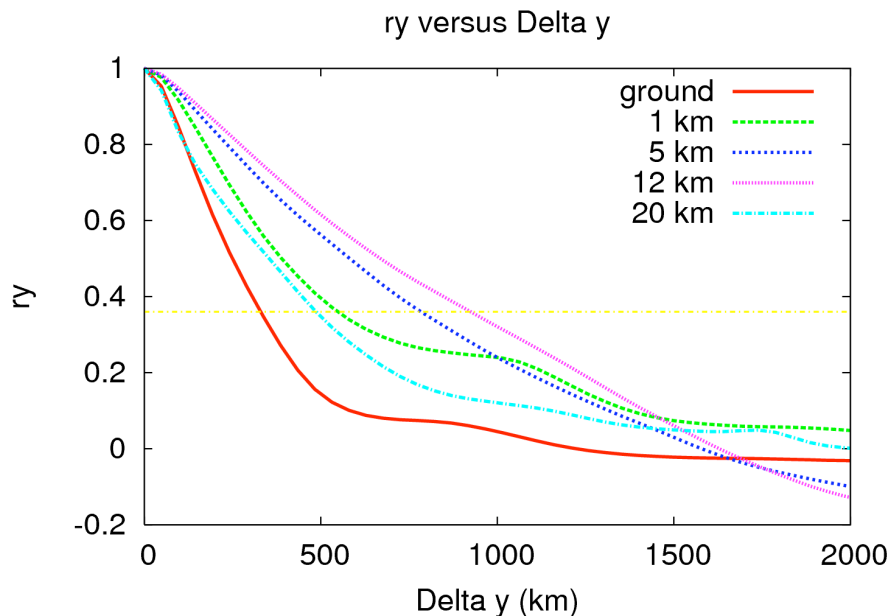


Vertical correlations extend throughout boundary layer (level 14 ~ 1.5 km).



# Spatial autocorrelations

**Met Office** In limited-area mode, the Met Office variational system requires that a horizontal correlation length be prescribed for the control variable, i.e., total dust.

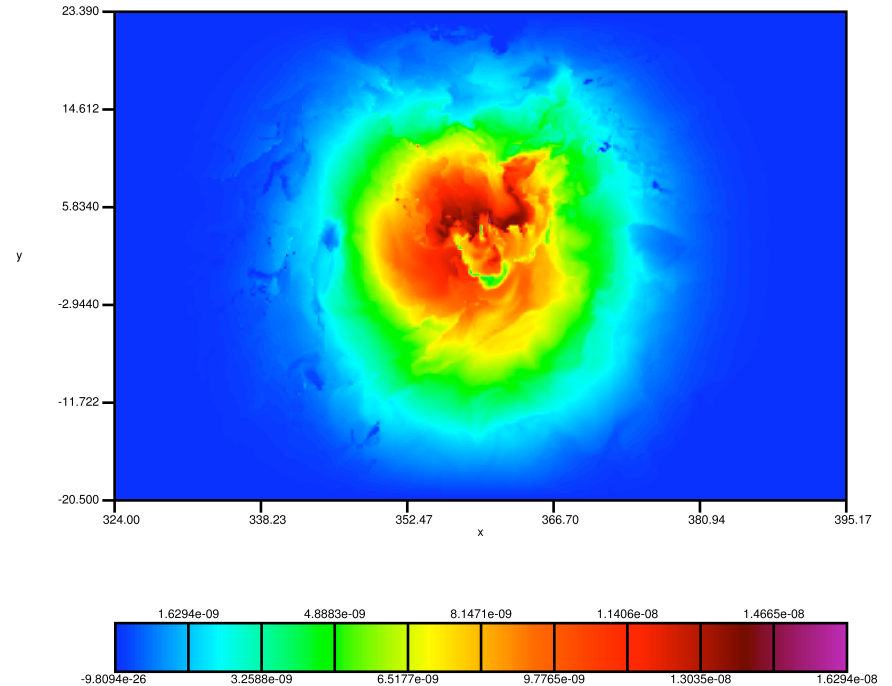
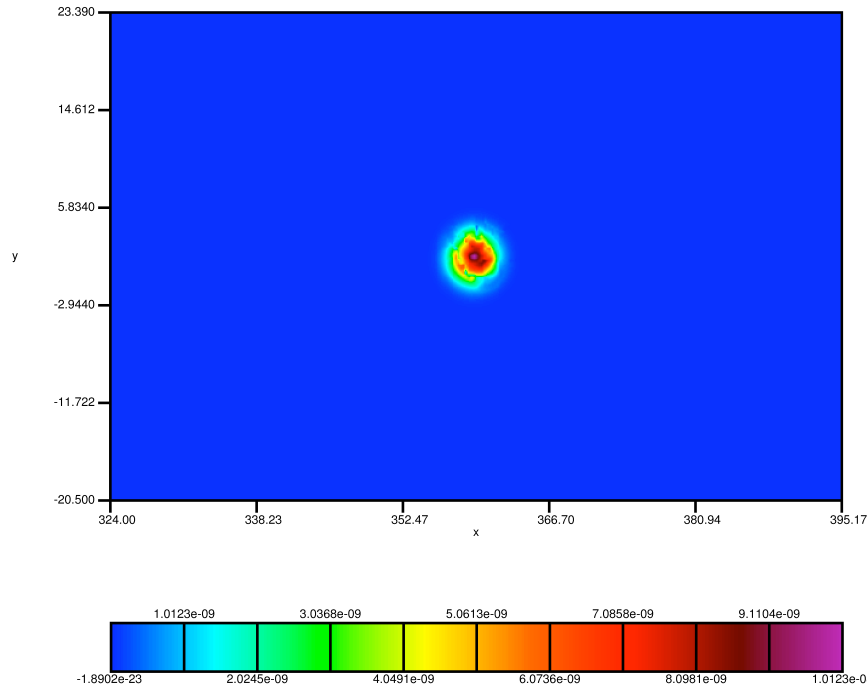


- Correlation lengthscale cannot be specified unambiguously.
- We use  $L_{td} = 300\text{km}$ . By default,  $L_{\psi} = 400\text{km}$  in the global model.



# Pseudo-observations

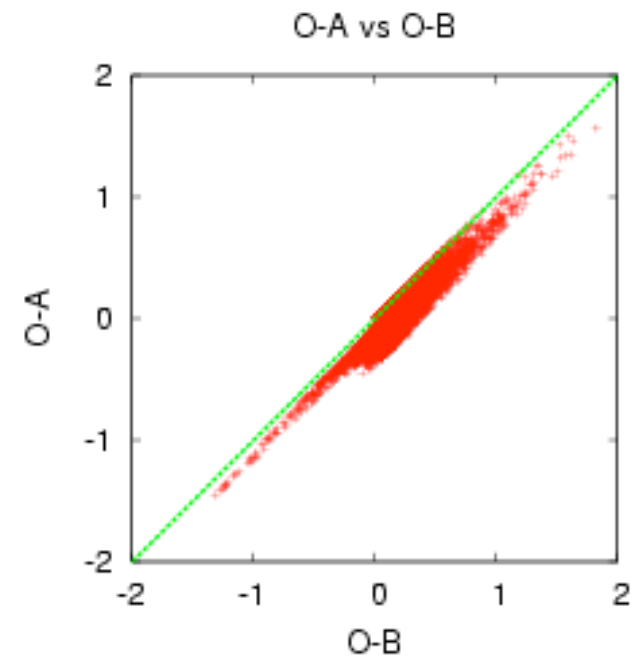
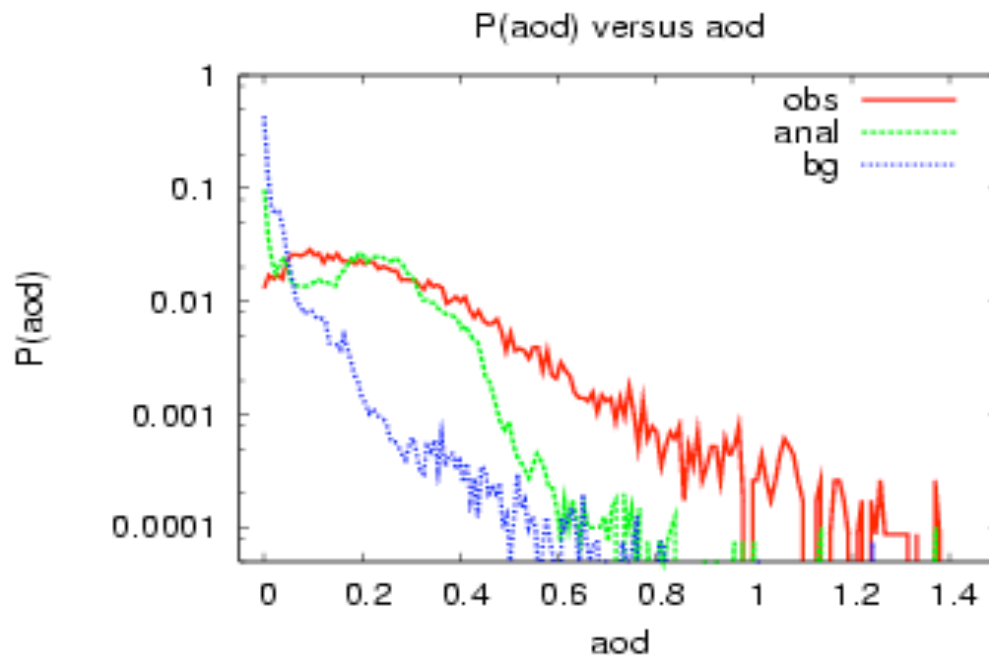
Increment to dust bin 1 at ground



- Horizontal structure is a product of the isotropic horizontal correlations and the anisotropic background.

# Single-cycle Analysis and Background

23-Jan-2010



- Analysed pdf is intermediate between analysed and observed pdfs.
- fails to capture long tail

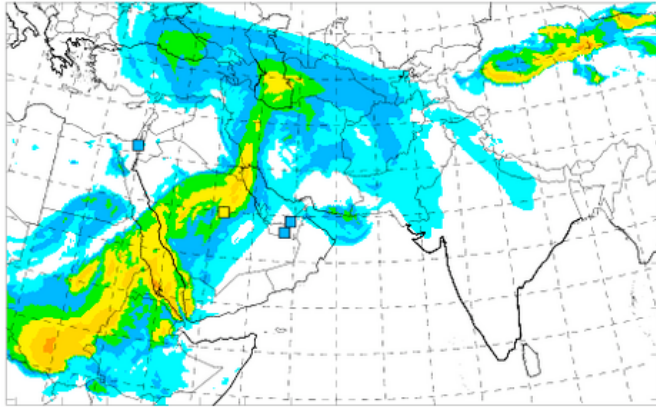
- Analysis error is smaller than background error for  $O-B > 0$



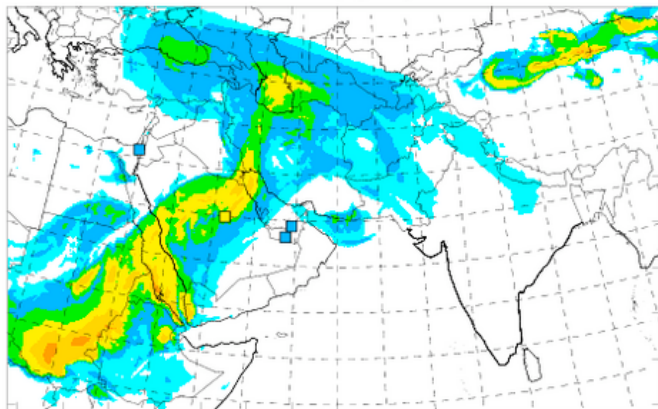
# Trial results (1)

Trial period: 10 Jan'10 – 10 Feb'10

Forecast mineral dust AOT 550nm with Aeronet obs  
PS23 Setup (sgotp): 20100119 12:00



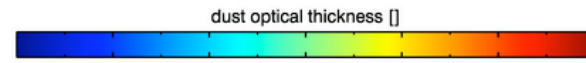
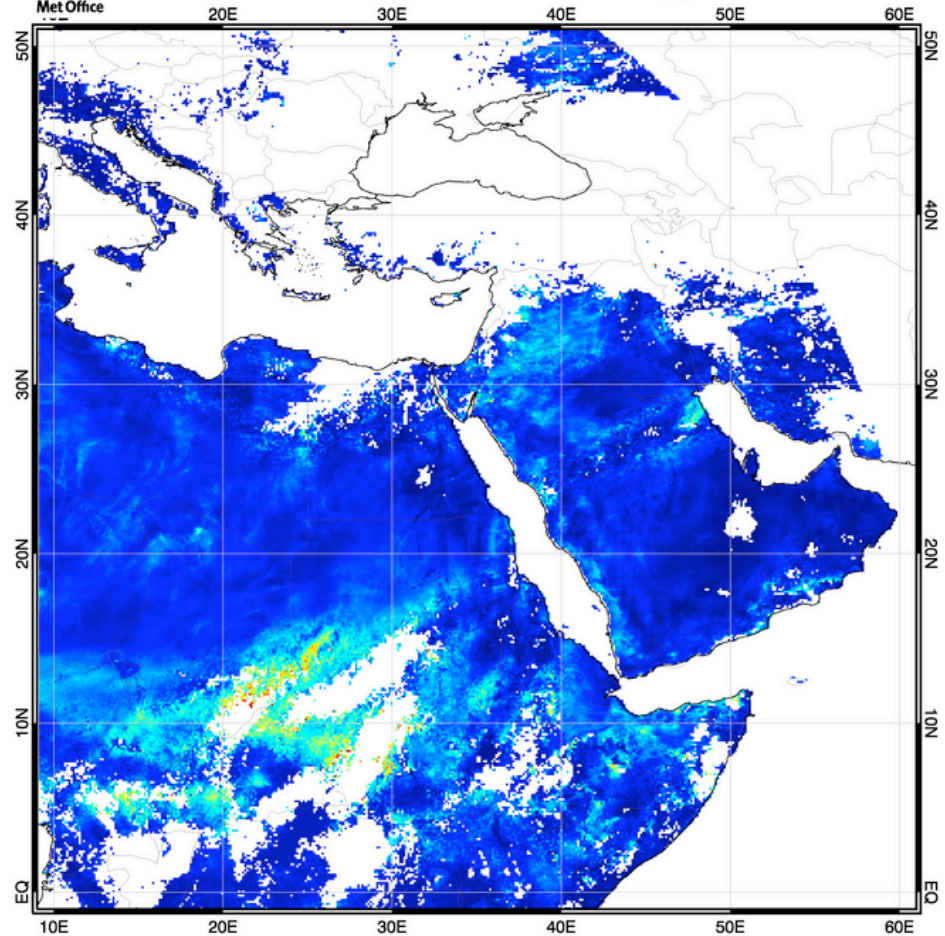
Forecast mineral dust AOT 550nm with Aeronet obs  
PS23 with SEVIRI assim. (sgote): 20100119 12:00



© C



MSG\_20100119\_DAILY\_A ( $\tau_{550}$ )

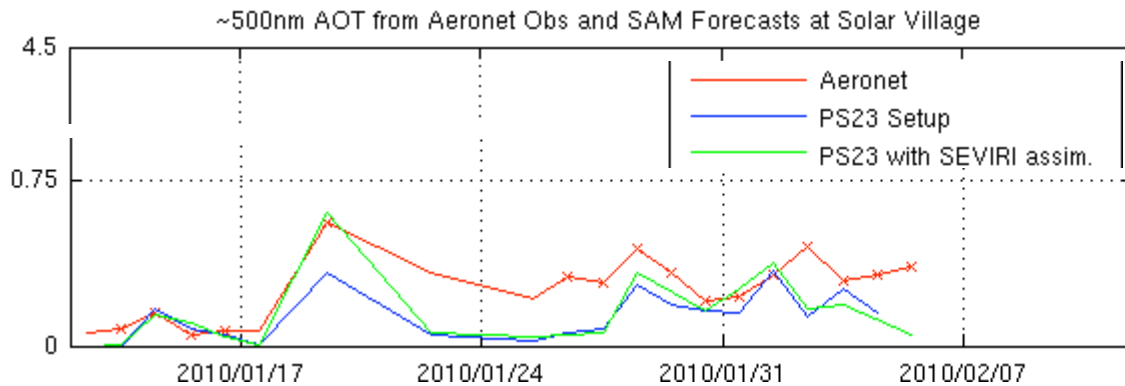


© Met Office/EUMETSAT 0.0 0.3 0.7 1.0 1.3 1.7 2.0 Mon Mar 29 13:38:48 2010

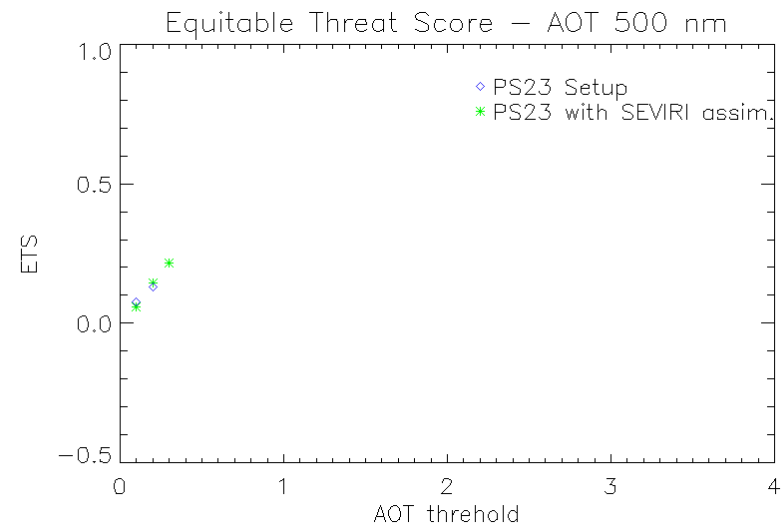
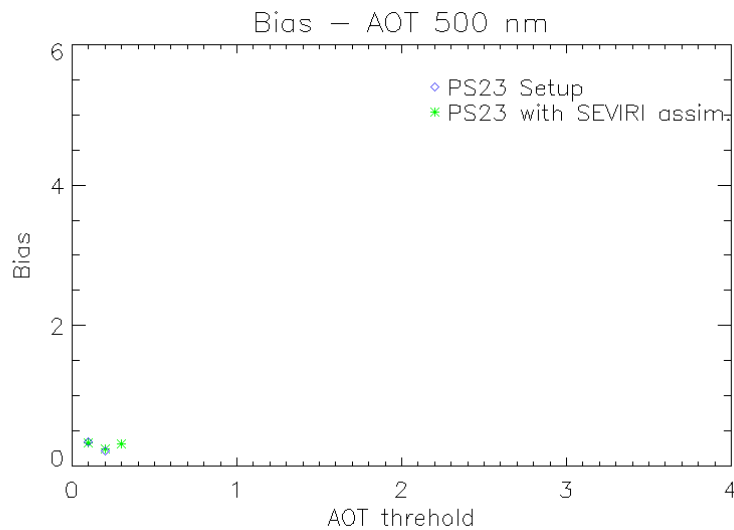




# Trial results (2)



Captured the dust event over Solar Village; Timeseries over other AERONET locations shows little or no improvement





# Future Work

## Short term

- 4d-Var
- More extensive trialling: major dust events

## Longer term

- Alternative control variables
- Alternative data sources (e.g. MODIS/VIIRS)
- Volcanic ash assimilation (?)



Met Office



# Questions and answers



# Control variable and adjoint

Variational systems minimise the control-variable cost function using **adjoint equations**.

- Assumptions (Benedetti et al. 2009):
  - Total dust mixing ratio,  $r_T = \text{constant}$
  - Fractional mixing ratio  $f_i = r_i/r_T = \text{constant}$
- Tangent-linear equations

$$\begin{pmatrix} \tilde{r}'_T \\ \tilde{r}'_i \end{pmatrix} = \begin{pmatrix} 1 & 0 \\ f_i & 0 \end{pmatrix} \begin{pmatrix} r'_T \\ r'_i \end{pmatrix}$$

- Adjoint equations follow immediately. For real variables, the adjoint operator is equivalent to the **transpose**.